A GUIDE
TO THE
FOSSIL INVERTEBRATE ANIMALS
IN THE DEPARTMENT OF
GEOLOGY AND PALAEONTOLOGY
IN THE
BRITISH MUSEUM (NATURAL HISTORY),
CROMWELL ROAD, LONDON, S.W.

WITH 7 PLATES AND 96 TEXT-FIGURES.

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PLAN OF GALLERIES CONTAINING FOSSIL INVERTEBRATE ANIMALS.
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TABLE OF CONTENTS.

<table>
<thead>
<tr>
<th>Section</th>
<th>Page</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table of Contents</td>
<td>iii</td>
</tr>
<tr>
<td>List of Plates</td>
<td>vi</td>
</tr>
<tr>
<td>List of Illustrations in Text</td>
<td>vi</td>
</tr>
<tr>
<td>Preface</td>
<td>ix</td>
</tr>
<tr>
<td>Geological Time-scale</td>
<td></td>
</tr>
<tr>
<td>Opposite p.</td>
<td>1</td>
</tr>
<tr>
<td>Plan and Principles of Arrangement</td>
<td></td>
</tr>
<tr>
<td>Gallery XI. Stratiographical Series</td>
<td>4</td>
</tr>
<tr>
<td>Historical Collections</td>
<td>6</td>
</tr>
<tr>
<td>Dynamical Series</td>
<td>12</td>
</tr>
<tr>
<td>Tracks and Markings</td>
<td>14</td>
</tr>
<tr>
<td>Gallery X. Protozoa</td>
<td>14</td>
</tr>
<tr>
<td>Class Foraminifera</td>
<td>16</td>
</tr>
<tr>
<td>Eozoont</td>
<td>25</td>
</tr>
<tr>
<td>Class Radiolaria</td>
<td>26</td>
</tr>
<tr>
<td>Porifera (Sponges)</td>
<td>29</td>
</tr>
<tr>
<td>Class Calcarea</td>
<td>31</td>
</tr>
<tr>
<td>Class Hexactinellida</td>
<td>32</td>
</tr>
<tr>
<td>Class Demospongiae</td>
<td>35</td>
</tr>
<tr>
<td>Class (?) Octactinellida</td>
<td>38</td>
</tr>
<tr>
<td>Class (?) Heteractinellida</td>
<td>39</td>
</tr>
<tr>
<td>Coelentera</td>
<td>43</td>
</tr>
<tr>
<td>Class Scyphozoaa</td>
<td>44</td>
</tr>
<tr>
<td>Class Hydrozoa</td>
<td>45</td>
</tr>
<tr>
<td>Class Anthozoa</td>
<td>48</td>
</tr>
<tr>
<td>Gallery VIII. Echinodermia</td>
<td>58</td>
</tr>
<tr>
<td>Class Crinoidea</td>
<td>58</td>
</tr>
<tr>
<td>Class Cystidea</td>
<td>63</td>
</tr>
<tr>
<td>Class Blastoida</td>
<td>66</td>
</tr>
<tr>
<td>Class Edrioasteroida</td>
<td>67</td>
</tr>
<tr>
<td>Class Asteroidea</td>
<td>68</td>
</tr>
<tr>
<td>Class Ophiuroidea</td>
<td>70</td>
</tr>
<tr>
<td>Class Echinoidea</td>
<td>71</td>
</tr>
<tr>
<td>Class Holothurioidea</td>
<td>76</td>
</tr>
<tr>
<td>Echinoderms as Rock-formers</td>
<td>76</td>
</tr>
<tr>
<td>TABLE OF CONTENTS.</td>
<td></td>
</tr>
<tr>
<td>-------------------</td>
<td></td>
</tr>
<tr>
<td>Gallery VIII. ANNELIDA .</td>
<td>76</td>
</tr>
<tr>
<td>Class CHAETOPODA .</td>
<td>76</td>
</tr>
<tr>
<td>ARTHROPODA .</td>
<td>81</td>
</tr>
<tr>
<td>Class TRILOBITA .</td>
<td>82</td>
</tr>
<tr>
<td>Class ARACHNIDA .</td>
<td>86</td>
</tr>
<tr>
<td>Order EURYPTERIDA .</td>
<td>86</td>
</tr>
<tr>
<td>Order XIPHOSURA .</td>
<td>88</td>
</tr>
<tr>
<td>Order SCorpionida .</td>
<td>89</td>
</tr>
<tr>
<td>Order Anthracomarti .</td>
<td>91</td>
</tr>
<tr>
<td>Order Araneae .</td>
<td>92</td>
</tr>
<tr>
<td>Class CRUSTACEA .</td>
<td>92</td>
</tr>
<tr>
<td>Sub-Class BRANCHIOPODA .</td>
<td>93</td>
</tr>
<tr>
<td>Order Phyllopoda .</td>
<td>93</td>
</tr>
<tr>
<td>Sub-Class OSTRACODA .</td>
<td>93</td>
</tr>
<tr>
<td>Sub-Class CIRRIPEDIA .</td>
<td>94</td>
</tr>
<tr>
<td>Sub-Class MALACOSTRACA .</td>
<td>95</td>
</tr>
<tr>
<td>Group Phyllocarida .</td>
<td>95</td>
</tr>
<tr>
<td>Group Syncarida .</td>
<td>95</td>
</tr>
<tr>
<td>Group Schizopoda .</td>
<td>96</td>
</tr>
<tr>
<td>Order Isopoda .</td>
<td>96</td>
</tr>
<tr>
<td>Order Amphipoda .</td>
<td>97</td>
</tr>
<tr>
<td>Order Stomatopoda .</td>
<td>97</td>
</tr>
<tr>
<td>Order Decapoda .</td>
<td>98</td>
</tr>
<tr>
<td>Sub-order Macrura .</td>
<td>98</td>
</tr>
<tr>
<td>Sub-order Brachyura .</td>
<td>100</td>
</tr>
<tr>
<td>Class DIPLOPODA .</td>
<td>103</td>
</tr>
<tr>
<td>Class CHILOPODA .</td>
<td>104</td>
</tr>
<tr>
<td>Class INSECTA .</td>
<td>104</td>
</tr>
<tr>
<td>BRACHIOPODA .</td>
<td>108</td>
</tr>
<tr>
<td>Order Atremata .</td>
<td>113</td>
</tr>
<tr>
<td>Order Neotremata .</td>
<td>113</td>
</tr>
<tr>
<td>Order Protremata .</td>
<td>113</td>
</tr>
<tr>
<td>Order Telotremata .</td>
<td>113</td>
</tr>
<tr>
<td>BRYOZOA OR POLYZOA .</td>
<td>116</td>
</tr>
<tr>
<td>Sub-Class GYMNO LAEMATA .</td>
<td>117</td>
</tr>
<tr>
<td>Order Trepostomata .</td>
<td>117</td>
</tr>
<tr>
<td>Order Cryptostomata .</td>
<td>117</td>
</tr>
<tr>
<td>Order Cyclostomata .</td>
<td>118</td>
</tr>
<tr>
<td>Order Cheilostomata .</td>
<td>118</td>
</tr>
<tr>
<td>Gallery VIII.</td>
<td>MOLLUSCA</td>
</tr>
<tr>
<td>--------------</td>
<td>------------</td>
</tr>
<tr>
<td></td>
<td>Class</td>
</tr>
<tr>
<td></td>
<td>Amphineura</td>
</tr>
<tr>
<td></td>
<td>Gastropoda</td>
</tr>
<tr>
<td></td>
<td>Scaphopoda</td>
</tr>
<tr>
<td></td>
<td>Lamellibranchia</td>
</tr>
<tr>
<td>Gallery VII.</td>
<td>Class Cephalopoda</td>
</tr>
<tr>
<td></td>
<td>Nautiloidea</td>
</tr>
<tr>
<td></td>
<td>Ammonoidea</td>
</tr>
<tr>
<td></td>
<td>Coleoidea or Belemnoidea</td>
</tr>
<tr>
<td>INDEX</td>
<td></td>
</tr>
</tbody>
</table>
LIST OF PLATES.

I. Plan of Galleries containing Fossil Invertebrate Animals. Frontispiece.

II. William Smith, LL.D., aged 69. (Lent by Dr. Henry Woodward) To face page 7

III. Siliceous skeletons of Sponges from the Upper Senonian Chalk of Hanover. To face page 42

IV. A slab of Lias shale from Boll, in Wurtemberg, covered with the remains of a large Crinoid, Pentacerinus Hiemeri. To face page 63

V. Cambrian Trilobites from China. (Lent by the Editor of the Geological Magazine) To face page 86

VI. Fossil Shells: Trigonia in Corallian rock from Weymouth; part of a shell-bank in London Clay at Fareham, Hants. To face page 135

VII. Ammonite Marble: Marston Stone, Lower Lias near Yeovil, full of Amblyoceras planicosta. To face page 167

In Plates IV, VI, and VII, the scale is given by a foot-rule divided into inches and photographed with the objects.

LIST OF ILLUSTRATIONS IN TEXT.

FIG. PAGE
1. Living Foraminifera: Miliola and Botalia 16
2. Foraminifer from the Upper Chalk: Anomalina. (From Chapman, after Eley) 17
3. Examples of Foraminifera. (From Nicholson, after Brady) 21
4. Foraminifera as Rock-formers. (From Chapman) 22
5. Two generations of a Nummulite. (After De la Harpe) 24
6. Radiolaria, Recent and Fossil. (After Haeckel and Rüst) 26
7. Radiolarian rock from the Lower Culm, Cornwall. (From Hinde and Fox) 28
8. Fossil Sponge spicules: various Silicispongiae and Calcispongiae. (After Hinde) 30
10. Reconstruction of Ventriculites. (After Minchin) 34
11. Fossil Sponge spicules: Silicispongiae, Tetractinellida Choristida. (After Hinde) 36
<table>
<thead>
<tr>
<th>FIG.</th>
<th>LIST OF ILLUSTRATIONS IN TEXT.</th>
<th>PAGE</th>
</tr>
</thead>
<tbody>
<tr>
<td>12.</td>
<td>Fossil Sponge spicules: Silicispongiae, Tetractinellida Lithistida. (After Hinde)</td>
<td>37</td>
</tr>
<tr>
<td>13.</td>
<td>Ditto</td>
<td>Ditto</td>
</tr>
<tr>
<td>14.</td>
<td>Fossil Sponge spicules: Silicispongiae, Monactinellida. (After Hinde)</td>
<td>38</td>
</tr>
<tr>
<td>15.</td>
<td>A Calcisponge, Peronidella, Great Oolite</td>
<td>39</td>
</tr>
<tr>
<td>16.</td>
<td>Lithistid sponges from Upper Greensand, Warminster. (After Hinde)</td>
<td>40</td>
</tr>
<tr>
<td>17.</td>
<td>Verruculina, Chalk of Flamborough</td>
<td>41</td>
</tr>
<tr>
<td>18.</td>
<td>A fossil jelly-fish, Rhizostomites, from Solenhofen. (After Walcott and Brandt)</td>
<td>44</td>
</tr>
<tr>
<td>19.</td>
<td>An early form of Graptolite, Didymograptus</td>
<td>46</td>
</tr>
<tr>
<td>20.</td>
<td>A Favorisitid coral, Syringolites, Silurian. (After Hinde)</td>
<td>51</td>
</tr>
<tr>
<td>21.</td>
<td>Silurian Anthozoa, possibly Aleyconaria, Wenlock Limestone: Halysites, Autopora</td>
<td>52</td>
</tr>
<tr>
<td>22.</td>
<td>Palaeozoic Anthozoa: Omphyna (from Prestwich); Syringopora</td>
<td>53</td>
</tr>
<tr>
<td>23.</td>
<td>A Devonian operculate coral: Calceola</td>
<td>54</td>
</tr>
<tr>
<td>24.</td>
<td>Zoantharian corals of Bajocien age: Latomaeandrae, Montlivaltia. (From Prestwich)</td>
<td>55</td>
</tr>
<tr>
<td>25.</td>
<td>Zoantharian corals of Upper Corallian age: Isasraea, Thecosmilia. (From Prestwich)</td>
<td>56</td>
</tr>
<tr>
<td>26.</td>
<td>Zoantharian corals from the British Chalk: Syn helia, Parasmina, Stephanop hyllia. (From Prestwich)</td>
<td>56</td>
</tr>
<tr>
<td>27.</td>
<td>Perforate Zoantharian corals from the Lutetian, Bracklesham: Turbonia, Dendrophyllia, Gonio pora. (From Prestwich)</td>
<td>57</td>
</tr>
<tr>
<td>28.</td>
<td>A simple form of Crinoid: Botryocrinus</td>
<td>60</td>
</tr>
<tr>
<td>29.</td>
<td>Types of Cystidea: Aristocystis, Echinosphaera, Protocrinus, Lepadocrinus</td>
<td>64</td>
</tr>
<tr>
<td>30.</td>
<td>A typical Blastoid: Orophocrinus</td>
<td>66</td>
</tr>
<tr>
<td>31.</td>
<td>Edrioaster Bigsbyi. (After Bather)</td>
<td>67</td>
</tr>
<tr>
<td>32.</td>
<td>Palaeozoic starfishes: Schuchertia, Palasterina</td>
<td>69</td>
</tr>
<tr>
<td>33.</td>
<td>A Palaeozoic Ophiurid: Lapworthura</td>
<td>70</td>
</tr>
<tr>
<td>34.</td>
<td>Types of Fossil Echinoidea: Palaechinus, Archaeo cidaris, Cidar is, Hemicidar is, Salenia, Dysaster, Enallaster, Catopygus. (From Messrs. Black)</td>
<td>73</td>
</tr>
<tr>
<td>35.</td>
<td>The tubicolous Polychaeta: Ortonia</td>
<td>78</td>
</tr>
<tr>
<td>36.</td>
<td>Supposed coiled tubes of Polychaeta: Spiorbis</td>
<td>79</td>
</tr>
<tr>
<td>37.</td>
<td>A worm-casting from the Solenhofen Stone: Lumbricaria. (After Baier)</td>
<td>80</td>
</tr>
<tr>
<td>38.</td>
<td>Diagram of a Trilobite: Dalmanites</td>
<td>82</td>
</tr>
<tr>
<td>39.</td>
<td>Reconstruction of a Trilobite: Triarthrus. (After Beecher)</td>
<td>84</td>
</tr>
<tr>
<td>40.</td>
<td>Examples of Trilobites: Agnostus, Olenus, Stau rocephalus</td>
<td>85</td>
</tr>
<tr>
<td>41.</td>
<td>Restoration of a Eurypterid: Simonia. (After Laurie, from Woods)</td>
<td>87</td>
</tr>
<tr>
<td>42.</td>
<td>A modern Xiphosura: Limulus</td>
<td>89</td>
</tr>
<tr>
<td>43.</td>
<td>Silurian primitive Scorpions: Palaeophonus. (After Pocock, from Lankester)</td>
<td>90</td>
</tr>
<tr>
<td>44.</td>
<td>Eophrynus Prestvicii from the Coal Measures. (After Pocock)</td>
<td>91</td>
</tr>
<tr>
<td>45.</td>
<td>A typical Crustacean: Glyphaea. (From Woods)</td>
<td>92</td>
</tr>
<tr>
<td>46.</td>
<td>Fossil Crustacea: Dromilites, Palaeocorystes, Eryon, Meco chir us, Cypridea, Loricula</td>
<td>94</td>
</tr>
<tr>
<td>47.</td>
<td>Anthrapalaemon Woodwardi, Coal Measures. (From H. Woodward)</td>
<td>96</td>
</tr>
<tr>
<td>48.</td>
<td>Isopods: Palaega, Aega</td>
<td>97</td>
</tr>
<tr>
<td>49.</td>
<td>An ancestor of the Crabs: Palacinachus. (After H. Woodward)</td>
<td>101</td>
</tr>
<tr>
<td>50.</td>
<td>A fossil crab: Bhachiosoma. (After H. Woodward)</td>
<td>103</td>
</tr>
<tr>
<td>51.</td>
<td>A fossil millipede: Eu phoberia</td>
<td>104</td>
</tr>
<tr>
<td>52.</td>
<td>Cockroaches of the Coal Measures: Eto blattina, Progonoblattina. (After Seudd and Heer)</td>
<td>106</td>
</tr>
<tr>
<td>53.</td>
<td>Wings of Neuroptera from the Coal Measures: Lithosialis, Bro diea</td>
<td>106</td>
</tr>
<tr>
<td>54.</td>
<td>A Neuropterous (?) insect from the Coal Measures: Lithomantis. (After H. Woodward)</td>
<td>107</td>
</tr>
<tr>
<td>55.</td>
<td>Shell of a Silurian Brachiopod, Atrypa</td>
<td>108</td>
</tr>
<tr>
<td>FIG.</td>
<td>LIST OF ILLUSTRATIONS IN TEXT.</td>
<td>PAGE</td>
</tr>
<tr>
<td>------</td>
<td>--------------------------------</td>
<td>------</td>
</tr>
<tr>
<td>56.</td>
<td>Internal cast of a Brachiopod shell, Camarophoria</td>
<td>110</td>
</tr>
<tr>
<td>57.</td>
<td>Muscle-scars in a shell of Lingula</td>
<td>110</td>
</tr>
<tr>
<td>58.</td>
<td>Types of Brachiopod Orders: Iphidea, Orbiculoidea, Clitambonites</td>
<td>112</td>
</tr>
<tr>
<td>59.</td>
<td>A spire-bearing Telotretamatus Brachiopod: Spirifer</td>
<td>114</td>
</tr>
<tr>
<td>60.</td>
<td>A loop-bearing Telotretamatus Brachiopod: Magellania</td>
<td>114</td>
</tr>
<tr>
<td>61.</td>
<td>Structure of a Bryozoan zooid</td>
<td>116</td>
</tr>
<tr>
<td>62.</td>
<td>A Trepostomatous Bryozoan: Callopora. (After Ulrich)</td>
<td>118</td>
</tr>
<tr>
<td>63.</td>
<td>A Cryptostomatous Bryozoan: Fenestella. (After Ulrich)</td>
<td>118</td>
</tr>
<tr>
<td>64.</td>
<td>Cyclostomatous Bryozoa: Stomatopora, Berenicea. (After Gregory)</td>
<td>119</td>
</tr>
<tr>
<td>65.</td>
<td>Chelostomatous Bryozoa: Membranipora, Onychocella. (From Gregory)</td>
<td>120</td>
</tr>
<tr>
<td>66.</td>
<td>Shell of Chiton squamosus</td>
<td>124</td>
</tr>
<tr>
<td>67.</td>
<td>A Pteropod, Cleodora pyramidata</td>
<td>125</td>
</tr>
<tr>
<td>68.</td>
<td>Shells of non-marine Mollusca: Unio, Corbula, Pisidium, Helix, Paludestrina. (From B. B. Woodward)</td>
<td>128</td>
</tr>
<tr>
<td>69.</td>
<td>Eocene Gastropods and Lamellibranchs: Cyprina, Pholadomya, Crassatellites, Cyprea, Clavatides, Cardia</td>
<td>130</td>
</tr>
<tr>
<td>70.</td>
<td>A Senonian Lamellibranch, Spondylus spinosus</td>
<td>132</td>
</tr>
<tr>
<td>71.</td>
<td>Cretaceous Lamellibranchs: Neithia, Actinoceramus</td>
<td>132</td>
</tr>
<tr>
<td>72.</td>
<td>Lamellibranchs of the Lower Lias, Hippopodium, Gryphaea</td>
<td>137</td>
</tr>
<tr>
<td>73.</td>
<td>Pleurotomaria Quoyana and P. platyspira</td>
<td>138</td>
</tr>
<tr>
<td>74.</td>
<td>Carboniferous Lamellibranchs: Carbonicola, Posidonomya</td>
<td>139</td>
</tr>
<tr>
<td>75.</td>
<td>Lower Palaeozoic Lamellibranchs and Gastropods: Cardiola, Pterinaea, Bellerophon, Platyceras, Maclurea</td>
<td>140</td>
</tr>
<tr>
<td>76.</td>
<td>Conularia quadrisulcata, Coal Measures</td>
<td>141</td>
</tr>
<tr>
<td>77.</td>
<td>Miocene Pteropods: Hyalaea, Vaginella</td>
<td>144</td>
</tr>
<tr>
<td>78.</td>
<td>Cephalopods: Rhyneoliths, Nautilus, Loligo</td>
<td>145</td>
</tr>
<tr>
<td>79.</td>
<td>Shells of Endoceras. (After Holm and Foord)</td>
<td>147</td>
</tr>
<tr>
<td>80.</td>
<td>Primitive Nautiloida: Piloceras, Orthoceras</td>
<td>148</td>
</tr>
<tr>
<td>81.</td>
<td>Initial chambers of Cephalopod shells. (From Foord and Crick)</td>
<td>150</td>
</tr>
<tr>
<td>82.</td>
<td>A Nautilus and an Ammonite</td>
<td>151</td>
</tr>
<tr>
<td>83.</td>
<td>Aptychus of an ammonite</td>
<td>153</td>
</tr>
<tr>
<td>84.</td>
<td>Animal and shell of a Belemnite. (After D'Orbigny)</td>
<td>154</td>
</tr>
<tr>
<td>85.</td>
<td>The Belemnite and its descendants: sections of shells</td>
<td>155</td>
</tr>
<tr>
<td>86.</td>
<td>Cephalopods with thin enclosed shell: Belemnoteuthis, Dorateuthis</td>
<td>156</td>
</tr>
<tr>
<td>87.</td>
<td>Female Argonaut</td>
<td>157</td>
</tr>
<tr>
<td>88.</td>
<td>Actinoceras Bigsbyi, siphuncle. (From Foord)</td>
<td>158</td>
</tr>
<tr>
<td>89.</td>
<td>Three-lobed aperture of Gomphoceratid</td>
<td>159</td>
</tr>
<tr>
<td>90.</td>
<td>Four-lobed aperture of Gomphoceratid</td>
<td>159</td>
</tr>
<tr>
<td>91.</td>
<td>Ascoceratidae. (From Foord, after Lindström)</td>
<td>160</td>
</tr>
<tr>
<td>92.</td>
<td>Palaeozoic Nautiloida; Ophidioceras, Hercoceras, Apheleceras, Vestinatulus</td>
<td>162</td>
</tr>
<tr>
<td>93.</td>
<td>Goniatites: Pronorites, Glyphioceras, Agathiceras. (From Foord and Crick)</td>
<td>164</td>
</tr>
<tr>
<td>94.</td>
<td>Ceratites nodosus</td>
<td>166</td>
</tr>
<tr>
<td>95.</td>
<td>Ammonites: Trachyceras, Lytoceras, Coeloceras, Phylloceras.</td>
<td>168</td>
</tr>
<tr>
<td>96.</td>
<td>Cretaceous Ammonoidea: Cricoceras, Heteroceras, Turritilites, Macrosiphaptes, Hamites, Scaphites, Baculites</td>
<td>172</td>
</tr>
</tbody>
</table>
PREFACE.

The "Guide to the Fossil Invertebrates and Plants" having run out of print, it is in part replaced by this book dealing with the Invertebrate Animals, while it is proposed that the Fossil Plants shall form the subject of a separate Guide. By describing the Galleries and their contents in a different order from that followed in the previous Guide, it is attempted to present the whole as a connected story. At the same time the book does not profess to be a complete systematic treatise, but relates only to the specimens actually exhibited in the Galleries.

As in the other Guides to the fossil animals, the present one assumes on the part of the reader at least so much knowledge of the existing world of life as is conveyed by the corresponding Guides to the Department of Zoology. Since, however, many of the groups of animals herein dealt with are entirely or almost entirely extinct (e.g. Graptolites, Cystids, Trilobites, Ammonites), they have been treated more fully than those which are more familiar.

The book has been written by Dr. Francis Arthur Bather, Assistant Keeper of Geology, who desires to acknowledge the kind help that he has received from his colleagues: Mr. R. Bullen Newton and Mr. G. C. Crick in connection with the Mollusca; Mr. W. D. Lang in connection with the Bryozoa and Coelentera; and Dr. W. T. Calman in connection with the Arthropoda. Thanks are also due, for similar assistance, to Mr. S. S. Buckman, Mr. H. W. Burrows, Mr. C. D. Sherborn, and Mr. B. B. Woodward. For permission to use copyright illustrations the Trustees are indebted to Messrs. A. & C. Black, Messrs. Archibald Constable & Co., Messrs. Longmans & Co., the Cambridge University Press, the Council of the Geological Society, and Dr. Henry Woodward, F.R.S. (Editor of the Geological Magazine).

E. RAY LANKESTER,
Director.

British Museum (Natural History),
23 April, 1907.
A GEOLOGICAL TIME-SCALE, WITH EXAMPLES OF FOSSILIFEROUS ROCKS.

<table>
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<th>Era</th>
<th>Epoch</th>
<th>Age</th>
<th>British Rocks</th>
<th>European Rocks</th>
<th>Range in time of life-groups</th>
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<td>TERTIARY</td>
<td>HOLOCENE</td>
<td>PRESENT DAY</td>
<td>Blown sand, alluvium, beaches, clay, peat, shell-beds, etc., as now forming.</td>
<td>Palaeocene Limestone.</td>
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<td>White Lias, Purbeck Beds.</td>
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<td>Waterstones, Elgin Sandstone, Teign Marsh Beds.</td>
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<td>Red Marls and Sandstones, Purnish Beds.</td>
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<td>Millstone Grits; Yoredale Beds.</td>
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<td>Bala, Chirbury and Cadocian Series.</td>
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<td>Arcadian, Shelsley, and Shelsley Series.</td>
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<td>OLENIDIAN</td>
<td>Arthenian Series.</td>
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<td>OLONELLIAN</td>
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A large series of rocks of which only the uppermost have yielded fossils, and those for the most part obscure, as the worm burrows in the London Clay.

RELATIVE LENGTHS OF EPOCHS AS REPRESENTED BY THICKNESS OF ROCKS.

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<tr>
<th>ERAS</th>
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<tbody>
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<td>PLEISTOCENE</td>
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<tr>
<td>HOLOCENE</td>
<td>2,500 ft.</td>
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<td>CENOZOIC</td>
<td>6,000 ft.</td>
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<td>PERMIAN</td>
<td>1,500 ft.</td>
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<td>JURASSIAN</td>
<td>7,000 ft.</td>
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<td>SILURIAN</td>
<td>15,000 ft.</td>
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<tr>
<td>CAMBRIAN</td>
<td>11,000 ft.</td>
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Note: The divisions on the timescale are not exactly comparable, but those for the period of erosion, which is based on the division of ages, is not exact.
### A GEOLOGICAL TIME-SCALE

#### PALÆOZOIC OF PRIMARY.

| PERMIAN | THURINGIAN | PUNJABIAN | ARTISKIAN | MAGNESIANS LIMESTONES | RED MARLS | MILLSTONE GRIT | MANTEL GRIT | LONDON CLAY | UPPER CHALK | MIDDLE CHALK | LOWER CHALK | KELLAWAYS RED MARLS | KELLAWAYS RYET | KEMPSTON RYET | CORALLIAN LIMESTONES | CORALLIAN CLAY | CORALLIAN FLINT | BARTONIAN CLAY | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | BARTONIAN | 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A GUIDE

to the

FOSSIL INVERTEBRATE ANIMALS.

This book is a guide to the fossil remains of Invertebrate Animals and to a few subsidiary collections contained in Galleries Nos. VII, VIII, IX, and XI of the Geological Department. These Galleries may be approached by turning to the left, either immediately on entering the large Gallery of fossil Mammals, or on reaching the end of it. Both these turns to the left lead into the long Gallery No. IV containing fossil Reptiles, out of which lead the four Galleries here described (See Frontispiece-Plan, Plate I).

Gallery VII is the westernmost and is nearest the main entrance, and Gallery XI is furthest away, towards the east. The entrance to each Gallery is at its south end. The Wall-cases and Table-cases in the Galleries are numbered consecutively, No. 1 being, in the wider Galleries, on the left hand of the visitor as he enters, and the highest number being on his right. Following this order of the Cases, the contents of these Galleries are briefly as follows:—

Gallery VII: Cephalopods (Ammonites, Belemnites, &c.).

Gallery VIII: All Molluscs (except Cephalopods), Bryozoa, Brachiopods, Arthropods (insects, lobsters, trilobites, &c.), ringed worms, Echinoderms (sea-urchins, star-fishes, &c.).

[Gallery IX: Geological Library and Students' Collection, closed to the public.]

Gallery X: Coelentera (corals, graptolites, &c.), Sponges, Protozoa, Plants (these last to be described in a separate Guide-book).

Gallery XI: Stratigraphical series of rocks and fossils; Forms and Structures produced by geological agents, and in some cases simulating fossils; Tracks and Footprints left by various animals; and Collections of Historical interest.
For the purposes of this Guide, it is more convenient to take the collections in the reverse order from that given above, namely, to begin with Gallery XI. The description of the specimens exhibited in that Gallery serves to explain the meaning of fossils and the use made of them in interpreting the structure and history of the earth. We then proceed to Galleries X, VIII, and VII, beginning with the simplest forms of life and passing gradually to those more highly organised. Under each group of animals too, the description generally begins with the older fossils and traces the history of the group down to our own day. In order to follow this method of description, it is, in Galleries X and VIII, necessary to visit the Cases in the reverse order of their numbering.

The collections in Galleries X, VIII, and VII are arranged, in the main, according to a zoological classification, the specimens belonging to the various large groups of the Animal Kingdom being placed together. In some groups it has been found more convenient to subdivide the specimens according to the geological epochs to which they belong, and under each of those epochs again to arrange them in zoological order. In other groups the zoological system is the dominant one throughout, all species of each genus being placed together. It is also the general rule that the specimens from British localities are exhibited in the Table-cases, and those from foreign localities in the Wall-cases. This, however, is not rigidly adhered to; for example, among the larger specimens mounted on blocks and placed usually towards the back of the Wall-cases may often be found several British specimens, which have to be exhibited there on account of their size.

Although all the animals whose fossil remains are herein described are often spoken of collectively as Invertebrata, or backbone-less animals, in contrast to the Vertebrata, or animals with a backbone, yet it must be remembered that this common usage does not represent a scientific classification. It is nearer the truth to say that each of the larger groups named above—Mollusca, Arthropoda, Echinodermata, Coelentera, and the rest—has the same independence and importance as a division of the Animal Kingdom as has the whole group of Vertebrata. Just as the Vertebrata are divided into Classes, namely, Mammals, Birds, Reptiles, Amphibians, and Fishes, so is each of these great groups, or Phyla as they are termed, divided into Classes. Each Class
is again divided into Orders, and sometimes Sub-Orders, and each of these into Families. The names of the Orders are usually printed in large type on cards fixed in the Exhibition Cases. The names of the Families are printed on labels mounted on wood and inserted in their places in the series; they may be recognised by always ending in —ideae. The Families again are divided into Genera, each composed of a number of allied Species. The generic names are printed and mounted on wood in similar fashion to the family-names.

The names on the labels of the separate specimens consist of the name of the genus (e.g. Conchidium) followed by the trivial name (e.g. biloculare), the two together making up the name of the species. This is followed by the name of the author who first described and named the species, e.g. Conchidium biloculare, Linnaeus. When the species has been transferred to a genus other than that in which it was placed by the original author, then that author's name is placed within brackets or followed by "sp." Thus Brünnich in 1781 described as Trilobus caudatus a trilobite which is now placed in the genus Dalmanites; therefore we write Dalmanites caudatus (Brünnich). Since 1839, however, this trilobite has generally been placed in the genus Phacops, and appears in the text-books as Phacops caudatus. In this Guide-book the generic name by which a species is generally known is added to the correct name within square brackets, e.g., Dalmanites [Phacops] caudatus (Brünnich). When a genus has been split up into sub-genera, the name of the sub-genus to which a species belongs may be inserted within round brackets between the name of the genus and the trivial name, thus, Orthis (Dalmanella) elegantula.

Many of the exhibited specimens bear small discs of green or red paper. A green disc indicates that the specimen bearing it either is the original specimen on which the species to which it belongs was based, or has been described and figured in some scientific work, to which a reference is given on the label. Specimens marked with red discs have been merely noticed or briefly described in some published work.
STRATIGRAPHICAL SERIES AND HISTORICAL COLLECTIONS.

Gallery XI. Hanging on the wall of Gallery XI, immediately to the left of the entrance is a diagram showing very broadly the Geological Epochs during which the rocks found in Great Britain were formed, the newest being at the top and the oldest at the bottom of the column. A more elaborate list, with the Epochs divided into Ages, is given as a table (facing p. 1) in the present Guide.

Adjoining the diagram, in the wall-cases on the west side of the gallery, is the Stratigraphical Series, which is a collection of the various kinds of rock found in Britain, arranged in order of age. Along the top of the Cases is a diagram showing the succession of these rocks from the newest to the oldest, as they might be seen in a continuous section across the country from east to west. Examples of the rocks themselves occur on the shelves below, where will also be found numerous small sections of the strata, as observed in various parts of England. Affixed to the Cases is a series of small maps, each coloured to show the tract of country occupied by the one or two rock-groups of which specimens are exhibited in the adjoining Case. In the long section, the numbers placed beneath the beds give their approximate thicknesses in feet. It must not, however, be supposed that all these beds occur in such regular succession right through the country, the fact rather being that one is found in one district, while another is better developed elsewhere, as indeed may often be gathered from the names applied to the beds. Certain gaps in the section, as between Pliocene and Eocene, and again between Permian and the Coal Measures, represent intervals of time, during which there were being deposited rocks, which are found in other parts of the world, but for one reason or another do not occur in the British area.

It is plain that when rocks have been deposited, as we know that they now are being deposited at the bottom of the sea, then the underlying rocks are older than those above them. As concrete examples of the way in which one layer of rock is found lying on another, there are placed on the floor between the Wall-cases in various parts of the Gallery several examples of the cores of rock brought up from below by deep borings. Thus at Dover the boring for coal went down through the Chalk at the surface, through several rocks
underneath this, till it came to the sandstones of the Coal Measures, of which cores from depths of 1262 feet and 2234 feet are here exhibited, as well as a piece of coal from 2039 feet. At Ware, in Hertfordshire, a boring through the Chalk brought up from a depth of 825 feet the core of Wenlock Shale here shown. At the far end of the Gallery are cores of Carboniferous Limestone obtained beneath the Lower Jurassic rocks close to Northampton, from depths of 805 to 828 feet. We learn from these borings that the rocks found on the surface in the western and north-western parts of England, pass beneath other rocks and stretch under the south-east of England and presumably under the sea until they come again to the surface in Belgium and the north-east of France. Thus we have proof quite easy to understand that in this country the older rocks pass generally from north-west to south-east under newer ones, as shown in the long section at the top of the Wall-cases.

Among the specimens selected in illustration of the various beds are many containing the remains of animals or plants. Thus, the very first specimen at the top left-hand corner of Case 1 contains fragments of bone embedded in a stalagmitic deposit which formed on the floor of Brixham Cave; and this indicates that the animals to which the bones belonged lived, or at least died, in the cave, where their remains were gradually covered by the limy deposit. Close by is a piece of an old beach from Brighton, in which is embedded part of a horse's leg-bone. Below these are other specimens of beach-deposits, in which may be seen the remains of shells. Lower down in the Case are rocks of more sandy nature, such as are now being formed off shore, and in them also may be seen shells, as well as the remains of other marine animals and plants. These either lived at the bottom of the sea or sank to it when dead, and were then gradually covered by sand or clay produced by the wearing of the land and deposited on the sea-floor. All such remains or traces of animals and plants found in the rocks are called fossils.

Although the specimens exhibited in this series are not intended to give anything like a complete idea of the animals living in former periods of the earth's history, still as the visitor passes down the Gallery, he will readily observe that the fossils contained in the fragments of rock gradually change in character. Those in the first Case are, as has been seen, similar to animals living at the present day; but
already in Case 2, where are exhibited specimens from the Chalk, many forms have an unfamiliar appearance, and indeed belong to types of life which no longer exist. A like strangeness characterises the Jurassic fossils, but is still more noticeable among the older rocks: thus Case 4 contains some of the curious plants from the Coal Measures, while in Case 5 are fragments of Old Red Sandstone with the strange fishes characteristic of that period. Closer inspection would show that this change was gradual and continuous, and that each of the successive beds of rock was characterised by fossils differing from those found in the beds above and below. Sometimes the bed itself may change in mineralogical character, while the fossils remain the same. Therefore, when once a geologist knows the fossils characteristic of the various strata he can, if set down in any part of the country, readily determine on which bed in the geological series he is standing, if only he can find a few fossils.

The credit, at least so far as British geology is concerned, of first recognising this important fact is due to William Smith (Plate II), whose bust, a copy of that by Chantrey in All Saints' Church, Northampton, is on the eastern wall of the Gallery. The son of a small farmer and mechanic, Smith was born at Churchill, Oxfordshire, in 1769, and at an early age collected the fossils that occur in the rocks around his home. When the boy was eight years old his father died, leaving him to the care of an uncle who, noticing the studious habits of his nephew, gave him some money to buy books. By means of these he taught himself to such purpose that at the age of eighteen he obtained employment as a land surveyor in Oxfordshire and the neighbouring counties and, in 1793, was appointed to survey the course of the intended Somersetshire coal canal near Bath. Six years' work on this canal, added to his previous knowledge, enabled him to prove that the strata met with in this district followed each other in a regular and orderly succession, each bed being marked by its own characteristic fossils, and having a general tendency to slope or dip to the S.E. That this succession was no local phenomenon, and that the same fossils were throughout characteristic of the same beds, was subsequently proved by Smith in his journeyings over the greater part of Britain. The surveys made on these journeyings enabled him, in 1815, to publish the large map exhibited on the right hand of the entrance to this Gallery. This, the first geological map of England
and Wales, comprising also a part of Scotland, measures 8 feet 9 inches by 6 feet 2 inches. Several sections across England, published by Smith in 1819, are placed on the wall around his bust. Here also are reproductions of the first small sketch for the larger map coloured by him in 1801, of a map of the country around Bath coloured geologically by him in 1799, and of a table of strata dictated by him in the same year.* The original MSS. of these were presented by Smith to the Geological Society in 1831.

Smith's views on the value of fossils to the geologist and surveyor were enunciated in his works "Strata identified by organised fossils," of which four parts only were published (4to, 1816–1819), and "Stratigraphical system of organised fossils" (4to, 1817). A set of the plates from the former work is exhibited in the Case below Smith's bust and in a frame on the adjoining wall. The different colours to the backgrounds of the plates are the same as those employed by Smith in his geological map, and have continued in general use, with many of our common geological names for British formations, such as Lias, Greensand, Coral Rag, and Cornbrash, all of which were adopted by him from the local terms in use by quarrymen and others. The fossils illustrated in these works, with many others collected by Smith, are contained in the same cabinet, and form the most characteristic memorial of one who was justly termed by Adam Sedgwick "the father of English geology."

Besides the William Smith Collection, acquired by the Trustees in 1816 and 1818, there are arranged in the Table-cases of this Gallery eight other collections of special interest as bearing either on the early history of the British Museum or the study of geology and palaeontology in this country.

At the end of the Gallery will be found the oldest and, in some respects, the most interesting of these, under the heading The Sloane Collection. Here are still retained in their old association just one hundred specimens out of the large series that once formed the museum of Sir Hans Sloane (1660–1753), who by the terms of his will, may be considered the first founder of the British Museum, since he offered his collection to the nation for the relatively small sum of £20,000, in order "that it might be preserved and maintained, not only for the inspection and entertainment of the learned and the curious, but for the general use and benefit of the

Gallery XI. public to all posterity." The collection was purchased on these terms in 1753, and the British Museum, then in Montagu House, Bloomsbury, was opened to the public in 1759. The geological portion contained many thousand specimens of minerals and of "extraneous fossils, comprehending petrified bodies, as trees, or parts of them, herbaceous plants, animal substances," and the like. It included the large collections previously formed by William Courtén (1642–1702) and James Petiver (1658–1718). In 1857 the minerals were removed from the collection to the newly instituted Department of Minerals, and it is only the "extraneous fossils" that are now preserved in the Geological Department. Each specimen in the Sloane Collection had originally a number attached to it, corresponding to a carefully prepared MS. catalogue, still preserved in the library of this Department, and containing many curious entries concerning the various objects. In the course of over a century and a half many of the labels have become detached from the objects, or obliterated by cleaning, so that although other specimens from the Sloane Collection may be in the Department, it is no longer possible to identify them, and even among those here gathered together, there are some which cannot be referred to their original entry. So far as possible, however, the original words applied to the specimens by Sloane himself have been reproduced on the label, and thus the collection is of particular interest as showing the way in which such specimens were regarded by an eminent naturalist in the early part of the eighteenth century, and throws some light upon various names now disused, but then generally employed by scientific writers. Among the specimens attention may be directed to the chambered portion or phragmocone of a belemnite brought from Japan by Engelbrecht Kaempfer, some Echinites or fossil sea-urchins from Dr. Lavater, a coral from Mr. Beaumont, F.R.S., and especially the Echinites from Agostino Scilla's Collection. Scilla was a Sicilian painter, who in 1670 published an important book on fossils. By these specimens the Museum is connected with some of the famous collections in the early history of geology.

Adjoining the Sloane Collection, and in the same Table-case, is a collection of 124 Tertiary fossil shells obtained by Gustavus Brander (1720–1787) from the cliffs of Barton in Hampshire, and presented by him to the Museum in 1765. The collection was described by D. C. Solander, an officer of the Museum, in a work entitled, "Fossilia Hantoniensia
collecta, et in Musaeo Britannico deposita a Gustavo Brander," London, 1766. The specimens retain the original names given by Solander, underneath which are the names now in general use. Those figured in the book are distinguished by a disc of green paper, as previously explained (p. 3).

The next collection is that containing many of the specimens figured in the "Icones fossilium sectiles" (1820, 1825), an illustrated work on miscellaneous fossils in the British Museum, prepared by Carl D. E. Koenig (1774–1851), who was the first keeper of the Mineralogical and Geological Department, after its separation from the general natural history collections in 1825.

Following this is a collection of fossils from the Carboniferous Limestone of Bolland, formed by William Gilbertson of Preston. This owes its great scientific importance to the fact that the specimens in it were described and figured by Professor John Phillips in Volume II. of his "Illustrations of the Geology of Yorkshire" (1836). Thus this collection is particularly rich in the identical specimens upon which the various species were originally based, in other words, the type-specimens of the species; and it acquires additional importance from the fact that Phillips' own collection was stolen from him on his arrival in London by thieves, who are said to have thrown it into the Thames in their disgust at finding the booty was of no value to them. The Gilbertson Collection was purchased by the Trustees of the British Museum in 1841, but was for many years kept in the Zoological Department.

The next collection, from which only a selection is exhibited, was formed by a naturalist who devoted his entire life to the study and illustration of a single class of organisms, namely, the Brachiopoda. This was Thomas Davidson (1817–1885), whose great monograph on the British fossil Brachiopoda was published by the Palaeontographical Society between 1850 and 1886. The collection contains many of the specimens therein described, as well as an excellent series from foreign localities; it also includes the specimens described in Davidson's Monograph of recent Brachiopoda (Trans. Linnean Soc. 1886–1887). The entire collection of 22,831 specimens was bequeathed by him to the Trustees of the British Museum and handed over by his son, William Davidson, Esquire, in 1886, with Davidson's original drawings, and his library relating to the subject.
The next three Table-cases contain the greater part of the collection which formed the basis of the "Mineral Conchology of Great Britain," a work by James Sowerby (1757–1822) and his son, James de Carle Sowerby, of which successive parts, issued between June, 1812, and January, 1846, amounted to seven volumes in 8vo., illustrated with 648 plates, engraved by the authors and, in some of the later parts by G. B. Sowerby and by J. W. Salter, afterwards Palaeontologist to the Geological Survey. The collection comprises about 5000 fossils, from all parts of England and from every geological formation, many of them named and described for the first time in the "Mineral Conchology," and therefore the type-specimens of the species to which they are referred. Many of the green discs indicating figured specimens were actually fixed by James Sowerby. The ammonites of this collection, being inconveniently large for exhibition in these Table-cases, have been removed to the general collection of Cephalopoda in Gallery VII. The collection was purchased by the Trustees of the Museum from Mr. J. de Carle Sowerby in 1861.

The two collections which follow owe their inception to a society known as The London Clay Club, founded in 1838 by a few London geologists—namely, J. S. Bowerbank, Searles V. Wood, John Morris, Alfred S. White, Nathaniel Wetherell, J. de Carle Sowerby, and F. E. Edwards. Originally intending to illustrate the British Eocene Mollusca, they eventually in 1847 founded the Palæontographical Society for the purpose of monographing all the fossils of the British Isles. Here is exhibited the collection of Eocene Mollusca, begun by Frederick E. Edwards (1799–1875) about 1835, and continually increased until a few years before his death. It was purchased by the nation in 1873. Starting with the fossils of the London Clay, Edwards extended his researches to the Eocene strata of Sussex, Hampshire, and the Isle of Wight, where he was assisted by Mr. Henry Keeping. This collection served as the basis of six memoirs contributed to the monographs of the Palæontographical Society, 1848–56, and of various other papers published by him. The Eocene bivalves in the collection were described by Searles V. Wood in the volumes of the Palæontographical Society for 1859, 1862, 1870, 1877. About 500 species were thus described and figured, but the collection also contains many new and undescribed forms to which manuscript names were applied by Edwards. A catalogue of the collection, by
Mr. R. Bullen Newton, was published by the Trustees in 1891. The first publication of the Paleontographical Society was Part I. of the “Crag Mollusca”—a monograph by Searles V. Wood (1798–1880) published between the years 1848 and 1861, with supplements in 1871, 1873, and 1879. The collection on which this work was based was begun in 1826, and took about thirty years to form. It represents the Molluscan fauna of the Red and Coralline Crags of the neighbourhood of Woodbridge, and from Aldborough, Chillesford, Sudbourn, Orford, Butley, Sutton, Ramsholt, Felixstowe, and many other localities in Suffolk, also from Walton-on-the-Naze in Essex. The collection was presented by Mr. Wood to the British Museum in 1852, and a supplementary collection was given by Mrs. S. V. Wood, jun., in 1885.

Before leaving these collections, the visitor may again be reminded that their importance lies in the fact that they contain the specimens described in certain classical memoirs, and form therefore the ultimate evidence for the truth of these works. This is particularly the case when a student wishes to make certain of the actual form which was in the mind of the original author when he was describing a new species. Owing to the subsequent discovery of many closely related and intermediate forms, it is usually impossible for a species founded by the older writers to be recognised from their descriptions and figures alone. The specimens themselves must be seen. Consequently these original or type-specimens, as they are called, are of the highest importance in scientific study. A very large number of type-specimens in the various groups of fossils are preserved, and most of them exhibited in the general collection of the Department, where they may be readily found in their proper zoological and geological positions. The similar specimens in this gallery have been kept apart, either owing to the historical interest of the original collections, or in deference to the wishes of their former owners.

For more complete information concerning the numerous collections of fossils that have found their way into the British Museum, reference should be made to “The History of the Collections contained in the Natural History Departments of the British Museum,” vol. i., 1904.
DYNAMICAL SERIES; TRACKS AND MARKINGS.

At the further end of the Gallery are exhibited several illustrations of forms produced by natural agencies, as a rule unconnected with animal or vegetable life, and yet frequently simulating fossil organisms. Some of these illustrate the greater geological agents. The kind of movement that takes place in mountain building is shown by some models constructed by Lord Avebury and presented by him (see Quart. Journ. Geological Society, lix, p 348, and lxi, p. 345). Movements of this kind naturally crumple and contort the rocks, and fragments bent and folded in this way are exhibited in the adjoining Wall-case. Besides crumpling, there is a shearing action, and in some of the slates may be observed trilobites greatly distorted, proving the considerable movement that the particles of rock have undergone. The specimens of "ruin marble" below are also due to slight cracks and displacements of the original rock-bands, a phenomenon even more clearly exemplified in a brightly coloured rock from Johannesburg. The "landscape marble" underneath, also probably owes its origin to subsequent disturbance of the original strata, in some cases perhaps combined with the action of organisms. Earth-movements acting on less compact rocks, such as those containing pebbles or boulders, frequently produce a striation and facetting of the stones, as exemplified in some curiously facetted pebbles of Carboniferous age found in the Punjab. One of these is here exhibited. Beside it are boulders or pebbles polished or striated or faceted, either by the action of ice or by that of wind-blown sand, or even by animals rubbing against them. There are also shown examples of rock-weathering by other agencies, such as atmospheric weathering, and borings by land-shells, sea-shells such as Pholas saxicava, worms, sponges, white ants, and other organisms. Among these specimens the most interesting is a portion of one of the columns of the temple of Jupiter Serapis at Puzzuoli, familiar to all readers of Lyell's "Principles of Geology." The marble has been perforated by boring marine shells (Lithodusmus), which attacked it at a time when, owing to the subsidence of the land, the temple had been submerged more than 20 ft. beneath the sea. The floor of the temple was originally 15 ft. above the level of the sea, and, since submergence, has again been raised to about its original level.
Among the forces which produce perplexing and curious forms and markings in the rocks are those to which the general name of concretionary action is frequently applied. The term is a convenient one, since, as it implies nothing, it cannot well be incorrect. A beautiful example of it is furnished by the dendritic markings seen on many flat surfaces, and often resembling mosses or sea-weeds. The manner of their production is illustrated in some artificial preparations by Professor William Watson. The iron pyrites (sulphide of iron) so common in the Chalk constantly assumes shapes which may be mistaken for fossils, and the same substance often does replace the original constituents of shells and similar objects, which are then said to be pyritised. The passage of water through the rocks, often carrying minerals in solution, is a frequent cause of spotting and of banded structure. The banded flints, of which some fine specimens are exhibited, are exceedingly common examples of such a process.

Concretions are frequently formed around some organic nucleus, as a plant or a fish, in consequence of the chemical action produced by the decaying matter. Some of the larger concretions often become cracked, and the cracks again filled up by the infiltration of carbonate of lime or a similar substance, which forms numerous partitions or septa through the concretion, which is then known as a septarium. In addition to the examples of septaria shown in Wall-case 6, there is a polished section of a fine one from the Oxford Clay of Weymouth, fixed on the end wall of the Gallery. The curious forms assumed by flint in the Chalk should by this time be well known, and yet there are many people who still offer curiously shaped flints to the British Museum under the impression that they are rare fossils. A few specimens obtained in this way are here shown as a warning to others. It will, however, be noticed that many flints do actually contain fossil animals. Among the extraordinary forms assumed by flints, particular mention may be made of the pot-stones or Paramoudras, common in Norfolk. These huge, roughly cylindrical masses, usually with a central cavity, are arranged in vertical columns in the Chalk, and often pass through it for long distances. Specimens and illustrations of these are exhibited.

We come next to a fine series of concretions collected, chiefly by Mr. G. Abbott, from the Magnesian Limestone of Upper Permian Age in Sunderland and Marsden in Durham.
Several of these have a curious structure which produces in them a close resemblance to the skeletons of corals. Some larger masses of rock exhibiting the same structure are in the N.E. corner of the Gallery.

Further examples of concretions and other markings produced in various ways, not themselves organic, but curiously simulating organic objects, such as a human skull, the tooth of an elephant, a dog's head, plants, and fruits, are exhibited in order to impress upon beginners in the study of fossils the truth that here, if anywhere, things are not always what they seem.

The term "fossils," as has already been said, is applied not only to the remains of animals and plants, but to various traces left by them. The footprints of many animals with which we are more or less familiar are easily recognised, and many such exhibited in Wall-cases 8, 9, and 10 on the east wall are described in the Guide to the Fossil Reptiles. The more lowly animals, however, produce tracks which are less well known, and while certain markings found in the rocks can reasonably be explained by reference to the tracks and imprints of animals or plants now living, others still lack a convincing explanation. Here may be seen tracks ascribed to marine worms, crustaceans, and jelly-fishes; others, which have been ascribed to fossil plants and have received learned names accordingly, are now supposed to be either the tracks of some animal, such as a worm, or even the markings left by currents or eddies in the water.

Markings obviously ascribable to such inorganic agencies—for example, ripple marks and the prints of rain drops—have been found in rocks of all ages, appearing just like the "ribbed sea sand" of to-day, or the rain prints newly formed on any wet surface of mud or sand, such as the stretches left when the tide goes out at the Bay of Fundy.

PROTOZOA.

Entering Gallery X, either from that last described or from the Gallery of Fossil Reptiles, No. IV, we pass down its left or western side to the far end. Here are exhibited the remains of the lowest forms of animal life that are preserved as fossils. These are the Foraminifera and the Radiolaria, two sections of the Phylum or great group Protozoa (first, i.e. simplest, animals).
The Protozoa are animals of simple structure and usually of minute size. In them the soft substance of the body is not divided into cells, such as build up the body-tissues of all other animals, and consequently they have no definite tissues. They have been called unicellular animals, but a single Protozoan often comprises more than is to be found in any one cell of a multicellular animal. In the latter each cell is trained for some special service and has dropped the qualities not essential to that service, whereas the body of the Protozoan has to fulfil all functions of the animal economy. Microscopic investigation, moreover, has shown that the minute drop of viscous protoplasm which constitutes the body of even the simplest Protozoan has really a most complicated structure. By keeping this in mind we shall better appreciate the significance of those exquisite skeletons formed by the Foraminifera and Radiolaria and frequently preserved as fossils.

A fuller account of living Protozoa is given in the Guide to the Coral Gallery in the Department of Zoology. It is there explained how in some Protozoa the outer surface of the body is hardened, so that the animal retains always a definite shape, and can usually take in food only at one permanent opening. Further, the surface is generally provided with definite lash-like or hair-like processes capable of rapid rhythmic movement. Those Protozoa are called Corticata, and since they have no skeleton to be fossilised, need no further mention here. The rest of the Protozoa have no hard skin, so that either the whole body may change its form within certain limits, or portions of its protoplasm may be extended as lobes or threads and again withdrawn. By stretching out a lobe, and then as it were flowing into it, the animal moves, and therefore these extensions are called pseudopodia (false footlets). Fragments of food are caught up in them and taken into the central body at any point. These Protozoa are called Gymnomyxa (naked slimes) or Rhizopoda (root-feet). They may be divided into the Classes: I. Lobosa, with lobose pseudopodia, e.g. Amoeba. II. Heliozoa or Sun Animalcules, with fine radiating pseudopodia. III. Foraminifera, with pseudopodia branching and again uniting so as to form a network. IV. Radiolaria, with fine radiating pseudopodia, as in Heliozoa, but with the central protoplasm enclosed in a porous membrane or capsule. In each of these groups there are some genera that construct a skeleton and others that do not; but in the
Lobosa and Heliozoa such genera are few, and their remains have not been found in any rock. In the Radiolaria, on the other hand, the large majority form a skeleton, while in the Foraminifera there are very few without it well developed. These two groups, therefore, are abundantly represented in the fossil state.

Class Foraminifera.

The skeleton deposited by the protoplasm of these animals generally consists of carbonate of lime, and is called the shell or test. Sometimes, especially in fresh-water species, the shell consists only of chitin, and then is never found fossil. In several a chitinous or thin calcareous coat is strengthened by grains of sand or other foreign particles, as is the case of a caddis-worm, and some tests of this arenaceous composition

![Diagram of Living Foraminifera](image)
are preserved in the rocks. This shell encloses the main mass of protoplasm, but the pseudopodia stretch out through certain openings. Either they are confined to the aperture of the shell, or the whole shell-wall is pierced by minute tubes through which the protoplasm can pass (Fig. 1). Shells of the first kind are called Imperforate, and may be chitinous, arenaceous, or calcareous. Shells of the second kind are called Perforate, and are generally calcareous. The Imperforate calcareous shells are chalky-white and rather like porcelain in appearance, whence they are called Porcellaneous. The Perforate calcareous shells are of more glassy appearance, and are therefore called Hyaline. Both types of calcareous shell are made of calcite, but in porcellaneous shells this is less pure. Attempts to classify Foraminifera according to the composition of the shell have not proved satisfactory, since genera or species which in other respects appear closely allied differ in this respect.

The form of the shell varies greatly. It may be simply flask-shaped, without any divisions, e.g. Lagena; or the animal may grow out of the opening of such a single chamber and fashion for itself another chamber adjoining the first, and as it grows it may build chamber after chamber (Fig. 2 a). Shells of the former type have been called Monothalamia (single-chambered), those of the latter type, Polythalamia (many-chambered). A hundred years ago,

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![Fig. 2.—A Foraminifer from the Upper Chalk, Anomalina ammonoides.](image)

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Gallery X.
many of the Polythalamia were supposed to belong to the Cephalopoda, that Class of Mollusca which contains the nautilus and ammonite with coiled and chambered shells so like those of some Foraminifera, but on an enormously larger scale (see p. 151, Fig. 82). A. D. d'Orbigny, however, whose attention had been directed to the shells found by his father in the sea-sand near Rochelle, and who had studied similar sands and fossiliferous deposits from various parts of the world, published a memoir in 1826, pointing out that these microscopic forms differed from the true Cephalopoda in having no tube or siphuncle passing through the chambers, but simply one or many holes or foramina through the walls or septa that separate one chamber from the next; he therefore distinguished them as "Foraminiferes" (hole-bearers).

By observation of the living animal, Dujardin in 1835 discovered the more essential difference that, whereas the body of a nautilus is an elaborate structure confined to the last-formed chamber of the shell, the body of a chambered Foraminifer fills every chamber and is of simple protoplasm, connected throughout. Thus the flinty casts of the chambers of some Foraminifera found in the flint-meal of Chalk flints represent the form of the original animal, without the pseudopodia (Fig. 2 b).

The diversity of form assumed by the shell may be studied in the two series of plaster models shown on the top shelf of the Wall-case. One of these series was made by A. D. d'Orbigny and issued to his subscribers, about 1825, in four "Livraisons" or sets of twenty-five each. The label sent therewith described them as "Models of microscopic cephalopods, recent and fossil, representing one example from each of the chief divisions of a new classification based on the mode of growth of the shell. The diameter of these models is from 40 to 200 times that of the original shells. . . . The coloured models represent the fossil shells; the white models, the recent shells. The position and shape of the siphuncles [the openings between the chambers] are indicated by the marks or black spots." The models are labelled with the names originally attached to them by d'Orbigny, and are arranged in the numerical order in which he sent them out. Unfortunately that order was one of pure convenience and did not correspond with his classification; it is, however, the order followed in a subsequent description of the models and revision of their names by W. K. Parker, T. Rupert Jones, and H. B. Brady (Annals and Mag. Nat. Hist., July, 1865).
Many species marked by d'Orbigny as fossil are now known to live also in modern seas, and many recent ones have since been found fossil. In the following classification these models are referred to as “O 1, 2, 3, &c.”

The adjoining series of models, prepared by Prof. A. E. Reuss and Dr. Anton Fritsch of Prague in 1861, was intended to supplement the series by d’Orbigny. These are arranged in the order of the catalogue issued with them, the classification being that of Reuss. Since an account of it is given in the paper by Parker, Jones, and Brady, quoted above, we need say here only that Nos. 1–30 are Imperforata; Nos. 31–100, Perforata; Nos. 1–18 have arenaceous shells, while all the rest are calcarceous. These models are referred to as “R 1, 2, 3, &c.”

The classification followed in the arrangement of the actual specimens, whether British or foreign, is that used by Brady in his report on the Foraminifera collected by H.M.S. ‘Challenger.’ According to this, the genera are arranged in ten families, nowadays often raised to the rank of Orders. These are: (1) Allogromiidae, hornycoralliaceous and mostly freshwater; none fossil. (2) Miliolidae (Fig. 3, a, b, c), comprise all porcellaneous shells; this division exhibits a great many plans of shell-building, including several that are also found among arenaceous and hyaline Foraminifera; examples of nearly all these occur as fossils. (Models, O 8, 18, 31–33, 90–97, 100; R 16–21, are Miliolidae; O 22, 81, R 26, 27, are Hauerininae; O 16, 20, 21, 24, 48, R 15, 22–25, 29, 30, are Peneroplidae; O 50, Alveolininae). (3) Astrorhizidae, arenaceous and irregular, usually monothalamous and seldom truly septate; Saccammina is well-known in Carboniferous Limestone, but there are few other fossils of this family. (4) Lituolidae (Fig. 3, d, e), arenaceous, truly septate and more regular; the shell-form is diverse; and among exhibited fossil examples of the various types are Lituola (R 1–4, 13), Trochammina, Nodosinella, Stacheia, and Endothyra. Except for a few Endothyraeae, all the preceding families have essentially imperforate shells; the remaining families are essentially Perforata. (5) Textulariidae, usually arenaceous, but the smaller species are hyaline; shells usually composed of a double or triple series of alternating segments; Textularia (Fig. 3, l, O 7, 28, R 75) is a common fossil (O 2, 25, 56–58, 66, R 5–7, 10–12, 14, 28, 76, belong to the same sub-family); Bulimina (Fig. 3, n, O 9, 68, R 8, 9, 64, 85),

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PROTOZOA—FORAMINIFERA. 19

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**Gallery X.**

**Wall-case 9B.**

c 2
Gallery X. *Virgulina* (O 64, R 58, 65), and *Bolivina* (R 79) are exhibited fossil representatives of Bulimininae; *Cassidulina* (Fig. 3, m, O 41, R 81, 82) is the type of a third sub-family. (6) *Cheilostomellidae*, calcareous, thin, and finely perforate, with successive chambers in single, alternate, or triple series. *Allomorphina* (R 74), *Cheilostomella* (R 73), and *Ellipsoidina* are Cretaceous or Tertiary fossils. (7) *Lagenidae* (Fig. 3, f–j), hyaline and finely perforate, simple shell-wall; starting with a flask-shaped or spheroidal chamber as in *Lagenia* (R 33–35), other types are formed by the addition of chambers in straight series, e.g. *Nodosaria* (O 1, R 36–38), or curved e.g. *Marginulina* (O 6, 55, R 55, 57, 59, 60, 72), or spiral, e.g. *Cristellaria* (O 14, 19, 44, 47, 82–85, R 61–63), or in alternating double or triple series, e.g. *Polymorphina* (O 23, 29, 30, 61–63, R 67–70), while in *Ramulina* the chambers are joined by branching tubes; all the genera mentioned are widely distributed as fossils; (others are O 3, 4, 26, 27, 51–54, 60, 67, R 40–54, 56, 66, 77). (8) *Globigerinidae*, hyaline shell of a few swollen chambers spirally arranged, *Globigerina* (Fig. 3 k, O 17, 76, R 91; others are O 43, 65, R 71, 92, 95). (9) *Rotaliidae* (Fig. 3, o, p), hyaline shell typically composed of chambers wound spirally like a snail-shell, and either free or attached; (*Spirillina*, R 31; *Rotaliinae*, O 10, 12, 13, 15, 34–39, 42, 49, 69–75, 77–79, 89, R 83, 84, 86–90). (10) *Nummulitidae* (Fig. 3, r–t), shell consists of a series of chambers coiled in one plane, and is often thickened by extra layers of shell-substance connected with a system of canals distinct from the ordinary communications between the chambers and the perforations of the shell-wall; among important fossils are *Fusulina* (R 96), *Nonionina* (O 11, 46, 86, R 94, 96), *Polystomella* (O 45, R 93), *Amphistegina* (O 40, 98, R 97), *Operculina* (O 80, 88, R 98), *Heterostegina* (O 99, R 99, 100), *Nummulites* (O 87), and *Orbitoides*. The complication of structure sometimes attained in this family may be studied in an enlarged model of a *Nummulite*.


The Epoch mentioned after each genus is the oldest in which it is found; all, except *Archaediscus*, survive to the present. The figures are variously enlarged, from 10 to 100 diameters. (From Nicholson's "Paleontology," after H. B. Brady.)
Fig. 3.—Examples of Foraminifera:

**Miliolidae**—a, *Cornuspira*; b, *Quinqueloculina*; c, *Peneroplis*: all Tertiary to Recent.

**Lituitidae**—d, *Litula*, Carboniferous to Recent; e, *Trochammina*, Lower Lias to Recent.


**Globigerinidae**—k, *Globigerina*, Cambrian to Recent.

The marine Foraminifera, with which geologists are chiefly concerned, are found on sea-weed and similar objects on the sea-floor, from shore pools down to great depths, and from arctic to tropical waters, sometimes fixed and sometimes free; they live chiefly on diatoms and algae. Most of the Globigerinidae float in the warm surface-water of the great oceans down to a depth of 500 fathoms, and stretch out their pseudopodia along delicate spines; these eat also minute animals.

The empty shells are found in all kinds of marine deposits. Numbers are drifted ashore, as at Rochelle and at Dog's Bay, Connemara, whence 124 kinds have been obtained. An ounce of sand from the Adriatic yielded 6,000 shells. Deposits dredged from the sea-bottom contain each a special assemblage varying with the nature of the bottom, depth, and temperature. Such are the coral sands of the Pacific, and the greensands formed at about 500 fathoms. In the latter the empty shells become filled with a green siliceous mineral (glauconite) and often disappear, leaving their casts behind. In the deeper parts of the ocean, especially where the surface is warm, is found an ooze mainly consisting of the shells of Globigerinidae and other pelagic forms (Fig. 4a); its extent is estimated at 49 million square miles, and its thickness must be enormous. Examples of some of these deposits dredged by H.M.S.
"Challenger" are shown in a special Case in the middle of Gallery X. This Gallery.

It is natural, therefore, that Foraminifera should not merely be common fossils, but that they should have helped to build up large masses of rock. From Cambrian to Devonian times they are rare, and no specimens of those periods are exhibited. Among the Carboniferous specimens may be seen limestones composed of the Astrorhizid genus Saccammina from Britain, of the Lituolid Endothyra from Indiana, and the Nummulitid Fusulina from Russia; also isolated shells of Archaeodiscus, a primitive Nummulitid, of the Lagenids Dentalina and Nodosaria, and of Textularia. The last two genera also occur among the foreign Permian specimens, and Trochammina, a Lituolid, among those from Britain. The Jurassic marls and shales contain immense numbers, mostly of the small perforate and arenaceous forms not easily placed on exhibition. Among the British specimens from the Oxford clay is an oyster-shell covered with the irregular adherent Lituolid, Webbina. The Cretaceous system furnishes greensands, in which, as above explained, the actual shells are rarely preserved; but some fragments with the large Rotaliid Patellina are shown in both the British and foreign series. Its characteristic rock, however, is the white Chalk, which in some parts approaches a Globigerina ooze, and contains numerous shells of Globigerina, Cristellaria, Nodosaria, Textularia, Lituola, and other genera (Fig. 4b). These may be preserved in flint, and many such figured in the Rev. H. Eley's "Geology in the Garden" (1859) are in the British Case. From the Maestrichtian Chalk are shown the flat circular Orbitoides and the spurred Calcarina. As examples of Eocene limestones, mainly formed by Foraminifera, may be noted the Paris building-stone with Miliola, an Alveolina limestone from Persia and from Selsea, a French rock with Orbitolites, and another with Orbitoides from both Biarritz and India. During the same period were formed the various Nummulitic limestones, of which numerous examples are shown from countries round the Mediterranean, also from S.E. Africa and India, while in the British series are specimens of the slighter development at Alum Bay and Bracklesham. Here also are many smaller forms obtained by washing the London Clay, and others of Pliocene Age from the Coralline Crag of Suffolk.

The Nummulites have attracted the attention of learned writers from Strabo downwards, but have recently acquired
Gallery X. fresh interest owing to a curious problem connected with the reproduction of the Foraminifera. Examination of some of the pieces of nummulitic limestone will show that nummulites of two sizes are usually associated in the same rock. These were formerly supposed to be distinct species, and so received distinct names. On splitting the shell of a nummulite it is found that the spiral series of chambers starts from a central spherical chamber, and in these paired forms it has been observed that in the larger shell the central sphere is of microscopic size, whereas in the smaller shell it is readily visible to the naked eye (Fig. 5). In

![Fig. 5.—Two generations of a Nummulite. Sections across (a) *Nummulites laevigatus*, showing the small central chamber, and (b) its other form known as *Nummulites Lamarcki*, showing the large central chamber. Both from the Eocene of Stubbington; enlarged 10 diameters. (Copied from the original figures by De la Harpe, 1881.]

consequence of their universal association it was inferred that the large circles with small spheres and the small circles with large spheres were really two forms of the same species; and it was then found that similar dimorphism, i.e., composition of a single species by two forms, occurred in the shells of many other genera of Foraminifera. By observation of living individuals of one of these (*Polystomella*), Schaudinn and Lister have proved that the dimorphism results from the alternation of two modes of reproduction: the small-sphered form extrudes from the shell the whole of its protoplasm, which then separates into spheres, and round each of these is deposited a shell which proves to be the central chamber of a large-sphered form. In such a form the protoplasm ultimately divides into a multitude of minute two-tailed particles or spores, which are ejected. When a
spore from one individual meets a spore from another it fuses with it, and round the united pair grows a shell which proves to be the central chamber of a small-sphered form.

The two forms of shell are to be seen in the following two pairs of so-called species; in each case the name which has to be adopted for the whole species, and under which it is exhibited, is the one printed in italics.

<table>
<thead>
<tr>
<th>LARGE CIRCLE, SMALL SPHERE.</th>
<th>SMALL CIRCLE, LARGE SPHERE.</th>
<th>FORMATION AND COUNTRY.</th>
</tr>
</thead>
<tbody>
<tr>
<td>Nummulites complanatus, Lamarck.</td>
<td>N. Tchihatcheffi, d'Archiac.</td>
<td>Eocene and Miocene, France, Bavaria, Hungary, Egypt, India.</td>
</tr>
<tr>
<td>N. perforatus (de Montfort).</td>
<td>N. Lucasanus (De France).</td>
<td>Eocene, India.</td>
</tr>
<tr>
<td>N. gizehensis (Forskål).</td>
<td>N. curvispirus (Savi and Meneghini).</td>
<td>Eocene, Egypt.</td>
</tr>
<tr>
<td>N. laeavigatus, Lamarck.</td>
<td>N. Lamarcki, d'Archiac and Haime.</td>
<td>Eocene, France, England (Bracklesham).</td>
</tr>
<tr>
<td>Assilina exponens (Sowerby).</td>
<td>Assilina mammillata (d'Archiac).</td>
<td>Eocene, India, Bavaria.</td>
</tr>
</tbody>
</table>


EOZOON.

This green serpentinous rock, in wavy layers, which was formerly thought, chiefly by W. B. Carpenter and J. W. Dawson, to have been built up by a colonial Foraminifer, called by them Eozoön (the dawn animal), is found in some of the very oldest rocks in Canada and Bohemia. A similar mineral structure, however, also occurs in much later rocks, including some of undoubted igneous origin. The organic nature of Eozoön is therefore upheld no longer, but the hypothesis had its uses in the impetus which it gave to the microscopic and chemical study of rocks.
Class RADIOLARIA.

The central mass of protoplasm with the nucleus is surrounded by a horny membrane, forming a capsule. This, in its simplest form, is spherical, and scattered evenly over its surface are numerous pores through which the protoplasm within it communicates with that outside it. In a higher stage the pores are grouped into several areas, but these in their turn are evenly distributed, so that there is still no definite axis. All these constitute the Porulosa of Haeckel.

In the other Radiolaria, which Haeckel names Osculosa, the pores are concentrated in a basal region called the osculum; thus the capsule has a central axis with a basal and an apical pole; smaller oscula may occur near the latter. In each of these divisions a skeleton may be developed, either of pure silica or of other substance.

On these bases the Radiolaria are divided by Haeckel into four Orders: 1. Spumellaria: Porulosa, with pores scattered; skeleton of solid silica. 2. Acantharia: Poru-
losa, with pores grouped; skeleton of strontium sulphate. 3. **Nassellaria**: Osculosa, with one basal osculum; skeleton, when present, of solid silica. 4. **Phaeodaria**: Osculosa, with the capsular wall projecting around the basal osculum as a tube; two or more smaller oscula may be at the apical pole, extracapsular protoplasm contains dark pigment granules (phaeodium); skeleton of silica combined with organic matter.

Only the glassy skeletons of pure silica—as occurring in Orders 1 and 3—are found fossil. These skeletons, though manifesting extreme diversity of shape, conform in general plan to the structure of the central capsule, as may be seen from the greatly enlarged models in Table-case 15. Thus in Spumellaria the skeleton is usually a sphere of lattice-work (Fig. 6 c), or several such spheres one within the other (Fig. 6 b), and joined by cross-bars which radiate from the innermost sphere but do not meet at the centre, and which project as spines (Fig. 6 b and c). The sphere may be pulled out to an ellipsoid, or compressed to a discoid. In Nassellaria the skeleton is generally in the shape of a bell or of an elongated cone, which may be transversely constricted at intervals (as in Fig. 6 d); there may be a spine at the apex and others projecting from the basal margin. All these skeletons are so minute that their form can scarcely be distinguished by the naked eye.

Radiolaria live only in the sea, where they float in all parts and at all depths, but mainly near the surface of tropical oceans. On death their skeletons sink to the bottom, but those not made of pure silica are dissolved by the seawater; and where the ocean is very deep the calcareous shells of the Foraminifera are also dissolved as they sink. Hence at depths of from two to four miles the ooze of the ocean-floor is formed almost entirely of Radiolaria; examples of this, dredged by H.M.S. ‘Challenger,’ are shown in a case in the middle of this Gallery. Similar radiolarian ooze has been deposited in past geological epochs, and when found among the rocks bears witness as a rule to an upheaval of that part of the earth’s crust from a great depth. In the rocks of Tertiary age, such dried oozes are known as Tripoli stone (*Kieselguhr*), and occur in many parts of the world. The Radiolaria from several of them were described by Ehrenberg (1838–1873) under the name Polycistines. On the bottom shelf of Wall-case 9b, in the corner, is a large core of the Miocene radiolarian marl of Barbados, from which
400 species have been described; and a glass slide with Radiolaria from a similar rock in Cuba, with illustrative drawings, is shown in Table-case 15. In these soft Cainozoic deposits many Radiolaria belong to species still living, and their skeletons are as perfect as those in modern ooze. In the Mesozoic and Palaeozoic rocks, however, the oozes have been changed, by pressure, or heat, or the percolation of water, into quartzites, cherts, and flinty shales, so different in appearance that it is not long since their radiolarian origin was discovered. This was done by examining thin sections of the rock under the microscope, when in some, less altered than most, the skeletons were recognised. Usually, however, the skeletons themselves have been dissolved, and there can only be detected spots of transparent silica formerly deposited in the cavity of the skeleton. In this way Radiolaria have been found in siliceous rocks as far back as to the Cambrian period. In illustration of this are exhibited specimens of the radiolarian chert and shale of Carboniferous age, found in the Lower Culm of Devonshire and Cornwall (Fig. 7), and
Ordovician cherts from Cornwall and the south of Scotland. Drawings of the species found in these rocks by Dr. G. J. Hinde are exhibited alongside. Other specimens of Radiolarian rock from foreign localities are also shown. Radiolaria are occasionally found scattered through rocks other than cherts, but as a rule their delicate skeletons have been dissolved away.

See further Haeckel's "Report on the Radiolaria collected by H.M.S. 'Challenger,'" 1887, in which the fossil species are also dealt with.

PORIFERA (Sponges).

Next the Protozoa come the fossil remains of sponges, the foreign ones being displayed in the Wall-cases, and the British ones in the Table-cases. Among the latter will be seen many of those preserved in flint, and familiar either to collectors in the chalk-pits or the searchers for pebbles on the beach. To explain the structure of these a rather long description is necessary.

Sponges are animals most of which live in the sea at all depths, while one family alone is found in fresh water. Since they have no organs for locomotion, sense, or reproduction, they are usually indefinite in shape as well as very variable in size. The only organs readily seen are several holes, often mounted on slight projections at various parts of the surface; from these holes, which are called "ostia," a current of water is always issuing. Closer examination discovers scattered over the whole surface a far larger number of small openings through which the water is as constantly being drawn in; these are called "ostia." Study of thin sections of a sponge under a microscope shows how this flow of water through the sponge is brought about. The water that enters the ostia is led by irregular winding canals into the deeper parts of the sponge. These incumbent canals open by very minute pores into a number of small round chambers, whose walls are furnished with little lashes (or "flagella") in constant motion. From each such flagellated chamber the lashes are continually driving the water through a wide opening into another set of canals, the excurrent canals, which lead it out to the oscula. The sponge, unlike the animals hitherto considered, is a combination of many cells, modified for various services. Thus the outer surface and the walls of the canals are coated with flattened cells; the cells lining the
flagellated chambers bear each a lash surrounded by a collar (collar-cells), so that they closely resemble certain Protozoa; interspersed among these are other cells, each of which is pierced by one of the very minute pores mentioned above; the substance surrounding the chambers is a jelly containing various cells, among them the germ-cells and the spicule-cells. For further information as to the soft parts, recourse should be had to the exhibit in the Zoological Department.

All this soft mass of the sponge is supported by a skeleton built up from the tiny spikes or rods deposited by the spicule-cells. In the common bath-sponge, as in many modern sponges, the substance of the skeleton is horny, and appears incapable of preservation in the fossil state. In

![Fig. 8.—Fossil sponge spicules: Silicispongiae, Heteractinellida, skeletal spicule (a); Silicispongiae, Tetractinellida and others, flesh spicules (b–g); Calcispongiae, skeletal spicules (h–n). a is from Asteractinella; b, c from Geodia; j–l from Tremacystia. a is enlarged 13 diam.; b–g, 66 diam.; h, 26 diam.; i, 114 diam.; j, m, n, 184 diam.; k, l, 80 diam. (After Hinde.)](image)

other sponges the skeleton is either calcareous, i.e. composed of carbonate of lime (calcite), or siliceous, i.e. composed of flinty spicules. The latter are further distinguished by being deposited around an axial filament of softer tissue, which disappears in the fossils, leaving an axial canal. The siliceous sponges constitute the larger and more important group, and are the better preserved as fossils; in some of them the siliceous skeleton partly gives place to horny fibres, a change which suggests how the true horny sponges arose. The spicules, whether of calcite or of silica, are built on certain plans which are utilised in classification. The main types are: (1) Monaxons (Fig. 14), spicules of rod-like form, that grow out from a single point of origin either
in one direction (one-rayed) or in two directions (two-rayed), but in the latter case the two rays lie along a single axis; the axis is straight or curved. (2) Triaxons (Fig. 9), spicules that grow out from a single origin in both directions along three axes at right angles to one another, thus producing six rays, some of which may, however, be suppressed. (3) Tetraxons (Fig. 11), spicules that grow out in one direction only along four axes which meet at equal angles; thus there are four rays, one or more of which may be suppressed. In the latter case tetraxon rays may be distinguished from those of triaxons by meeting at an angle of about 120° instead of 90° or 180°. (4) Polyaxons (Fig. 8, a, d), which grow out along several axes radiating from a common centre. The modifications and associations of these four types can be gathered from the accompanying figures. These types are more conspicuous in the larger spicules that build the skeleton; there are also smaller flesh spicules or skin spicules of more irregular shape (Fig. 8, b–g).

Frequently sponges are divided into only two Classes:—
(I.) Those with spicules of calcite, called **Calcispongiae** or Calcarea; (II.) all the rest, called **Silicispongiae** or Silicea. There is however considerable difference between those Silicea having six-rayed spicules or triaxons, and all the others, so that the following classification has been proposed.

**Branch: CALCISPONGIAE.**

**Class I.—CALCAREA.**

Skeleton of calcareous spicules, either monaxon or tetraxon or both (Fig. 8, h–n).

Grade A.—**Homocoela.** Body a simple sac, which branches in the adult, but retains the simplest type of canal system, and is lined throughout by collar-cells. Although the most primitive forms, these have not yet been found fossil, probably because of their small size and imperfect skeleton.

Grade B.—**Heterocoela.** The canal-system is broken up into separate flagellated chambers, to which the collar-cells are restricted. According to the degree of complexity of these chambers, and according to the shapes of the spicules, they are divided into six families, of which only one, the Pharetronidae, is largely represented in the fossil state. The Sycettidae, in which the chambers are radially arranged round the central cavity, are doubtfully represented.
by the Jurassic Protosycon. The diminutive Leucandra Walfordi from the Middle Lias is the sole fossil example of the common recent family Grantidae, for in them the spicules are loosely and irregularly distributed chiefly around the chambers.

The Pharetronidae have a skeleton of fibres formed by spicules arranged side by side and interlocking, but not fused (except in the Lithoninae). In some a relatively large three-rayed or four-rayed spicule is enveloped by smaller thread-like spicules; in others the spicules are approximately equal throughout. In many there is an outer layer formed of a close felt of spicules. The more important genera are Corynella (Trias to Cretaceous), Holcospongia (Jurassic), Elasmostoma and Peronidella (Jurassic and Cretaceous), and Pharetrospongia (Cretaceous). All these belong to the Sub-family Dialytinae. The Lithoninae, in which the main spicules are fused, contain a few Tertiary and Cretaceous genera, of which the Chalk Porosphaera is the best known.

The structure of the skeleton has often been greatly altered in fossilization. The spicules have lost their outlines, and the fibres now appear as if entirely formed of granular or fibrous calcite. In other cases the fibres have been replaced by silica, so that they remain after treatment with acid, but all trace of spicules has been obliterated. In some specimens of Pharetrospongia, preserved in solid flint, the outer portion of the fibres has been replaced by silica, whilst their interior still retains the original calcite. The structure even in the best preserved specimens is hardly recognisable, unless in thin sections under the microscope.

The Calcarea of to-day are marine shallow-water forms; but some Pharetrones may have lived in deeper water.

Branch: **SILICISPONGIAE.**

Class II.—**HEXACTINELLIDA.**

Skeleton of siliceous spicules (Fig. 9), all triaxon and therefore primitively six-rayed (hexactine). Canal-system simple, with thumbnail-shaped chambers. The body-wall, which is relatively thin but may be thickened by folding, surrounds a wide funnel-shaped or cylindrical cavity (cloaca) into which the chambers discharge their currents (Fig. 10). In this wall the chambers are suspended by rafters (trabe-culae) of soft tissue between an outer dermal membrane and an inner gastric membrane. The spicules are formed by cells
in the trabeculae; some support the dermal membrane, some the gastral membrane, while others stretch across the body-wall between the two membranes; in Lyssacina yet other spicules project from the dermal membrane.

Modern Hexactinellids all live in the deep sea, fixed to the bottom or moored in the ooze by long tufts of rooting spicules. The best known is the beautiful Venus's Flower-basket (*Euplectella*).

**Sub-Class I.—** **LYSSACINA.** The spicules of the skeleton either remain separate or are united at a late period of growth in an irregular manner by siliceous masses or small transverse rods (*synapticulae*).

**Order I.—** **Hexasterophora.** In some of the spicules in the middle layers of the body-wall the rays branch, forming rosette-like bodies called hexasters. The chief

![Diagram of fossil sponge spicules](image)

Fig. 9.—Fossil sponge spicules: Silicispongiae, Hexactinellida. *a*—*f* are skeletal spicules; *g*, *h* are flesh spicules. Six-rayed spicules are shown in *a*, *b*, *e*, *f*; in *c* one ray is suppressed. Axial canals are seen in *a*, *e* and *f*; and *a* and *f* have lantern nodes. *e* and *f* illustrate the union of spicules to form the square mesh of Dictyonina, *e* being from *Sestrodictyon* and *f* from *Coeloplychium*; *g* is called a pinule, and *h* an amphidisc. *a* is enlarged 66 diam.; *b*, *c*, *f*, 40 diam.; *e*, 47 diam.; *g*, 134 diam.; *h*, 114 diam. (After Hinde.)

families are Euplectellidæ, Asconematidæ, and Rossellidæ. Fossil representatives of the last two have been found in Eocene rocks.

**Order II.—** **Amphidiscophora.** There are no hexasters, but some spicules in the limiting membranes are in the form of rods with toothed disc-like expansions at their ends; they are called amphidiscs (Fig. 9 *h*). There is always an anchoring root-tuft. The living forms belong to the family
Gallery X. Hyalonematidae or Glass-rope Sponges, some of which are found fossil as far back as the Silurian. There are a number of extinct families: Protospongidae (Cambrian and Silurian); Dictyospongidae (Silurian and Devonian); Plectospongidae (Ordovician and Silurian); Brachiospongidae (Silurian); Pollakidæ (Carboniferous and Cretaceous), and others. Here also have been placed the peculiar Receptaculitidae (Ordovician to Carboniferous), but it has been shown that their spicules were probably calcareous, and, although still exhibited with the sponges for the sake of convenience, they are now not considered to be sponges at all. Some suppose them to be calcareous algae.

Sub-Class II.—DICTYONINA. The six-rayed spicules of the middle layers of the body-wall overlap by their ends and are then fused by a deposit of silica into a network, or rather rafterwork, with square meshes (Greek dictyon, a net). Owing to their strong skeleton many Dictyonina are well preserved as fossils, representing the following families:—
Euretidæ or Craticularidæ (Jurassic to Recent): mostly cup-shaped or funnel-shaped; spicular nodes simple and imperforate; canal openings large, simple, ending blindly in the skeleton. Mellitionidæ (Cretaceous, Eocene). Coscinoporidae (Cretaceous to Recent): in addition to the cup and funnel shape, many have thin walls folded into a series of flanges, e.g. in the Cretaceous Guettardia; surface canal-openings small, usually arranged in quincunx. Staurodermidæ (Jurassic and Cretaceous): mostly funnel-shaped, with an irregular skeletal mesh and a definite dermal layer in which are large cross-shaped spicules.

Callodictyonidæ (Cretaceous). Ventriculitidæ (Jurassic and Cretaceous): mostly funnel-shaped, with a thin wall thrown into vertical folds which are usually arranged radially (Fig. 10); spicular nodes hollow or lantern-shaped (Fig. 9 a); the base of the sponge has root-like extensions of spicular fibres. Coeloptychidæ (Cretaceous): mushroom-shaped with a thin wall thrown into radial folds and enclosed in a perforate dermal layer; canal-openings in rows on the ridges of the under surface; lantern nodes. Meandrospongidae (Cretaceous to Recent): pear-shaped, sack-shaped, or nodose masses, with a thin wall thrown into numerous folds, which join one another irregularly and are often partly or wholly enclosed by a fine spicular membrane, e.g. Camerospongia, Cystospongia, and Plocosephyia.

Class III.—DEMOSPONGIAE.

Silicispongiae without triaxon spicules. These are the commonest sponges of the present day, most familiar in the freshwater sponge and the bath-sponge, but found in all waters in the most varied surroundings. Palaeontologists, however, are only concerned with those that retain siliceous spicules. According to the form of those spicules they may be divided, somewhat artificially, into two Sub-Classes.

Sub-Class I.—TETRACTINELLIDA. Demospongiae typically with tetraxon spicules (Fig. 11).

Order I.—Choristida (to which the term Tetractinellida is sometimes restricted). Spicules four-rayed and not joined into a rigid network. The simplest spicule has the form of a caltrop (Fig. 11 a, e). In others one ray is elongated, forming a shaft from which the other rays project as three prongs; this trident shape is called a triaene and is subject to much further modification (Fig. 11 f–k). With these...
typical shapes other forms of spicule are associated (Fig. 8 b–g). The spicules are generally arranged in radial bundles, but since they are not fixed together, the skeleton has generally fallen to pieces in course of fossilization. Thus only a few forms have been preserved entire, e.g. Pachastrella convoluta from the Upper Chalk of Flamborough (Fig. 11 a, b). Detached spicules, on the other hand, are abundant, and form the main constituents of beds of sponge-rock in the Lower and Upper Greensand and in the Upper Chalk.

**Order II.—Lithistida.** Branching secondary spicules (desmas), which may or may not be modified tetraxons, interlock to form a rigid skeleton (Figs. 12, 13); triaenes also may be present (Fig. 13 i). Owing to the firm manner in which their skeletons are built up, Lithistida abound as fossils and are better represented in the Museum than even the Hexactinellida. They are very diverse in form and size, and their spicules also show great variety, some apparently being modified tetraxons, others monaxons, and others polyaxons. It seems therefore that the Order has been derived from both Tetractinellid and Monactinellid ancestors, and its classification is naturally difficult. Pending a satisfactory division into families, the following Sub-Orders are used by palaeontologists:—**Tetracladina** (Cambrian to Recent): desmas four-rayed with their ends produced into twig-like processes which interlock (Figs. 12 d, 13 c, d); *Aulocopium, Siphonia* (Fig. 16 c), and *Jerea* are well-known examples.
Eutaxicladina (Silurian to Jurassic): desma with a thickened central node from which proceed three or more rays, expanded at the ends so as to join others (Fig. 12 h); chief genus, Astylospongia. Anomocladina (Jurassic to Recent): desma a rod with swollen ends from which proceed three or more simple or branched rays, uniting as in Eutaxi-

Fig. 12.—Fossil sponge spicules: Silicispongiae, Tetractinellida. Skeletal spicules of Lithistida. Rhizomorina: a, b, Seliscothon; c, Cnemidiostrum. Tetracladina: d, Aulocopium megamorina; e, f, Doryderma; g, Carterella. Eutaxicladina: h, Astylospongia. a, b, c, d, h, enlarged 40 diam.; e, 26 diam.; f, g, 20 diam. (After Hinde.)

Fig. 13.—Fossil sponge spicules: Silicispongiae, Tetractinellida. Skeletal (a-e) and dermal (f-i) spicules of Lithistida. Eutaxicladina: a, Mastosia. Anomocladina: b, Cylindrophyma. Tetracladina: c, d, Callopegma; e, Plinthosella. e, enlarged 26 diam.; the rest 40 diam. (After Hinde.)

cladina (Fig. 13 b); Cylindrophyma is common in Upper Jurassic rocks. Megamorina (Silurian to Recent): desmas relatively large, curved, branching rods, either intertwining or joined as in Eutaxicladina (Fig. 12 e, f, g); Doryderma (Fig. 16 b) is abundant in the Upper Chalk. Rhizomorina (Cambrian to Recent): desma small, usually elongate, curved, with irregular branches; these end in minute facets, which
Gallery X. are closely apposed to the axis and branches of adjoining spicules so as to form loosely arranged fibres or an irregular meshwork (Fig. 12 a, b); the Cretaceous Verruculina (Fig. 17) is characteristic.

Sub-Class II.—Monactinellida. Demospongiae with monaxon spicules.

The chief modifications of this simple type of spicule are shown in the accompanying figure (Fig. 14 a–q). With these are associated smaller flesh-spicules, serving as grapnels in those forms that have a harder outer skin or cortex (Fig. 14 r, s, t). In these latter, constituting the Order Hadromerina or Spintharophora, the spicules are loosely and irregularly arranged. In the other Order, Halichondrina, the siliceous spicules are bound into a skeleton by the horny substance spongin. Consequently, though the spicules abound in siliceous deposits, chiefly of Tertiary age, complete sponges are rarely preserved. Chalina, Reniera, and the freshwater Spongilla are among fossil representatives of the Halichondrina. The Hadromerina are important to geologists from including the boring Clionidae, whose tubes are often seen traversing fossil shells; indeed they are also active in the disintegration of rocks.

Class IV. (?) Octactinellida.

Silicispongiae, in which the normal spicule has eight rays, of which six radiate in one plane from a common centre, while the other two radiate from the same centre at
right angles to this plane. The Silurian *Astracospongia* is the only genus.

**Class V. (?) Heteractinellida.**

Silicispongiae, in which the normal spicule is large, with from six to thirty rays, radiating variously from a common centre (Fig. 8 a). Contains only the Carboniferous *Tholiasterella* and *Asteractinella*.

Although the spicules are aberrant, it is often maintained that the Octactinellida and Heteractinellida should be placed with the Hexactinellida.

Turning now to the Table-cases with the British fossil Sponges, we find first the oldest known sponge, *Protospongeia fenestrata*, from the Cambrian rocks of St. Davids, S. Wales. All that is preserved are fragments of the meshwork, in which the silica is now replaced by pyrites. They show, however, very primitive triaxon spicules, which enable one to refer the genus to the Hexactinellida *Lyssacina*. Here also are placed the Silurian *Amphiscopnia, Plectoderma*, and *Hyalostelia*. A specimen of the last shows a root-tuft.

Devonian sponges are represented only by the strange *Lodanella*, which certainly has the form of a simple type of sponge, presumably one of the Calcarea, since no spicules are preserved.

The Lower Carboniferous rocks of Yorkshire and Ayrshire have yielded large root-tufts and clearly preserved body spicules of *Hyalostelia*. Among the other spicules here shown, special note should be taken of the Heteractinellid genera, *Asteractinella* and *Tholiasterella*. Lithistid, Tetractinellid, and Monactinellid spicules from Ayrshire, and a Calcisponge from Fife are also exhibited.

The Calcispongiae first appear in force in a series from the Inferior Oolite, Great Oolite, and Corallian; they form massive or stout branching stocks. The marine Jurassic rocks have also yielded Hexactinellid, Tetractinellid, and Lithistid sponges, while from the Purbeck beds comes a freshwater *Spongilla* to represent the Monactinellida.

Among Cretaceous rocks the Lower Greensand of Faringdon, Berks, has long been famous for its beautifully
preserved Calcisponges, of which a fine series is shown. The chert beds of this age in Kent and Surrey are mainly composed of detached spicules of Tetractinellida and Lithistida. A few larger fragments of the latter group from Sevenoaks are the only representatives of the complete skeleton. From the Gault there are only some specimens of *Jereia* and a fine example of the Hexactinellid genus *Craticularia* in pyrites. The Upper Greensand and the Chalk Marl, on the other hand, contain great numbers, which preserve their original form and structure. Those from the Greensand of Warminster, Wilts, first studied by Miss Etheldred Benett of that place, are shown in great quantity. Among the Calcispongiae are good specimens of *Pharetriospongia Strahani* from Cambridge. The folded walls of *Plocoscynthia* are conspicuous among the Hexactinellida, but the Lithistid Tetractinellida are the most richly represented. Among these last, particular attention may be directed to the peculiar lobate forms of the Lithistid *Hallirhoa* (Fig. 16 a) with long stems, the perfect *Siphoniae* from Blackdown.
(Fig. 16 c), the large goblet forms of *Pachypoterion*, and the cylindrical and branching examples of *Doryderma* (Fig. 16 b) from near Warminster.

Next follow the sponges from the Chalk, which, though perhaps more numerous, are not so well preserved. Those from the Upper Chalk of the South of England have had their skeletons almost entirely replaced by iron-peroxide; very frequently also they are now enclosed in nodules of solid flint, in part retaining their forms, but their interior structures are merged in the flinty matrix, and only show the course of the larger excurrent canals and the cloacal cavity. In the sponges from the Upper Chalk of Flamborough, Yorkshire, the form is usually preserved, and also the main features of the canal system, but the spicular structure is now scarcely recognisable. Calcispongiae are represented by *Elasmostoma* and *Pharetrospangia*, but are not numerous except in the case of the little globular *Porosphaera*, which has recently been placed here by Dr. G. J. Hinde. The Hexactinellida all belong to the Sub-Class Dictyonina, and among them the best known is the funnel-shaped *Ventriculites* with its folded walls (Fig. 10), but one should notice also the wide flange-like walls of *Guettardia* and the mushroom-shaped *Ceclopychium*. Among Demospongiae, the Choristid Tetractinellida are represented by *Pachastrella* and *Stelletta* and numerous detached spicules of *Geodia* (?). Lithistida are numerous, and here may be observed the large size and the projecting canal-openings in *Stichophyra* and *Verruculina* from the Flamborough Chalk (Fig. 17). Monactinellida include the borings of *Cliona* in Molluscan shells. These last are the only representatives of sponges in the British Tertiary series.

Passing now to the fossil sponges from foreign localities, exhibited in the Wall-cases, one sees among the Ordovician specimens the large branching Lyssacine Hexactinellid, *Brachiospongia digitata* from Kentucky.

The Silurian of North America has furnished the Monactinellid *Climacospongia* and numerous examples of the Lithistid genera *Astylospongia*, *Palaeomanon*, and *Hindia*. The Lithistid *Aulocopium* is from Gotland in the Baltic.
From this island and from Tennessee comes the noteworthy Octactinellid, *Astraeospongia*.

The Devonian of North America yields casts of *Dictyophyton*, a Lyssacine Hexactinellid.

From the Triassic strata of St. Cassian there is a series of small Calcispongiae, which have been referred to *Stellispongia, Corynella*, and other genera of Pharetrones.

The Upper Jurassic rocks of Württemburg and Switzerland, known as the White Jura, contain a large number of Hexactinellida and Lithistida, some layers being mainly composed of these sponges. They retain their outer form, but, in nearly all, the siliceous skeleton has been replaced by calcite. Examples of numerous genera are exhibited.

Among the Cretaceous sponges, special notice should be taken of the fine specimens of *Coelopistylium* from the Westphalian Upper Chalk, and of the exquisite series of Hexactinellida and Tetractinellida from the Upper Senonian Chalk of Hanover prepared by Dr. A. Schrammen by carefully dissolving out the siliceous skeleton from the calcareous rock: alone among fossil sponges do these present the beauty so frequent in the skeletons of recent forms. Those illustrated in the accompanying Plate III are: (1) *Calyptrella tenuissima*; (2) a species of *Rhizopoterion*; (3) *Sporadoseinia Decheni*, the upper figure showing the interior with the canals, into which open the ostia seen in the lower figure; (4) *Aphrocallistes alveolites*.

Tertiary sponges are so rare that one must not overlook the Calcispongiae, *Typhlocid, Bactronella*, and *Plectroninia*, which are modified Pharetrones from the Eocene of Australia.

Siliceous skeletons of Sponges from the Upper Senonian Chalk of Hanover.
(Slightly enlarged.)

(To face p. 42.)
COELENTERA.

The rest of the Gallery is occupied by the fossil COELENTERA (Hollow-guts). This great division of the animal kingdom, which comprises sea-anemones, corals, jelly-fishes, and their allies, owes its name to the fact that the body-wall of each individual encloses a single hollow chamber (not many as in sponges) which forms the digestive cavity of the animal, and is not traversed by any separate gut such as is found in echinoderms and all the other groups that follow. The body-wall consists of an inner and an outer layer of cells, the "endoderm" and "ectoderm" as they are respectively called, and between them a jelly-like layer, the "mesogloea," which is very thick and noticeable in an ordinary jelly-fish. In a jelly-fish, too, one may readily observe that the organs of the body are regularly arranged on lines radiating from the centre (see Fig. 18). A radiate symmetry of this kind is found in most Coelentera. In many Coelentera the individuals or zooids unite to form colonies.

Setting aside the Ctenophora, which some exclude from the Coelentera, and of which no fossils are known, the Coelentera may be divided into three classes: Hydrozoa, Scyphozoa, and Anthozoa.

The Hydrozoa include the fresh water polyp Hydra, the Hydroid zoophytes such as Sertularia the sea-fir, the corals Millepora and Stylaster, and various forms without fossil allies. All of these have a simple tubular body-wall, enclosing a cylindrical cavity which is not divided by partitions; the mouth is not pushed inwards; the reproductive cells are derived from the ectoderm and are set free directly to the exterior.

The Anthozoa (Flower-animals, also called Actinozoa or Rayed animals) include the sea-anemones, nearly all the stony corals, precious coral, sea-fans, sea-pens, and dead men's fingers. In these the mouth is pushed inwards, so as to form a tube; the body-cavity is divided by partitions, called mesenteries, which consist of mesogloea covered with endoderm; the reproductive cells are derived from this endoderm and are set free into the body-cavity whence they pass to the exterior through the mouth.

The Scyphozoa (Cup-animals) comprise the larger jelly-fishes. Some place them with the Hydrozoa, which they
Gallery X. resemble in the simple mouth and absence of mesenteries; others place them with the Anthozoa, which they resemble in the origin of the reproductive cells. Other features, however, distinguish them from both these Classes.

Class **SCYPHOZOA**.

Many Coelentera being soft-bodied animals can leave no fossil traces except impressions that they may have formed on a sandy shore; these of course can only be left by free-swimming jelly-fishes or medusae, not by fixed forms. Such impressions are actually known in various rocks from the

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**Fig. 18.**—A fossil jelly-fish, *Rhizostomites lithographicus*, one of the Scyphomedusae, from the Kimmeridgian of Solenhofen. *a*, imprint on the exposed surface of the Lithographic Stone, about \( \frac{1}{4} \) natural size. (From E. Ray Lankester’s “Extinct Animals.” After Walcott.) *b*, diagram interpreting the marks seen in *a*. (After Brandt.) *c*, diagram to show how the imprint is formed by the jelly-fish settling down on the mud; a vertical cut has been made through the mud and through the middle of the jelly-fish. (After Walcott.)
Cambrian Epoch onwards, and all appear to belong to a single Sub-Order of Scyphozoa, the Rhizostomata (Root-mouths), so called because the four lips of the mouth are drawn out and each fused by its edges into a tube pierced by small openings through which the animal sucks in its food. Examples of such fossils from Cambrian, Ordovician, and Jurassic rocks are exhibited. Among them a large one from the lithographic stone of Solenhofen, of Kimmeridgian age, called Rhizostomites, resembles that depicted and explained in our Figure 18. Dr. C. D. Walcott has published a full account of fossil jelly-fishes in the Monographs of the United States Geological Survey (1898). Except for these imprints, the Scyphozoa are not known as fossils.

Class HYDROZOA.

Among the Hydrozoa many of the colonial forms are protected by a thin horny coat of chitin secreted by the ectoderm and called “periderm.” In some this covers the branches of the colony and forms little cups or thecae, into which the polyps can be withdrawn; these are the Calyptoblastea (covered buds). In others the periderm does not expand into cups; these are the Gymnoblastea (naked buds). An example of the former is the sea-fir, Sertularia abietina, often cast on our shores. The periderm has a main stem, with branches diverging from each side alternately; the sides of the stem and of all the branches are clothed with a row of cups or thecae, also in alternate arrangement. In each theca lives a polyp, which stretches out twenty-four tentacles and is connected with its fellows by a cord of flesh that passes inside each branch and down the main stem. It is strange that, though capable of preservation, no traces of the chitinous periderm of any such hydroid should have been found in either Cainozoic or Mesozoic rocks. Not until we pass back to early Palaeozoic times, Cambrian to Devonian, do we find organisms that bear any resemblance to the Calyptoblastea. These, which are called Cladophora (branch-bearers) or Dendroidea (tree-forms), have numerous slender forking branches, connected by transverse processes, and bearing little thecae, some for the ordinary polyps, others modified possibly for reproductive cells. Some of the genera, such as Dendrograptus and Callograptus, seem to have been fixed to the sea-floor like a Sertularia. Dictyonema, however, which forms fan-shaped or funnel-shaped colonies, has been
offered to grow up from a sharply-pointed conical cup, the point of which was sometimes attached to a long thread. Such individuals may have been attached to some other animal or plant, but cannot have been directly fixed to the sea-floor. The substance of these fossils is supposed to have been chitinous, like the periderm of Calyptoblastea, when the animals were alive; but actually it is not so. This fact and the absence of any representatives of the group from Devonian to Pleistocene times suggest that, though the Dendroidea may be Hydrozoa, still they cannot be closely related to the Calyptoblastea.

Contemporaneous with the Dendroidea, and, some suppose, derived therefrom, is the Order Graptitoidea—the graptolites, a name given by Linnaeus from the likeness which the fossils bear to writing on the slates in which they are usually found compressed. The pointed conical chamber observed in Dictyonema is also characteristic of all graptolites, and is called the sicula (Fig. 19). In the growth of a colony, a theca (with its contained polyp, be it always understood) budded out from the sicula; and from this theca another budded; and each of these by further budding gave rise to a long line of thecae connected by a common canal. A single branch of such a simple form looks like the blade of a fret-saw, with a straight back, and the thecae forming the teeth, which are directed away from the sicular end. The sicula was almost certainly attached to something, and it is probable that in such forms as these it was attached to floating sea-weed; sometimes it was fixed by a small disc-like expansion; sometimes it hung by a long flexible thread. The earlier graptolite colonies branched many times, e.g. Bryograptus; but the number of branches was gradually
reduced, e.g. Loganograptus, Tetrograptus, until there arose a
form with only two saw-blades to a single handle, e.g. Didymo-
graptus. If these hung from sea-weed the polyps would be
mouth downwards, a position obviously ill adapted for
securing food-particles showered from the floating weed.
This was remedied in two different ways. In one way, the
two blades, instead of hanging down like a pair of tongs,
gradually opened until at last they were directed upwards.
In the other way, the polyps of the first two thecae seem to
have stretched upwards and so to have bent their thecae in
that direction; the rest followed suit, and grew upwards
along the thread from which the sicula hung, e.g. Diplo-
graptus. Thus was formed quite a different type of graptolite.
Further developments took place in these latter: the colony
acquired a median supporting rod or virgula; this ended
often in a disc, which, it is supposed, was hollow, and served
as a float. The colonies were often compound, and many vir-
gulae with their thecae were attached to a single disc. If the
supposition that such forms took to a free-floating existence
be correct, we shall understand why succeeding forms, pre-
sumably their descendants, should have had their periderm
formed of a meshwork of fine strands, e.g. Retiolites: this is a
well-known way of obtaining lightness without loss of strength.

Such are the main lines of graptolite evolution up to
their abrupt end at the close of the Silurian Epoch. But
there were many subsidiary lines; and all these, combined
with the wide distribution of each successive form owing to
its floating life, have rendered the graptolites of enormous
value to the geologist in determining the succession of layers
in great thicknesses of rock, and in tracing those layers over
a large extent of country, even when much disturbed by
later earth-movements. A monograph of the British species
is in course of publication by the Palaeontographical Society.

Fossils that can with certainty be referred to the
Hydrozoa are very few, and are not older than Cretaceous.
They are confined to forms in which the ectoderm secretes
outside the polyp or zooid a number of small calcareous
rods; these grow together into a solid mass, leaving tubes in
which are the polyps. Most of these forms are generally
called Hydrocorallines, lately divided into Stylasterina and
Millepora, of which only the latter are found fossil; and,
indeed, Millepora itself is not known in rocks older than
some Pleistocene raised beaches. The Eocene fossil Axopora
is but a doubtful ally.
The Order **Gymnoblastea** includes both solitary and colonial forms; among the latter the family Podocorynidae contains genera whose roots form an incrustation, somewhat similar to that above described, but without tubes for the polyps; these generally coat gastropod shells. Fossils referred to *Hydractinia*, one of these genera, are found in the Pliocene and Upper Cretaceous rocks of England and elsewhere.

There are a number of fossils that present a general resemblance to the massive Hydrocorallines or the encrusting *Hydractinia*, and, in the absence of convincing evidence for or against, they are usually placed with the Hydrozoa. Such are the thick lamellar masses of *Ellipsactinia* and *Sphaeractinia* from Upper Jurassic rocks, the spheroidal *Parkeria* from the Cambridge Greensand and its ally *Stoliczkaria*, and some thin incrustations from the Chalk.

Apparently allied to some of these forms are the **Stromatopores**, of which a large series is shown in polished slabs of Devonian marble, and a smaller series from the Silurian, both of Britain. There are also several foreign specimens. These fossils usually began by incrusting some solid body, such as a shell, and often grew to enormous size, forming huge banks and reefs. They are composed of calcareous laminae, separated by distinct intervals across which run vertical or radial pillars. Both pillars and laminae are usually traversed by minute irregular canals. The whole mass is in some genera, e.g. *Stromatopora*, perforated by larger tubes, divided by horizontal platforms or tabulae; in these probably lived the principal polyps. The chief work on these obscure fossils is a monograph by H. A. Nicholson, published by the Palaeontographical Society (1886–1892).

**Class Anthozoa.**

Among living forms of this Class, a larger proportion have a skeleton capable of preservation in the fossil state. Recent Anthozoa are divided into two Sub-Classes: I. **Alcyonaria**, with 8 mesenteries and 8 tentacles, which bear short branches. II. **Zoantharia**, with 6, 12, 24, or a larger number of mesenteries and an equivalent number of tentacles, which are simple.

The **Alcyonaria**, with few doubtful exceptions, live in colonies, and the polyps are held together by a jelly (mesogloea). Cells wander from the ectoderm into this jelly,
and, by secreting spicules of a horns substance impregnated with carbonate of lime, give it greater consistency. The common 
Aleyonium digitatum (dead man's fingers) has such a skeleton; it falls to pieces on the death of the animal, but isolated spicules have been found fossil. Sometimes the spicules become tightly wedged together and form a compact skeleton which cannot be disintegrated; the precious coral 
(Corallium rubrum) is the hardest skeleton so formed, but has only been found fossil in some of the later rocks, and that rarely. The organ-pipe coral (Tubipora) is also well known, but is not found fossil. Some of the Gorgonacea form an axis partly horns and partly calcareous; and some calcareous segments referred to Isis, and a few other forms have been found in Tertiary and Cretaceous rocks. A horns axis supports the colony of the sea-pen, Pennatula, and in some allied forms this axis is partly calcified; a few such axes have been detected in Triassic and later rocks. In the living Heliopora, the skeleton is formed, not from spicules developed in cells, but from lamellae of calcite crystallised in an organic matrix produced by the disintegration of ectoderm cells. This genus is found fossil back to the Albian Age. With the exception of a supposed Triassic Pennatulid, fossils that can with certainty be referred to Aleyonaria are confined to rocks of Upper Cretaceous and later age; and yet, as was the case in Hydrozoa, there are a number of Palæozoic genera that resemble many of the recent forms.

In the Sub-Class zoantharia, the only living forms with a skeleton capable of fossilization are several genera differing a good deal in their structure but conveniently grouped together as Madreporaria. Several of them might be roughly described as sea-anemones with a skeleton. This skeleton is quite different from that of the Aleyonaria, except perhaps Heliopora. It consists of crystalline carbonate of lime secreted by special ectoderm cells, not within the cells but outside them, and outside the whole of the ectoderm. The skeleton of the Aleyonaria is internal, that of the Zoantharia is external, as is the shell of a snail. When the young coral-embryo settles down on the sea-floor, it deposits a layer of skeletal substance between its skin and the sea-floor, forming a plate, which soon is turned up at the edges like a saucer. The soft body of the coral may be wholly supported within the saucer, or it may pass beyond its rim. In either case the rim is still built up, and at the same time laminae of lime stretch out from it to the centre
Gallery X. of the floor, each lamina corresponding to a space between two mesenteries; then the lamina itself is built up and so forces the skin inwards into the said space. Thus the skeleton takes the shape of a cup or calyx, partly divided by vertical partitions called septa. Other growths from the wall or from the floor of the cup may be formed in like manner. These are: a single spike or column in the centre, called the "columella"; columns between the ends of the septa and the centre, called "pali"; ridges outside the wall, corresponding with the septa inside, and called "costae." As the calyx grows upwards, the polyp is sometimes pulled away from the bottom of it, but does not therefore stop the secretion from its skin; if a small piece only is pulled away from one side, the skin builds here an oblique partition called a "dissepiment;" if the whole base is pulled away, it deposits a horizontal or saucer-shaped, or sometimes funnel-shaped, partition called a "tabula." Such is the general structure of a solitary cup-coral. Corals may form colonies, either by the repeated budding of such a single form or by its dividing down the middle into two, each half again dividing, and so on. This process of fission, as it is called, is sometimes incomplete, and so arises a form like the Brain-coral, in which the cavities of the polyps and of the cups remain connected in serpentinous grooves (Fig. 24). There is much the same difficulty in connecting Palaeozoic genera with the Madreporaria of Tertiary and recent date, as we have already seen attaching to other groups of Coelentera. Similarly the modern reef-builders, Madreporidae and Poritidae, first appear in Tertiary rocks. Consequently the classification of the Order is far from settled. By means of the skeleton it is possible to divide the genera into three groups: Aporosa, Fungacea, and Perforata. The Aporosa are so called because the calyx-wall and the septa are not perforated by canals, and in colonial forms the polyps are either separate or connected only by superficial canals. The Fungacea include all forms like Fungia, whose skeleton has so many septa that it looks like the under side of a mushroom; the long thin septa of these forms are strutted by short cross-bars. The name Perforata is given to corals in which the skeletal substance is porous throughout, and the polyps of a colony are connected by deep-seated canals. It is not pretended that these groupings indicate relationship. Another method of division is according to the arrangement of the septa. In later Madreporaria these conform to the
general Zoantharian type, and are in multiples of 6. In many Palaeozoic corals they are in multiples of 4. The former have therefore been called **Hexacoralla**, and the latter, **Tetracoralla**. The wall of the Tetracoralla is often wrinkled, so that they have also been called **Rugosa**.

Owing to the difficulties of classification, all the fossil Anthozoa are placed together and arranged under the chief stratigraphical divisions, the British and foreign specimens being separated as usual.

Beginning with the oldest, we find **Ordovician corals** from the Llandeilo rocks of North Wales and the Stinchar Limestone of Scotland. *Streptelasma* belongs to the Zaphrentidae; *Lyopora* is allied to *Favosites*. These fossils are ill-preserved, and the genera will be better studied later.

The British **Silurian corals** are mainly from the Wenlock Limestone. First come a number of genera placed in the family Favositidae, though not all with equal reason. *Favosites*, the honey-comb coral, itself consists of five- or six-sided tubes, set closely together; each tube is divided by flat tabulae, and its cavity is in connection with that of its neighbours by small pores in the wall. *Syringolites*, a North American genus of the same age, has a similar structure, but its tabulae are funnel-shaped (Fig. 20). Some regard these genera as related to *Alveopora*, a modern perforate madreporarian. Others compare them with *Syringopora*, also shown in the Silurian series, but better studied in its

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**Fig. 20.**—A Favositid Coral, *Syringolites huronensis*, from the Silurian of North America. A, a fragment of a colony, natural size. B, a single calyx enlarged 8 diameters; the tabula bears tubercles in radiating lines, like the beginnings of septa, and is sharply bent downwards near the centre. C, a single tube split longitudinally and enlarged 6 diam.; shows the funnel-shaped tabulae. D, a single tube seen from the outside and enlarged 6 diameters; shows the mural pores that connect the cavities of adjoining corallites. (After Hinde.)
Carboniferous species (Table-case 5). Here the tubes are separate, but connected by cross-canals, and the tabulae are funnel-shaped (Fig. 22, b). The living alcyonarian, Clavularia, has a tubular skeleton with similar cross-canals, and the organ-pipe coral, Tabipora, has tabulae either flat or funnel-shaped and cross-canals running in the flat expansions that connect the tubes; therefore many place Syringopora and the Favositidae with these Alcyonaria. In all these genera the tubes of each colony are of equal size, and doubtless contained equally developed polyps. On the other side of this Case is Heliolites (sun-stone), in which the surface shows openings like little suns surrounded by smaller circular openings; in section the colony is seen to be formed of tabulate tubes of two sizes. Heliolites and its allies are explained by reference to Heliopora (see Table-case 1), a living Alcyonarian, in which the larger tubes contain complete polyps, and the smaller ones contain simple sacs of the common flesh of the colony. Halysites, the chain-coral (Fig. 21 a, b), consists of tabulate tubes, flattened, and joined by their edges, with no connecting pores; in some species all tubes are of equal size, in others some tubes are much smaller; it is probably an Alcyonarian. Aulopora, apparently an ally, grows in a low network over shells and corals (Fig. 21 c). A slab of Silurian limestone from Gotland, between Wall-cases 5 and 6, is largely composed of Favosites and Halysites.

In the next Case are Silurian Zoantharia Rugosa or Tetracoralla. Here come the conical Omphyma with root-like supports (Fig. 22 a), the broadly spreading cups of
Ptychophyllum, the disc-shaped Palaeocyclus, the massive colonies of Strombodes. In the next compartment is Cystophyllum, in which the lower part of the calyx is divided by dissepsiments into a vast number of tiny chambers or cysts. In the allied Goniophyllum and Rhizophyllum the calyx was closed by a movable lid. On the other side of the Case is Cyathophyllum, some specimens of which show several young budding from the calyx of the parent. Last is the allied colonial form Acervularia.

Of the Devonian corals, the Favositidae fill the rest of Case 7. Here is Pleurodictyum, frequently associated with a worm, as better seen in the specimens from the Eifel; a similar association is common in Heliopora. These and the Devonian Zoantharia are mostly of the same genera as the Silurian corals; but they grew more luxuriantly in reefs, which now form the massive limestones of Eifelian age in South Devon. Here the specimens have to be studied in polished sections, some of which are of great beauty; one may particularly note Pachypora cervicornis in the black
limestone of Newton Bushel, Smithia Pengellyi in the pink and grey reefs of Barton, and the massive Cyathophyllum helianthoides in the red rocks of Torquay and Plymouth. In the corner of Case 6 are specimens of Calceola sandalina, better shown in the Eifel series; this is allied to Rhizophyllum, and, from its triangular shape and large lid, closely resembles some brachiopods (Fig. 23).

The Carboniferous corals are the last of the Palaeozoic type. The foreign collection includes specimens from Arctic America, from Tournai and Visé in Belgium, and several brought from the Ural Mountains by Sir Roderick Murchison. Among the British Favositidae are the curious little Palaeacis and a good series of Syringopora. Close by is Monilopora crassa growing on crinoid stems, which become deformed in the attempt to grow over it. Chaetetes, which follows, is referred with doubt to the Aleyonaria. On the other side of the Case come Amplexus and Zaphrentis, showing marked bilateral symmetry in the arrangement of their septa, such as obtains in the mesenteries of the recent Zoanthidae, which have no skeleton. In Amplexus the septa are reduced in size, but tabulae are strongly developed. There follow several genera of Madreporaria Aporosa: the cylindrical Campophyllum, which may grow to a great length, as shown by specimens from Weston-super-Mare; Lonsdaleia and Lithostrotion, whose prismatic tubes build up large colonial masses, as those of the familiar Lithostrotion basaltiforme and of L. irregulare from Fermanagh; Dibunophyllum, of stratigraphical importance, shown also on a fine slab from
Durham; *Cyathophyllum*, largely represented in the Table-case, and by some polished slabs of *C. regium* from Bristol; *Olsiphyllum* and *Aspidophyllum*, which end the series. Henceforward these Tetracoralla disappear from the rocks; but the living Astracid coral, *Moseleya*, has when young a four-rayed symmetry which indicates relationship to the Cyathophyllidae. We may therefore suppose that the latter forms were the ancestors of the Astraceidae.

Neither **Permian** nor **Triassic corals** are represented in the British series, but there are shown specimens from the Zechstein, as well as the Klipstein collection from the Triassic beds of St. Cassian in the Tyrol.

The next coral fauna represented in the British islands is that of the Lias, the earliest of the **Jurassic corals**. This

![diagram](image)

Fig. 24.—Zoantharian Corals of Bajocian Age, from the Inferior Oolite of England. *a*, “*Latomeandrea*” Flemingi. *b*, *Montlivaltia trochoides*. (From Prestwich’s “Geology.”)

shows a great change; all the Palaeozoic genera have given place to normal representatives of existing families and genera, such as *Iasatraea* and *Montlivaltia* (Figs. 24 b, 25 b). The collections from the Inferior Oolite are richer, and the representatives of modern genera are increased by Fungiains, such as *Thamnastraea*, and also by a doubtful species (Fig. 24 a) representing the confluent Astraeids, in which the polyps and calyces are incompletely separated as in the Brain-coral. In the rocks of Bathonian age are many corals of similar type, the chief reef-builder being *Calamophylla radiata*. In the Corallian rocks true reefs are formed of *Thecosmilia*, *Thamnastraea*, and *Iasatraea*, of which large specimens are shown (Fig. 25). The structure of all these Jurassic corals, as of the succeeding Cretaceous and Tertiary genera, can be gathered from the diagrams placed in the Table-cases. An interesting series is that of *Iasatraea oblonga*,

**Gallery X.**

- Wall-case 4.

**Wall-case 3.**

**Table-case 4.**

**Table-case 3.**

**Wall-case 3.**
from the Portlandian of Portland and Tisbury, showing how the skeleton of the coral has been converted into chert in varying degrees.

Fig. 25.—Zoantharian Corals of Upper Corallian Age, from Wiltshire. 

a, Isastrea explanata; b, Thecosmilia annularis. Natural size. (From Prestwich's "Geology." ) (See Table-case 2.)

Table-cases 1 & 2. In the Cretaceous Epoch, corals were scarcer in England, for the conditions were less favourable to their growth. The faunas from the lower rocks are in Table-case 2, those from the upper rocks in Case 1. From the Lower Greensand comes Holocystis elegans, once regarded as the only Rugose coral of later age than Palaeozoic. In spite, however, of its four-rayed symmetry, it is now regarded as a normal Astraeid.

In the Gault and Chalk, the principal corals are small, simple forms, for the mud of the former, and the cold depths of the latter sea were fatal to reef-builders. The commonest type is conical in shape, such as Smilotrochus and Parasmilia (Fig. 26 c); some specimens of the latter have been split and show the structure of the calyx clearly. Occasionally
Parasmilia is elongate and cylindrical. Onchotrechus serpen tinus is as narrow and sinuous as a worm-tube. Other forms are discoid, such as Cyclocyathus Fittonii from the Gault, Trochocyathus Harveyanus from the Cambridge Greensand, and Microbacia coronula from the Chalk. Gorgonians are represented in the Chalk by Axogaster cretacea. The series of corals from the Upper Greensand of Haldon in Devonshire, recently enriched by the Vicary Collection, contains many specimens described by P. M. Duncan in a Monograph of the Palaeontographical Society. Here may be noticed Haldonia, Placosmilia, and excellent specimens of Heliopora. The Foreign Jurassic and Cretaceous collections chiefly illustrate the faunas of Central Europe.

Among the British Cainozoic corals the Eocene and Oligocene fauna is separated from that of Pliocene age. Here first appears the family of reef-building Zoantharia Perforata, the Poritidæ (Fig. 27 c), of which a Catalogue by

![Fig. 27.—Perforate Zoantharian corals from the Lutetian of Bracklesham. a, Turbinolia Dixoni; b, Dendrophyllia dendrophylloides; c, Goniopora Websteri. Natural size. (From Prestwich's "Geology.")](image)

H. M. Bernard has been published. Stephanophyllia, Dendrophyllia (Fig. 27 b), and Balanophyllia are also Perforata. The adjoining Stylocenia is an Astraed. Then come the characteristically modern Turbinolidae (Fig. 27 a), and the branching Oculina. Alcyonaria are represented by the axis of a sea-pen named Graphularia Wetherelli, and by the gorgonians Websteria and Mopsea, allied to Isis. The Pliocene Crags have yielded a few small corals of no great interest. The foreign series of Cainozoic corals includes representatives of the Indian faunas and many type-specimens of species from the West Indies.

In illustration of the rock-forming activities of corals there is placed in the centre of this Gallery a Case with specimens of the core obtained by a deep boring on the coral island of Funafuti, and other specimens showing the rate of growth of various corals.
We pass now to Gallery VIII, which contains fossils belonging to several of the great divisions of the animal kingdom. These are arranged in the following order, beginning on the right of the entrance and continuing down the east side of the Gallery, and then crossing to the west side: Echinodermata, Annelida, Arthropoda, Brachiopoda, Bryozoa, and Mollusca (except Cephalopoda). These groups will now be taken in that order.

**Echinoidea.**

**East Side.**

The Sea-Urchin, the Starfish, the Brittle-Star, the Feather-Star, and the Sea-Cucumber, all of which live in modern seas, are examples of this subkingdom. Though differing from one another in outward appearance, they resemble one another and differ from the rest of the animal kingdom in a number of characters which are briefly stated in a label on the wall near the entrance. For further information as to the anatomy and general appearance of forms now living the visitor should consult the Starfish Gallery of the Zoological Department and the Guide relating thereto.

The examples mentioned above represent each one of the Classes into which this subkingdom is generally divided. Taking them in the same order, these are Echinoidea, Asteroidea, Ophiuroidea, Crinoidea, and Holothurioidea. Remains of all these are found as fossils, and there are also at least three Classes, now extinct, and only known from fossils in Palaeozoic rocks. They are called Cystidea, Edrioasteroidea, and Blastoidea.

**Class Crinoidea.**

This Class comes first in the present arrangement of the Gallery. Specimens of some Recent species are placed in the Wall-case close to the entrance, and their flower-like appearance enables one to understand why these beautiful animals should be called by this Greek word, meaning Lily-shaped, and by the popular name Sea-lilies. Though they bear this name, and though many of them have a long stalk and may be fixed to the sea-floor, still they have nothing whatever to do with plants, but are highly organised animals, with a distinct digestive system shut off from the main-body cavity, with nerves, with blood-vessels, and with
ECHINODERMA—CRINOIDS.

a peculiar hydraulic system of water-vessels. The structure of the Crinoidea is illustrated by a series of specimens and drawings, with special reference to fossil forms, and study of this may serve as further introduction to the Echinoderma in general.

The first point to notice is that in crinoids as in all echinoderms, with one or two exceptions, the soft tissues of the animal have the power of depositing crystalline carbonate of lime. This may remain in the shape of minute separate spicules; or the spicules may grow together into a trellis-work, which forms rods and plates. The deposit is usually most abundant in the skin, where it may be built into a continuous skeleton. Often too, spines of the same substance are borne outside the test. This feature, rare in crinoids but characteristic of the sea-urchin, has given to the sub-kingdom its name Echinoderma, which is a Greek adjective meaning “urchin-skinned.”

The chief parts of the Skeleton of a Typical Crinoid are next shown, and are further illustrated by the accompanying figure (Fig. 28). What one may call the body of the animal is confined to the small portion labelled “cup,” on the top of which is the mouth. Since the creature does not move about, it needs some means of bringing food to the mouth, and this is provided by the arms. These are grooved on the inner surface, and water containing the animalculae on which the crinoid feeds is swept down the grooves to the mouth. The stem serves to raise the cup and arms away from the sea-floor and to sweep them through a larger field of food-supply.

Perhaps the Crinoidea are descended from animals that were neither fixed nor provided with a hard skeleton. In any case the result of fixation has been with these creatures, as with so many others, the development of radiate symmetry, caused originally by the food-grooves stretching out from the mouth in all directions. Apparently for mechanical reasons connected with the existence of a hard skeleton, the chief planes of this symmetry have come to be five in number. In other words the skeleton, and to some extent the soft parts and internal organs, can be divided into five similar portions grouped about a central axis. This division into fives or pentamerism, as it is termed, runs right through the Crinoidea and Blastoidae and all the free-moving Echinoderms, although modifications of it arise now and then. It should be understood that the forking of the arms is no modification,
since the arms remain only five however many times each may branch.

The arms serve not only for collecting food but also for respiration. The hydraulic system already mentioned extends
Echinodermata—Crinoids.

along each arm and arm-branch, in the form of a tube or water-vessel, shown in some of the diagrams exhibited. This gives off side-branches to little thin-walled tentacles, which serve as gills and as sense-organs. When touched they withdraw and the groove is closed over by the little covering-plates that are generally present. They are again extended by the pressure of the water in the hydraulic system, and this system is kept full of water by means of openings in the covering or lid of the cup, through which water is swept by minute vibrating lashes (cilia); frequently these openings are confined to a single sieve-like plate, the madreporite. One can understand how this system also was developed in connection with a fixed mode of life. But its importance, like that of the five-rayed symmetry, is due to the fact that it is also found in the free-moving Echinoderms. For these and other reasons it has been supposed that the chief characters of the Echinoderms as we know them in modern seas are due to their descent from a fixed ancestor.

We may now pass on to the general series of fossil Crinoids. The British specimens are grouped under Early Palaeozoic, Devonian, Carboniferous and Permian, Jurassic, Cretaceous, and Tertiary. The foreign specimens are under the same stratigraphical divisions, to which, however, the Trias is added. Some larger British specimens are also in these Wall-cases, and Wall-case 16 contains large slabs from both British and foreign localities.

Most of the British Lower Palaeozoic Crinoids consist of the varied series of forms from the Wenlock Limestone of Dudley. Here one may compare the specimens of Botryocrinus with the restoration (Fig. 28), and may note how pinnules are gradually evolved from simply forked arms. Adjoining are Mastigoocrinus, with its long scourge-like arms, and Thenarocrinus, both with a large extension of the cup-lid looking like a wicker-basket; this is the ventral sac, through which passed the end of the gut. Herpetocrinus is a curious form in which the stem coiled round the cup when the animal was at rest or dead, so that the fossils look like ammonites. In Calceocrinus the arms of one side increase in size while the others gradually disappear, so that the five-rayed symmetry of the cup is also partly lost, and the crown hangs down from the stem, looking like the head of some large-billed bird. Cyathocrinus and Gissoocrinus are simple types, from which Enalloocrinus is not far removed. By the union of the arm-branches in such a form arose Crotalocrinus
with its net-like arms. Then come crinoids in which the food-grooves were covered over and sunk beneath the covering of the cup; these are called Camerata or vaulted crinoids. Among them Pareichocrinus, with its large cup and bead-like stem, is the commonest. In Eucalyptocrinus wing-like processes grew up from the lid of the cup and formed recesses into which the arms were received when folded. Slabs containing many of these crinoids are on the lowest slope of Wall-case 18, and show how in those days many forms belonging to different genera and species lived close together. Here are also several specimens from N. America and Bohemia. The most remarkable is Scyphocrinus, which had an enormous and apparently top-heavy crown; its stem, however, was attached to a large hollow ball (the Lobolith of Barrande) which probably served as a float, so that the crinoid hung crown downwards.

The British Devonian Crinoids are not well preserved. Hexacrinus is the most conspicuous. Some foreign ones are in the Wall-case, but the most remarkable are those from the Lower Devonian slates of the Rhine district; these are preserved in pyrites, and are generally slender forms.

The British Carboniferous Crinoids come mostly from the neighbourhood of Bristol, from Derbyshire, Yorkshire, and Lancashire. Most of these are Camerata, among which Actinocrinus and Amphoracrinius are well-known types. The type-specimen of Actinocrinus loricatus, Schlotheim, is of interest as the first British crinoid calyx ever figured, and as having a longer history than almost any other specimen in this Museum. Described by J. Beaumont in 1676 as a root, it was called the Nave Encrinite by J. Parkinson (1808) and wrongly referred to Actinocrinus triacontadactylus by J. S. Miller. Another common genus is Platycrinus. The crinoids of this epoch are even more abundant in North America, and some exceptionally fine specimens are shown. One may note Gilbertsocrinus with its strange drooping appendages, the spiny Dorycrinus, and Euclidocrinus with a twisted stem like that of Platycrinus.

Triassic Crinoids are not found in Britain, but are fairly abundant in the Tyrol and in Bakony, from both which places the Museum possesses excellent series. The best known form, however, is the Lily Encrinite, Encrinitus liliiformis, from the Muschelkalk of Germany.

Conspicuous among Jurassic Crinoids is Pentacrinus, of
A slab of Lias Shale from Boll, in Wurtemberg, covered with the remains of a large crinoid, Pentacrinus Hiemeri.

[To face p. 63.]
which many magnificent examples from Lyme Regis and elsewhere are exhibited. Here we note how colonies were formed of many individuals of only one or two species, as is the case to-day. A portion of such a colony from the Lias of Boll, in Würtemberg, forms a beautiful picture in the middle of the case (Plate IV). The stem of this form is said to reach a length of 50 feet; a length of 15 feet is certainly common. The length of stem is perhaps to be explained by the fact that many of these Liassic Pentacrinus were attached to floating pieces of wood, and so hung crown downwards. A closely allied form is Isoecrinus, which grows in forests on the floor of some recent seas. Various species will be found among the Triassic, Jurassic, Cretaceous, and Tertiary crinoids. The elegant five-petalled stem-segments of both these genera are washed out of the rocks in many places, and to them the name Pentacrinus (five-lily) was first given by Agricola in 1546. The Pear Encrinite (Apiocrinus) from the Bradford Clay of Wiltshire has the top part of its stem greatly thickened. Near it is Millericrinus Pratti, which exemplifies the tendency, constant in crinoids, to loosen their attachment to the sea-floor and to become free-moving, with a shortened stem. The stem is reduced to a mere knob in Antedon and Actinometra, which, beginning in the Oolites, occur in vast numbers in modern seas. Though unstalked and free-moving when grown up, these crinoids are fixed by a stalk when quite young.

In the Cretaceous Crinoids, Marsupites and Uintacrinus, the stem is entirely lost, and it seems probable that the latter at all events was a free-swimming form. Both genera lived at almost the same time (Upper Senonian) and were widely distributed. Specimens are shown from the English Chalk, and a slab covered with Uintacrinus from North America is placed on the wall.

Tertiary Crinoids are not numerous. The most interesting specimens are those illustrating variation in the stem of Balanocrinus, another Pentacrinid.

Class CYSTIDEA.

The Cystids are of interest partly on account of their rarity, partly because they are all extinct, none having survived the Carboniferous Epoch, partly by reason of their diversity and strangeness of structure, but mainly because they are thought to comprise forms from which other classes
of Echinoderma have descended and to approach most nearly those pre-existing animals from which the Echinoderma were originally derived. The best British specimens are from the Wenlock Limestone; but other horizons are well represented by those from Bohemia, Russia, North America and other foreign countries. Among the older and more primitive are many, such as Aristocystis (Fig. 29 a), that appear to have been little more than plated sacks, without stem or arms:

Fig. 29.—Types of Cystidea. a, Sack form with scattered pores, Aristocystis. b, Example of Rhombifera, with food-groove skeleton slightly developed, Echinosphaera. c, Example of Diploporita, Protocrinus; the brachioles are restored on the right side; elsewhere are seen the facets that supported them. d, Example of Rhombifera, with food-groove skeleton highly developed, and with respiratory folds restricted to three "pectini-rhombs," of which one is shown near the top of the right-hand quarter; Lepadocrinus.

hence the name of the class, which means "sack-shaped." These, however, probably all had ciliated food-grooves stretching from the mouth, either along fleshy tentacles or along similar processes provided with a calcified support or skeleton. These processes are called brachioles, and there is no reason to suppose that they contained such extensions of the body-cavity, of the reproductive organs, or even of the hydraulic system, as occur in Crinoidea. Hence we suppose that a
special opening found in the hard body-wall of these forms served for the extrusion of the eggs; and we infer that the necessary process of respiration was effected by means of the pores which penetrate the plating of the sack or theca. Feeding and breathing are the two processes without which the life of an animal must stop, and the history of the Cystidea, as of most groups of animals, is the history of changes by which these processes were ever better and better carried out. Let us briefly contrast in these respects the two forms of the Crystal-apple, as the Swedes call a fossil so common in their country and found also in our Welsh Ordovician rocks, namely Echinosphaera (Fig. 29 b) and Sphaeronis, or the more advanced Protopcrinus (Fig. 29 c).

The numerous plates of Echinosphaera appear to be joined by fine lines, which represent canals in their substance and are arranged in rhombs. In Sphaeronis and Protopcrinus a number of small ovals are scattered over the plates, and each of these contains two vertical canals meeting near the inner surface. Both of these structures appear to have contained spaces that brought the soft tissues and interior fluids of the animal close to the outer aërated sea-water. In Echinosphaera the food-grooves were borne entirely on separate skeletal pieces; in Protopcrinus they pass over the surface of the plates before reaching the brachioles. Turn next to the Wenlockian Lepadocrinus, of which a reconstruction is here drawn (Fig. 29 d), and note that the canals joining the plates are now visible only in three places, where they are intensified as deep folds; here too the food-grooves are carried far over the surface on a series of special pieces from which the brachioles arise; to keep these away from the mud, the whole body is now raised on a stem. This then is a development of the Echinosphaera type. The other line of evolution leads to a form like Proteroblastus, in which there are five food-grooves passing right down the actual surface of the plated sack or theca, and bordered regularly by brachioles; the double canals are here concentrated on the plates that bear the brachioles. On these two distinct lines of evolution are based the two Orders: Rhombifera (with canals or folds in rhombic pattern) and Diploporita (with canals opening in double pores).

Notable Cystidea, perhaps to be regarded as a distinct Order, are the Anomalocystidae, of which Trochoeytis, Mitrocystis, and Placocystis respectively represent the Cambrian, Ordovician, and Silurian stages of evolution. The two
former are among the Bohemian and French Cystids, the last is well shown among the British fossils.

**Class BLASTOIDEA.**

The Blastoids (Bud-shaped) form a small class of Echino-derma, which arose at an early period, probably from Cystidea Diploporita, and flourished chiefly in Devonian and Carboniferous times. The Museum possesses a rich collection, which served as basis of a monograph written by Etheridge and Carpenter and published by the Trustees. A list of the specimens has also been issued. Since the fossils are too small to be seen clearly, only a few characteristic examples are exhibited. The general appearance of the blastoid skeleton and the terms applied to its more obvious parts are shown in Fig. 30. The brachioles border five food-grooves, of which the skeleton is rather complicated. The contiguous edges of the plates termed deltoids and radials were folded, and in most of the genera these folds projected far into the interior of the theca and thus enabled the aerated sea-water to come close to all the

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Fig. 30.—A typical Blastoid, *Orophocrinus fusiformis*, Carboniferous (Kinderhook) of Iowa, U.S.A.
internal organs; hence these folds are called hydrospires (water-breathers). In most blastoids the theca is borne on a stem and shows conspicuous five-rayed symmetry. In a few forms, however, the theca rested on the sea-floor, and this produced irregularity in its shape with a change in one of the food-grooves; examples of this are Eleutherocrinus, Pentephyllum, and Zygocrinus, which in other respects are dissimilar and not closely related.

The preceding Classes are essentially fixed forms, living with the mouth upwards and obtaining food by means of a current of sea-water swept towards the mouth along ciliated food-grooves. They are therefore termed PELMATOZOA (stalk animals) in opposition to the ELEUTHEROZOA (free-moving animals) such as star-fish and sea-urchins, which live with the mouth downwards and take more solid food into it without the help of ciliated grooves. In the Pelmatozoa hitherto discussed the food-grooves are supported on special structures usually stretched out from the body; and extensions from the hydraulic system may pass along these appendages, but do not penetrate the plated wall of the body (calyx or theca).

**Class EDRIOASTEROIDEA.**

We come now to a Class of Echinoderma, still living in the manner of Pelmatozoa, but with the food-grooves directly floored by thecal plates and in some cases at least with pores between those flooring-plates, giving passage, it is supposed, to extensions from the hydraulic system. These are the Edrioasteridea (sessile star-fish). Some of them, such as Agelacrinus and Lepidodiscus, were probably fixed for the whole of their existence, usually to some large shell. In these genera no pores have been found between the flooring-plates of the groove. Others, such as Edrioaster (Fig. 31) and Dinocystis, could probably shift their positions,
and in them pores are clearly seen. Pores, however, also
occur in the food-grooves of a stalked form, Steganoblastus.
The food-grooves of Edrioaster (Fig. 31) closely resemble
those of a starfish, except that they, as well as the mouth,
are protected by covering-plates like those of crinoids. It is
possible that these curious forms may throw some light on
the origin of starfish.

Since Edrioasteroidea are rare and of exceptional interest,
both British and foreign examples are exhibited together in
Table-case 30, and are supplemented by reproductions.

Class Asteroidea.

On the further side of Table-case 30, the free-moving
Echinoderms begin with the starfishes, generally regarded as
constituting the simplest Class of Eleutherozoa. In a starfish
the body is either markedly five-sided in outline or is more
or less star-shaped, in which case it is said to consist of a
central “disc” extended into “arms,” which vary in number
from 5 (e.g. the common cross-fish, Asterias) to over 40 (e.g.
the sun-star, Heliaster); the mouth is in the centre of the
body and is turned to the sea-floor; the anus is almost in
the centre of the upper surface, but is absent in a few forms;
the under side of the arms is grooved, and along each groove
runs a vessel of the hydraulic system; this vessel gives off
side-branches which end in free processes (podia) differing
from those of Pelmatozoa in that each terminates in a
sucker; between the flooring-plates of the groove are pores,
through which pass branches from the podia, each com-
municating with a swelling (ampulla) within the body. This
arrangement of the podia enables each one to be extended
for locomotion, and to be withdrawn into the groove by the
passage of the fluid from it into the ampulla; such an
arrangement is found in no Pelmatozoa except perhaps some
Edrioasteroidea; but from those forms starfish differ in
having the groove unprotected by covering-plates. The
remainder of the starfish skeleton consists usually of small
plates or bars which serve to strengthen and support the
stout but flexible skin.

The oldest Palaeozoic starfishes in the British collection
are Uranaster and Palaeaster, represented by casts in
Caradoc sandstone of Upper Ordovician age. From the
Wenlock beds comes the heavily plated and many-armed
Lepidaster. The Lower Ludlow shales of Leintwardine,
Herefordshire, have furnished a large number of starfishes belonging to the genera Palasterina (Fig. 32 b), Sturtzaster, Rhopalocoma, and Bdellacoma. Some obscure starfishes have been found in the Lower Devonian slates of Cornwall, but the fossils of this age are better studied in the beautiful series from Bundenbach in Prussia. In these the skeletons are altered into iron pyrites and imbedded in black slate, which has been cleaned away from them with most delicate care by Mr. B. Stürtz, the original describer of many of these specimens (see "Palaeontographica" 1886, 1890, and other papers mentioned on the labels). Palacasteriscus, Cheiroptaster, and Helianthaster may be mentioned, but all are beautiful and interesting.

Returning to the British series, we find some good speci-

![Image of starfishes](image)

Fig. 32.—Palaeozoic Starfishes. a, Schuchertia stellata, from the Ordovician of Ottawa; under surface with grooves and mouth. b, Palasterina primæva, Upper Silurian of Kendal; upper surface.

mens of **Jurassic star-fish**, notably Tropidaster from the Middle Lias, a massive Pentagonaster Sharpi from the Northampton Ironstone, a beautifully preserved Asterias Gaveyi and Solaster Moretonis from the Great Oolite. Astropecten is represented by species of various ages from Bajocian to Corallian, and some larger specimens of it are in Wall-case 17.

**Cretaceous Asteroidea** are best shown in the excellent series from the English Chalk, which have been described by W. P. Sladen and W. K. Spencer in a monograph of the Palaeontographical Society. Here the tesselated Calliderma Smithi and the fine group of Pentaceros bulbiferus are specially worth notice.

The **Tertiary star-fish** are represented chiefly by Pentagonaster and Astropecten from the London Clay, preserved in a pyrites that is regrettably liable to decompose.
The living Brittle-stars, Sand-stars and Basket-fish are separated from the star-fishes as a Class, because the arms are sharply marked off from the central disc, and have the grooves covered over by plates, and the flooring-plates of the grooves fused into a series of ossicles (little bones) like vertebrae, worked on one another by powerful muscles. Thus these arms can serve as limbs for locomotion; and the podia, not being needed for that purpose, usually serve only for respiration and touch. As a further result of this development, the arms no longer contain processes from the digestive and reproductive systems as they do in star-fish. In the Basket-fish the arms may branch, and are used for coiling round the stems of other animals or plants.

The Palaeozoic Ophiuroids do not show all these points of distinction from Asteroids; in many of them the arm-groove is not completely closed, and its flooring-plates are not yet fused into vertebrae. Species found in the Ordovician rocks of Bohemia are still more like Asteroids than any here exhibited. We begin with British Wenlockian forms, such as Lapworthura (Fig. 33) and Protaster from the Lower Ludlow shales. A slightly more advanced type is the little Sympterura from the Lower Devonian of Cornwall. The Ophiuroids of this age, must, however, be studied in the Stürtz Collection from Bundenbach, where explanatory labels are given.

Among the British Wenlockian Ophiuroidea the most remarkable are Eucladzia and its allies; for in them the arms do not extend beyond the disc, but to make up for this the few podia within the disc limits are of great size, and have a flexible armour of small plates.

From Carboniferous to Trias there are no British Ophiuroids, but on the lowest slope of Wall-case 17A may be seen Onychaster, from the Carboniferous rocks of Indiana,
with its coiled arms, a small *Hemiglypha* from the Muschelkalk, and a few other specimens.

Of **Jurassic Ophiuroids**, in the British Lias, the so-called Starfish bed of Pliensbachian age, exposed between Charmouth and Bridport, has yielded *Ophioderma Egertoni* and other well-preserved brittle-stars. Interesting forms have lately been obtained from the Corallian Calcareous grit of Yorkshire. The Foreign Jurassic series consists mainly of some elegant little species from the Kimmeridgian lithographic stone of Solenhofen, belonging to the genera *Geocoma* and *Ophiurella*.

Another *Geocoma* comes from rather similar rocks of **Cretaceous** age in the Lebanon. From the English Chalk there are *Ophioglypha serrata* and other species recently described by Mr. Spencer in the monograph referred to above. By this time, it will be noticed, the genera have quite a modern aspect.

*Ophioglypha Wetherelli*, from the London Clay, is the most important of the British **Tertiary** sand-stars; there is also an *Ophiolepis* from Pleistocene deposits of the Clyde basin.

**Class ECHINOIDEA.**

Owing to their abundance, especially in Mesozoic and Cainozoic rocks, and to the continuous change in structure during geological time, the fossil Sea-urchins, or Echinoids, are of great value to the stratigraphical geologist and of no less interest to the student of evolution.

The differences between a sea-urchin and a starfish have sometimes been illustrated by imagining a starfish with short rays, and therefore with a five-sided or globular shape; then suppose the grooves to grow upwards to the neighbourhood of the anus so that they supplant all the leathery loose-plated skin, except a small area just round the anus; let this area be surrounded by five plates, each pierced by a pore for the passage of the generative products, and one of them also serving as madreporite—then one would have something very like a sea-urchin. But there is an obvious difference: in the starfish the radial water-vessel lies in a groove outside the skeleton; in the sea-urchin there is no groove, but a series of plates flush with the rest of the test, and the water-vessel lies beneath these—that is, within the skeleton—and the podia pass out through pores between or in those plates. Thus the test of a regular sea-urchin is
Gallery VIII.

marked by five areas passing from near the anus to the mouth, and these areas are fringed by the podia so that they look like garden-paths or avenues (ambulacra). Thus it appears that the ambulacral plates, those that constitute these areas, are not the same structures as the flooring-plates of the groove in a star-fish. It would therefore in some respects be simpler to compare a sea-urchin with an Edrioaster in which the covering-plates of the food-grooves had become fixed, leaving passages for the podia, while the flooring-plates had gradually been absorbed; we must also suppose the Edrioaster to have turned upside-down, and its anus and water-pore to have moved to the surface now uppermost.

If now we examine the oldest British Silurian Echinoidea, namely, Echinocystis and Palaeodiscus from the Lower Ludlow shales, we shall observe that the anus has not yet reached the centre of the upper surface, that the ambulacra have not met regularly around either that centre or the anus, that in both genera the test was still flexible with its plates neither fixed in number nor regularly arranged, and that the pores for the podia are often between the ambulacral plates instead of surrounded by them as in later echinoids. Moreover, it has been maintained that some specimens of Palaeodiscus show traces of an inner set of plates corresponding to the flooring-plates of the groove in Edrioaster. These genera, however, had, as our specimens show, a well-formed jaw-apparatus of complicated structure, only a little simpler than that found in a Mesozoic Cidaris or in a recent Echinus. They must, therefore, have roved actively in search of food. The movable spines (radioles) borne by the plates are still small and not very different from those of some Asteroidea and Edrioasteroidea.

The Devonian rocks of Britain have yielded few remains of sea-urchins, but fossils from Germany (e.g. Lepidocentrus) show that, while the test remained flexible, the plates in each interradial area between the ambulacra were arranged in columns, and that often each plate bore one radiole larger than the others. In the British Carboniferous series are genera, such as Palaeochinhus (Fig. 34, 1) and Melonites, that still have the interambulacral plates in many columns; but in Archaeocidaris, or Echinocrinus, these plates are relatively larger (Fig. 34, 2) and are definitely arranged in four columns. In other respects Archaeocidaris closely resembles the earlier forms of Cidaridae found in the Trias of the Tyrol
ECHINODERMA—SEA-URCHINS.

73

and Hungary. In these, as in the later Echinoidea, the interambulacral plates were restricted to two columns, but the test remained flexible.

The British Mesozoic Echinoids include many specimens described by T. Wright in his Monograph published by the Palaeontographical Society. With the Jurassic series we find the beginning of more familiar types. *Cidaris Edwardsi* from the Lias has the jaws preserved, and the radicles of various sizes still on the tubercles to which they were attached by muscles when the animal was alive. Here is also *Acrosalenia*, with specimens of the *A. pustulata*

![Types of Fossil Echinoidea](image)

Fig. 34.—Types of Fossil Echinoidea. 1–5, Regular; 6–8, Irregular. 1, 2, Carboniferous; 3–5, Jurassic; 6–8, Cretaceous. 1, *Palaeechinus sphæricus*, side-view. 2, *Archæocidaris*, interambulacral plate and its radicle. 3, *Cidaris glandifera*, primary radicle. 4, *Hemicidaris intermedius*, side-view. 5, *Salenia petalifera*. 6, *Dysaster ringens*. 7, *Enalaster Grevenovi*. 8, *Catopygus columbarius*. 5–8 are seen from above, with the anus towards the spectator. The figures are all somewhat less than natural size. (By permission of Messrs. A. & C. Black.)

that was found in crowds near Cirencester. The Jurassic fossils continue with the more modified *Hemicidaris* and *Pseudodiadema*. Then come early forms of Diademina, such as *Hemipedina* and *Diademopsis*, those from the Lias being but slightly removed from some of the Triassic species. They give rise, however, to more elaborate forms in the Oolitic rocks, e.g. the large *Hemipedina (Phymopedina) mar- chamensis*, with its numerous tubercles. An interesting allied type is *Pelanechinus*, in which the plates were widened and flexibly united. In this species have been found the little pincer-like appendages called pedicellariae, well known in living urchins, but rarely preserved on the fossils.
All the echinoids thus far mentioned have a circular outline, with the mouth in the centre of the under surface, and the anus at the other pole surrounded by an apical system of plates. Such echinoids are called Regular. In Bajocian time some echinoids seem to have taken to moving generally in a single direction, and as one result the anus passed backwards from the apex, as may be seen in Pygaster. Then the mouth passed forwards, as in Hyboclypeus, and this resulted in modifications of the front ambulacrum and the hinder interambulacrum, e.g. Collyrites. At the same time in many forms the test became elongate, the jaws were lost, and the mouth developed instead a shovel-like lip, for the animal now took to burrowing through ooze and swallowing it on its way. Naturally the radioles became reduced to a coat of small, sometimes almost silky, spines. All such echinoids are called Irregular.

The foreign Jurassic echinoids are placed on the floor of Wall-case 16.

The British Cretaceous Echinoidea contain two distinct faunas—one from the Lower Greensand, and one from the Gault, Upper Greensand, and Chalk. The former is small, but the latter is the most interesting in the British series. Its most striking feature is the predominance of large specimens of Cidaris, of which a fine series of specimens from the Chalk is shown. One may note especially the example [E. 1952] of Cidaris sceptrifera with the apical plates, and those of Cidaris clavigera (33,455 and 39,998) which show the jaws in position and the radioles attached. Following the Cidaridae come the Saleniidae, which have an additional plate in the apical system. The Diadematidae are represented by a large series of forms, of which Cyphosoma Koenigi, from the Chalk, is the best known. The genera Glyphocyphus and Zeuglopleurus are the forerunners of sea-urchins with pitted tests, such as Temnopleurus. The most interesting specimens are those of an Echinothuria, a genus with imbricating plates, carrying further the type of structure begun in Pelanechinus and brought to a high development in Phormosoma, Asthenosoma, and other genera now living in the abysses of the ocean.

In the Irregular Echinoids of Cretaceous age the gradual change of form and ornament that takes place in all groups of fossils as they pass up through the rocks has of late received careful study. By these mutations geologists are able to recognise successive layers in the thick mass of
Conulus, Hagenovia, and these are particularly interesting in this respect. The species of Micraster and Echinocorys are worthy of attention are Discoidae, of which one specimen (40,341) shows the internal processes that serve for the attachment of the jaw-muscles, here much modified; Conulus [Galerites], in which the jaws have been lost; the unique specimen of the curiously-shaped Pygurus lampas from the Upper Greensand of Lyme Regis; Hagenovia and Infusaster, which in their elongate shape approach the modern deep-sea genus Pourtalesia.

The foreign Cretaceous Echinoids are partly in Wall-case 16 and partly on the lowest slope of 15 c. A specimen of Hemipneustes striato-radiatus from Belgium, mounted on a block on the top shelf of 15 b, is the largest sea-urchin in the collection.

The British Cainozoic Echinoidea are fewer and smaller than those of the Mesozoic Era. The Eocene specimens particularly are dwarfed and stunted in comparison with those that lived in Southern France at the same time. The Pliocene specimens from the Crags of East Anglia are larger and more numerous. Among these Temnechinus Woodi is represented by two forms, one of which has depressions at the upper ends of the interambulacra; these are supposed to have been for the reception and protection of the young, since several recent sea-urchins protect the brood in a somewhat similar manner. In addition to the ordinary North Atlantic forms, the Crag fauna contains various sea-urchins of West Indian type, such as Rhynchopygus Woodi, Agassizia equipetala, and Echinolampas subrostrata, and this implies a direct connection of warm shallow sea between the two regions.

The foreign Cainozoic Echinoids include a number of type-specimens from Malta and Australia. A series of the large Clypeaster from the Mediterranean basin and the West Indies is mounted on blocks on the top shelf. Two large specimens of Chelonechinus, a genus allied to Cystechinus which now lives in the ocean abysses, are of particular interest: one is from the radiolarian marls of Barbados, the
other from the soap-stone of Fiji, and they have been held to prove that those rocks were raised from great depths since the Miocene Epoch.

**Class Holothurioidea.**

The Sea-cucumbers, which form the last Class of Echinoderma, have no continuous skeleton, and are represented as fossils only by the spicules and minute plates deposited in the skin. These have been found so far back as in rocks of Carboniferous age. Spicules of Cucumaria from the Pliocene beds of St. Erth, Cornwall, and plates of Psolus from Scotch Glacial beds are exhibited.

An upright case in the middle of the Gallery contains a series of specimens intended to illustrate the importance of **Echinoderms as Rock-formers.** The back, or west side, of the case contains a single polished slab of Mountain Limestone full of stems and other fragments of Carboniferous crinoids. On the front of the case is a large slab of Silurian limestone from Gotland, with masses of crinoid stems showing on its weathered surface. Above this are samples of rock from various parts of the world, composed entirely or in great part of the skeletons of crinoids, of cystids, of blastoids, and of echinoids. The free-moving echinoderms, however, do not form so large a proportion of any rock as do the fixed forms. The latter often compose masses many feet in thickness and affording excellent building-stone.

The latest comprehensive account of Echinoderma, including fossil forms, is in Volume III. of "Treatise on Zoology," edited by E. Ray Lankester (London, 1900).

**Annelida.**

Among the numerous and diversely built forms of life that popular phraseology lumps together as worms, only the segmented or ringed worms have left in the rocks traces that can be identified by the palaeontologist. These worms constitute the group Annelida, and among them again it is only the Class **Chaetopoda** (bristle-feet) and, with few exceptions, only one Order of that Class, namely the **Polychaeta** (many-bristles), with which we have to deal. These animals are nearly all marine, and at any rate have
no representatives among freshwater fossils. They are all soft-bodied animals, and the only portions capable of preservation in the rocks are the bristles, used for locomotion, and the horny jaws. The bristles of course can only be identified when connected with other traces of the animal. The jaws, some of which in the fossil state were long known as conodonts, are so minute that they can as a rule only be found by the washing and microscopical examination of the softer rocks. Evidence for the former existence of the free-moving forms, known as Polychaeta Errantia, may also be furnished by impressions, borings, trails, or worm-castings, the last-mentioned being the mud passed through the animal's body for the extraction of food and then excreted in coiled heaps (Fig. 37). These traces are somewhat unsatisfactory, and many have been vaguely assigned to "worms" which are now believed to have been formed by other animals, such as arthropods or molluscs. Though it may sometimes be convenient to give them names, it must be remembered that this implies no knowledge of the animal to which they may have been due. The most abundant fossils assigned to these Annelida are the hard tubes which the sedentary forms build up, sometimes from sand-grains stuck together, sometimes of carbonate of lime deposited in layers by the skin. These, however, show so little characteristic structure or even shape that it is difficult to be sure that they were always formed by animals related to the modern makers of similar tubes—the Polychaeta Tubicola. Moreover, since simple tubes are fashioned by some other kinds of animals, for example boring molluscs, one cannot even be certain that all these fossils are due to polychaetes. In spite of these difficulties, fossil "worms" have some interest for the geologist, since many of them are sufficiently distinct to enable him to identify stratigraphical horizons by their means, while others have left their remains in such quantity as to build relatively large masses of rock, and others again throw light on the conditions under which the rock wherein they occur was deposited.

The obscurest fossils of any that have been referred to Annelida are, as might be expected, also the oldest. They are, in fact, the oldest traces of life in the Museum, and come from rocks believed to be of Pre-Cambrian age at Loch Fyne in Argyll. Some large slabs presented by the eighth Duke of Argyll are in Wall-case 8 of Gallery XI.; smaller specimens are with the other British Annelida in the present
Gallery. With these latter are other obscure fossils, named Salterella and Arenicolites from the Basal Quartzites of Lower Cambrian age in Sutherland. All these are supposed to have been worm-burrows like those of the living lob-worm Arenicola, but they may equally well have been made by plant-roots. Some supposed tubes of Serpulites from the same Quartzites have a slightly better claim to an annelid origin. On the top shelf of the Wall-case is another Arenicolites from the Upper Cambrian of Wisconsin; in this the burrows are seen to stop short at the level of successive layers of rock, as though the animals had been killed off, either by a period of drought or by the sudden deposition of a thick coating of sand. A similar form called Scolithus comes from the Potsdam Sandstone of the same age near Ottawa; the bit of rock exhibited on the bottom slope of the Wall-case shows over 80 burrows on a surface no bigger than a man’s hand.

Among British Silurian specimens may be noted the large Serpulites longissimus, of which a tube curved in an almost complete circle is at the back of the Wall-case, the small coiled Spirorbis (Fig. 36), and the ringed tubes of Cornulites often found in clusters. The similar tubes of Ortonia are attached to shells and such-like objects (Fig. 35). From the Lower Devonian of Cornwall come some peculiar...
bodies called *Nereitopsis*, from a supposed resemblance to *Nereis*, one of the Errantia; these fossils, however, if polychaetes at all, belonged to the Tubicola. The mass of tubes of *Serpula advena* from Caldy Island is of Upper Devonian age. Tubes referred to *Serpula* also come from the Carboniferous limestones of Ireland. *Spirorbis* (Fig. 36) is frequently associated with plant-remains from the Coal Measures; but some of the fossils to which this name is given are probably mollusc shells. The jaws and conodonts found in all these Palaeozoic rocks are too small for exhibition.

The *Jurassic* rocks contain many different shapes of tubes, referred to *Serpula* when adherent to other objects, to *Ditrypa* when free, and to *Galeolaria* when in dense clusters. Noteworthy examples of the first-mentioned are the tubes on a piece of jet from the Lias of Whitby, and those on the surface of a coral, *Thamnastrea*, from the Coral Rag. Similar tubes are common in Cretaceous rocks; the Serpulite Limestone of Brunswick is formed of them. More interesting are the tubes formed of fish-scales found in the English Chalk, originally described by Mantell as a fish—*Muraena lowesiensis*, and referred by Agassiz to another fish—*Dercetis elongatæ*. One worm seems to have burrowed into sponges, and is found as an irregular spiral coil inside hollow flints. In the Wall-case the most interesting Mesozoic fossils are those from the Kimmeridgian of Solenhofen. Here is a *Eunicites* with bristles and jaws, and numerous coiled castings known as *Lumbricaria* (Fig. 37), and supposed to be those of worms, though once regarded by Agassiz as derived from the intestines of fish. *Pyrgopolon* is a characteristically shaped tube, common at the top of the Cretaceous.

In the British Tertiary series, *Serpula heptagoma*, from the Barton Beds, should be noticed, for the opercula which closed the tubes are preserved, and thus the genus can be more accurately determined. The coiled tubes known as *Vermetus bogoriensis*, of which two fine slabs are exhibited in the Wall-case, are thought by some to have been formed by a mollusc. A large block of *Serpula tenuis* comes from the
Oligocene of the Isle of Wight. The *Eunice* found in a nodule of Pleistocene age from Greenland shows how well the Errantia can be preserved when circumstances are favourable. The rarity of such fossils proves once again the extreme imperfection of the geological record.

**ARTHROPODA.**

Next to the Annelida there are displayed the fossil remains of the Arthropoda, that great group of the animal kingdom which includes insects, centipedes, lobsters, barnacles, spiders, scorpions, and a host of less familiar forms. These animals have no internal skeleton, but the body is enclosed in a case made of a horny substance called chitin. In the island of Oesel thin films of this substance, the former clothing of a marine arthropod, have been preserved without change since Silurian times. Sometimes lime salts are deposited in the chitinous envelope, and render it even more fit for preser-
vation. The case is not like a solid box, but is divided into a number of segments, separated as a rule by softer flexible skin. In primitive forms the whole body is divided into a series of generally similar segments, each bearing a pair of limbs; but in later forms several segments fuse more or less completely, especially at the head end of the body. The limbs also are segmented and their segments united by flexible joints, whence the name Arthropoda (jointed feet). A more striking feature, however, is that though, in their essential structures, all these limbs are organs of locomotion, some at the front end of the body, around the mouth, are used for seizing and biting food: the feet have become jaws. In most arthropods that live in the water some limbs behind the jaw-limbs have developed plates or plumes, which serve as gills. Land arthropods breathe either by small lung-sacks or by long tubes called tracheae, which open to the air by holes, called stigmata, in the sides of the body-segments.

The great majority of arthropods now living are divided into the following Classes: Insecta, including flies, butterflies, beetles, and bugs; Chilopoda or centipedes; Diplopoda or millipedes; Crustacea, including crabs, lobsters, sand-hoppers, wood-lice, barnacles, and water-fleas; and Arachnida, including spiders, ticks, and scorpions. All these Classes are represented by numerous fossils back to Palaeozoic times; but many fossil arthropods are not obvious members of any of these Classes. Such are the trilobites, the Eurypterida (*Eurypterus, Pterygotus, &c.*), and the king-crabs, which last have persisted to our own day. Certain resemblances between these forms have led some writers to unite them in a single Class. It is now generally admitted that the king-crabs and Eurypterida are related to the Arachnida; but they may still be conveniently distinguished as Merostomata. The trilobites were perhaps allied to the Merostomata, and yet there are some features in which they resemble Crustacea. It may therefore be as legitimate as it is convenient to keep them apart as a Class Trilobita. These Classes will now be considered in the order in which the British specimens are arranged in the Table-cases, namely: Trilobita (Cases 25, 24); Arachnida (Cases 24, 23); Crustacea (Cases 23–20); Diplopoda, Chilopoda, and Insecta (Case 20). Setting aside the centipedes and millipedes, this order may be justified as that in which the Classes successively became dominant during geological time.
Class **TRILOBITA.**

The structure of the Trilobites (three-lobed forms) is illustrated by models. *Dalmanites caudatus* (Fig. 38) shows the ordinary view, namely, the back or upper side of the horny coat. Two grooves running lengthwise divide it into three lobes—a middle or "axial," and two side or "pleural" lobes. The principal organs of the body lie within the axial lobe, while the pleural lobes are expansions of this for the protection of the limbs. At first glance the coat appears made of several (about thirty) segments, gradually decreasing in size and importance from head to tail. Inspection shows, however, that these segments are more closely united into three regions: the head-shield, composed of five segments; the thorax, of eleven segments; and the tail-shield or pygidium, of about fourteen segments. The segments of the

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**Fig. 38.—Diagram of a Trilobite, Dalmanites [Phacops] caudatus, Wenlockian.** Where the black lines are drawn across is a definite suture or joint. The five segments of the head-shield are numbered.
head-shield are the most closely united, and are of the same number in all trilobites. Those of the thorax are flexibly joined, and in other genera vary from two to twenty-nine. The pygidial segments, which also vary much in number, are fused into a single shield, from which all traces of the original segmentation has sometimes disappeared. The axial portion of the head-shield, called the glabella, is separated by the dorsal or axal furrows from the pleural portions or cheeks, and is divided by side-furrows into lobes representing the five primitive segments. The part of each cheek next the glabella is firmly attached to it or “fixed.” From this the rest of the cheek is separated by the “facial suture,” and is “free.” The free cheeks, which in many genera are separate, here meet and fuse in front of the glabella. They bear the compound eyes, which stand up like semi-circular towers bearing facets on their outer walls. In many trilobites the eyes are not so conspicuous, and in some they seem to have been altogether absent. The hinder corners of the head-shield, called the genal angles, are here produced into spines. These are part of the fixed cheeks; but in some trilobites, where the free cheeks take a larger share in the head-shield, they are part of those cheeks. Each segment of the thorax consists of an axis and two pleura. The axis stretches forward beneath the axis of the segment in front, forming an articular surface. Each pleuron has a groove running obliquely from front to back and connected with the occasional overlap of the pleuron in front as the animal moved. Not all trilobites have the pleura grooved, but some have them also bevelled at the front corners, forming facets on which the hinder angles of the pleura in front could play; for these trilobites could roll up like a wood-louse. The pygidium is jointed on to the last thoracic segment; its shield has a smooth border, produced into a tail-spine.

The under side of the trilobite, with its appendages, has been made known chiefly by the labours of C. D. Walcott and C. E. Beecher. A restoration of Triarthrus, a genus of the family Olenidae, is exhibited (Fig. 39). The edge of the carapace is turned over on the under side and supports a thin membrane, in which are the two openings of the digestive tract: the mouth, beneath the glabella, and the anus, beneath the pygidial axis. In front of the mouth is the hypostoma or fore-lip, behind it is the metastoma or hind-
Specimens of other species showing these structures are exhibited. Every body-segment, except that in which the anus opens, bears a pair of appendages, attached to transverse thickenings of the ventral membrane. The front pair form whip-like antennae. The remaining pairs are branched, one branch being a crawling leg, the other branch bearing a fringe of bristles or of lamellae. The basal segments of the four pairs on the head served to bite food and to pass it into the mouth. The lamellate branches of the remaining limbs may have served partly for swimming, partly for breathing.

Trilobites lived only in the sea, some on reefs, some on muddy or sandy bottoms; some, it is inferred either from the absence or the extraordinary size of the eyes, in deep water. In the growth of an individual trilobite of simple structure, the free cheeks and the eyes borne by them are at first not seen on the upper surface of the head-shield. As the animal grows they appear at the edge, and gradually come to occupy more and more of the upper surface. Some early trilobites, however, such as Agnostus (Fig. 40 a), Harpes, and Trinucleus, never reach this stage, and may be separated as a Grade Hypoparia (under-cheeks) from those in which the free cheeks are visible on the upper surface. In these latter the free cheeks may be confined to the fore-

![Reconstruction of a Trilobite, *Triarthrus Becki*, from the Ordovician, Utica Slate of New York; natural size. (After Beecher. Table-case 25.)](image-url)
part of the shield, as in Calymene, Staurocephalus (Fig. 40 c), and Phacops (Fig. 38), or they may stretch right back so as to include the genal angles, as in Olenus (Fig. 40 b), Paradoxides, Triarthrus (Fig. 39), Ogygia, Bronteus, and Aculaspis.

Trilobites with the latter character are called Opisthoparia (back-cheeks), while those with the free cheeks in front only are called Proparia (front-cheeks).

Trilobites are found well developed in the oldest Cambrian rocks, being represented by Hypoparia, which soon die out, and by Opisthoparia. Three genera of the Opisthoparian Family Olenidae, Olenellus, Paradoxides, and Olenus, have given their names to the Lower, Middle, and Upper Cambrian Ages. With the Ordovician arise the Proparia, and the Trilobita as a whole attain their acme. Through Silurian and Devonian times they gradually decline in numbers and size, till in the Carboniferous Epoch only a single family remains, of which one genus, Phillipsia, struggles on to the Permian.

The British trilobites are allied to those of Scandinavia and Russia, rather than to those of Bohemia and the rest of Europe and eastern North America. They have been described mainly by J. W. Salter and H. Woodward in the Monographs of the Palaeontographical Society, and many specimens there figured are in the national collection. The large Paradoxides from the Middle Cambrian of St. Davids will attract notice. Angelina Sedgwicki from the
Tremadoc Slates is the usual text-book instance of how fossils may be distorted by earth-movements. The Lower Cambrian Hollybush Sandstone of Comley, Shropshire, has yielded fragments referred to the wide-spread genus *Olenellus*. *Trinucleus*, with its ornamented head-shield, is a characteristic Ordovician genus. *Asaphus tyrannus* and *Ogygia Buchi* are common in the Llandeilo Flags of the same Epoch. *Iliaenus*, with its smooth head-shield and pygidium, has even lost the axal furrows from the thorax, and forms a strong contrast to the spiny *Acidaspis* or the tuberculate *Encrinurus* of Wenlockian age. *Galymmene Blumenhachii* is "the Dudley Trilobite," and to its coiled individuals is due the name of the genus ("covered up"). *Homalonotus* is another form devoid of ornament and losing its furrows. In *Sphaerexochus*, *Deiphon*, and *Staurocephalus* (Fig. 40 c), the swollen glabella is a remarkable feature. Among Devonian trilobites note the fan-shaped pygidia of *Bronteus* and the tripartite head of *Trimeroccephalus*. The three Carboniferous genera *Griffithides*, *Phillipsia*, and *Brachymetopus* are well represented.

In the foreign collection one may notice slabs from the Cambrian of China covered with "petrified swallows," as the Chinese call these remains of *Stephanocaris*, *Drepanura*, and *Agnostus* (Plate V). Here are fragments and a restoration of the huge *Asaphus megistos* from the Ordovician of Ohio. The Bohemian collection obtained from J. Barrande is particularly valuable. Among the Cambrian genera one should note *Ptychoparia* and *Sao*; among the Ordovician, *Olenus* (Fig. 40 b), *Isotelus*, the large-eyed *Aeglina*, the deeply incised pygidia of *Areia*, and *Calymmena Tristani* which marks a horizon of Llandeilian age. *Calymmena senaria* is the species in which Walcott discovered the appendages by means of cross-sections. Of Silurian genera, *Proetius*, *Arethusina* and *Harpes*, from Bohemia, should be noticed. The Devonian series includes large pygidia of *Bronteus* from Bohemia, and several trilobites from South Africa.

**Class ARACHNIDA.**

In this Class, as already explained, we include the
**MEROSTOMATA.** First in this division comes the Order
*Eurypterida*, whose structure may best be studied in the exhibited models of *Eurypterus* and *Hughmilleria*, as well as
Cambrian Trilobites from China.

[To face p. 86.]
in the large specimens of *Pterygotus, Slimonia* (Fig. 41), and *Stylonurus* placed on the wall. The body is long, flattened and segmented, and is divided into two regions. The front one consists of the head and some body-segments fused under a single shield. The hinder region, or abdomen, consists of twelve distinct movable segments, the last of which bears a movable tail-plate or spine, the telson. On the upper surface or back of the front shield is a pair of small simple eyes (ocelli) near its centre, while a pair of compound, facetted eyes is placed on or near the front margin. On the under surface of the shield is the mouth, with a plate in front of it and another behind it (*b* in Fig. 41), as in the trilobites. In front of the mouth is a pair of limbs, differing greatly in size in different genera, but always ending in pincer-claws (chelae). At the sides of and behind the mouth are five pairs of limbs, variously modified in different genera for crawling, swimming, or grasping, but always agreeing in having the basal segment toothed to serve as a jaw. Thus there are six pairs of limbs to the front region, which is therefore supposed to consist of six body-segments. On the under surface of the first abdominal segment are the openings of the reproductive glands, covered by a

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![Fig. 41.—Restoration of the under surface of a Eurypterid, *Slimonia acuminata*, by M. Laurie. *b*, metastoma; *c*, leaf-like gills seen through the covering plates of the mid-body; *g*, telson; *i*-*vi*, appendages of the fore-body; *vii*-*viii*, segments of the abdomen, *vii* and *viii* covered by the genital operculum, *ix*-*xii* bearing gill-covers. Much smaller than nature. (From Woods' "Palaeontology," by permission of the Cambridge University Press.)
paired plate with a median process, the whole known as the genital operculum. This appears to correspond to two segments. The four following segments bear each a somewhat similar plate, to which were probably attached leaf-like gills. The six remaining abdominal segments have no appendages, except the telson. The surface of the chitinous envelope usually bears a scale-like ornament. The Eurypterida are first found in Ordovician rocks, and attained their maximum in both numbers and size about the beginning of the Devonian Epoch, when they seem to have frequented shallow waters and lagoons; they are found in the Coal Measures under circumstances indicating a brackish or freshwater habitat; the last survivor is associated with land-plants of Permian age. The British fossils belong chiefly to the genera Eurypterus, Slimonia, and Pterygotus. The remains of the last-mentioned, from the Old Red Sandstone of Scotland, are large and conspicuous objects, widely known through the writings of Hugh Miller. In the Silurian rocks of Oesel in the Baltic smaller species of Eurypterus and Pterygotus occur in a beautiful state of preservation. The great Stylourus and the smaller Hugh-milleria lived in North America during the Devonian Epoch.

Next come fossils of the Order Xiphosura (sword-tails), of which Limulus, the king-crab, is the living representative (Fig. 42). Here the fore-part of the body is proportionately much larger, and is covered by a domed shield of horse-shoe outline. Near its middle line is the pair of ocelli, and further back on each side, about halfway from the margin, is a compound eye. The hind part of the body is, in Limulus, covered by a single shield, with six spines at each side and with grooves on its back indicating that it is composed of certainly six segments and perhaps more. This is separated from the front shield by a strongly marked flexible articulation, and the bayonet-shaped telson is jointed to it behind. The under-surface of the fore-part has a central mouth surrounded by appendages, which scarcely differ from those of Eurypterida beyond the removal of the sixth pair from a share in biting the food. The six segments of the hind-part carry paired plate-like appendages, as in Eurypterida, the first forming the genital operculum, the rest bearing gills on their hinder surfaces. Limulus then differs from the Eurypterida mainly in the fusion of the abdominal segments and their reduction from twelve to six. In the very young Limulus, however, there are nine such segments,
not yet fused, and there are among the older fossils of this Order many that show a similar or greater approach to the Eurypterid plan. The first of these exhibited is the Silurian *Neolimulus*, with at least nine free segments; then *Hemiaspis*, in which the last three are narrower than the others and are followed by the telson. *Belinurus* from the Coal Measures has eight abdominal segments, of which the last two or three are fused; while in the contemporaneous *Euproöps* [*Prestwichia*] the segments are reduced to seven, and these are fused. If the Coal Measure fossils known as *Cyclus* are not larval stages of the contemporaneous Xiphosura, one can only say that they are just what one would expect those larvae to have been. The resemblance of all these early Xiphosura to trilobites is also too striking to be overlooked. *Limulus* itself first appears in the Trias; several specimens from the Solenhofen stone of Kimmeridgian age are shown.

As the Eurypterida were assuming a fresh-water existence before vanishing, the Order *Scorpionida* was making its appearance, being first represented in the Silurian rocks by what seems to have been an aquatic, if not actually a marine form. This is *Palaeophonius* (Fig. 43), found in both Scotland and Gotland. It consists of the same number of segments, arranged in the same way as those of Eurypterida, and
bearing similar appendages. The first two pairs of these have strong pincer-claws, the next four pairs are stout and end in a single claw, whereas in later scorpions they are thin and end in a double claw. The genital operculum is on the seventh segment, and on the eighth the appendages have been modified into a pair of organs corresponding to those which in later scorpions have a toothed edge and are known as pectines (combs). It is possible that the breathing organs on segments nine to twelve remained as in the Eurypterida. But in \textit{Eoscorpius} of Carboniferous age an important change has taken place in that the covering plates have closed over the lamellae of the gills, leaving only slit-like openings called stigmata. Thus when the animal emerged from the water the lamellae remained moist, and

Fig. 43.—Silurian primitive Scorpions, \textit{Palaeophonus}. \textit{a}, \textit{P. nuntius}, Ludlovian of Gotland, upper surface. \(\frac{3}{4}\) nat. size. (R. I. Pocock, after Thorell & Lindström.) \textit{b}, \textit{P. caledonicus [Hunteri]}, Ludlovian of Lanarkshire, under surface, about twice nat. size. (R. I. Pocock.) (Both blocks lent by Messrs. Constable, from Lankester's "Extinct Animals."
breathing took place by the admission of air to them through the stig mata. They are no longer gills, but lungs. Specimens of Eoscorpius from England, and a fine scorpion from Bohemia called Cyclophthalumus are exhibited. Later scorpions differ in no important respects from Eoscorpius.

Adjoining the British fossil Eoscorpius are some other Carboniferous Arachnida, mostly found on splitting open nodular concretions of ironstone that occur in the Coal Measures of Staffordshire and Lancashire. The most interesting genera are Anthracosiro and Eophrynus (Fig. 44), which belong to a group called Anthracomarti, apparently ancestral to the recent Pedipalpi (whip-scorpions and allies) and Opilio nes or harvest-spiders, and in some respects intermediate between them. The hind-part of the body is still

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**Fig. 44.—An example of the Anthracomarti, Eophrynus Prestvicii, from the Coal Measures of Dudley.** A, upper surface; B, side view of front shield, the front end being to the left; C, under surface. 1–11, appendages of the fore-body, the base of II bearing a biting process, mx; car, front shield with eye-pit, o; pr. st, sternal plates on its under surface; the upper (tergal) and under (sternal) segments of the hind-body are lettered 1–10tg and 1–9st respectively, that lettered gen. tg in A corresponds to the generative openings on the under side. Enlarged about 2 diameters. (From R. I. Pocock, Geological Magazine, 1902, by permission of the Editor. Table-case 23.)
articulated to the fore-part, and consists of flexibly joined segments. As in typical Arachnida the mouth is at the front of the body, and only the first two pairs of limbs take part in biting; the four other pairs are solely for walking. The breathing organs were probably lung-books, as in Scorpions and Opiliones, but there is still little evidence on this point. The Araneae or true spiders are also found among Carboniferous fossils; some from Bohemia are shown. These and other orders of Arachnida are, however, not very richly represented in the rocks until Cainozoic times, when they are met with in the Oligocene of the Isle of Wight and of Florissant in Colorado, in the lignites of Rott near Bonn, the Miocene of Oeningen in Baden, and in the Baltic amber.

Class CRUSTACEA.

These are almost all dwellers in water, breathing by gills. Their outer chitinous envelope is more often thickened by lime than is the case in the Classes previously described, and this crustaceous nature combines with their habitat to render them fairly common fossils. The annexed figure of a fossil lobster (Fig. 45) shows that in the more typical forms the envelope is composed of segments, in which an upper and under half are clearly distinguished. Several front segments are joined together and covered by a shield, part of which projects backwards as a carapace, and from the number of limbs borne by this part of the body the number of segments may be estimated; in the lobster, for instance, it is thirteen (not all preserved in the fossil figured). The six remaining

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**Fig. 45.**—A typical Crustacean; one of the Macrurous Decapods, Glyphea regleyana, of Oxfordian age. a–b, head; b–c, thorax, covered by the carapace; c–d, abdomen, bearing a telson, e, and tail-fins, f; g, eye; h, i, limbs of the head; k–o, walking limbs. Two-thirds nat. size. (From Woods' "Palaeontology," by permission of the Cambridge University Press.)
segments form the abdomen, at the end of which is the telson. In the number of the body-segments, in the manner in which they are united, in the limbs which they bear, and in the extent to which some of them are covered by a carapace, the Crustacea manifest much diversity. They may be divided into five Sub-Classes: Branchiopoda, Ostracoda, Copepoda, Cirripedia, Malacostraca.

The **BRANCHIOPODA** (gill-feet) are represented among fossils by only one Order, the **Phyllopoda** (leaf-feet), which owe their name to their numerous leaf-shaped gill-bearing limbs. In the large number of body-segments and the uniform nature of the limbs, as well as in some points of internal structure, modern phyllopods are considered to approach the supposed ancestors of Crustacea. The carapace may be absent, as in *Branchipus*, or may form a large shield over the fore-part of the body, as in the living *Apus*, or may be folded down the middle line, and the two halves held together by a muscle crossing them inside, as in the little *Estheria*. All recent phyllopods are essentially inhabitants of fresh water, but since they survive the change of this into salt water and even brine, as in the Great Salt Lake, their presence in any rock does not necessarily imply its fresh-water origin. *Protocaris* from the Cambrian of Georgia resembles *Apus*, and the doubtful *Anomalocaris* from the Cambrian of British Columbia has been compared to the abdomen of *Branchipus*. In Devonian shales *Estheria* and its relatives are common fossils. In Triassic and a few later rocks *Apus* is met with, and from the Oligocene of the Isle of Wight there is a *Branchipodites*, here exhibited.

The **OSTRACODA** (Fig. 46, 5) resemble *Estheria* in having a carapace of two valves united along the back by a membrane, and capable of being tightly closed by an internal muscle. The appendages, which differ from those of the Phyllopoda, can be thrust out between the lower edges of the valves, and used for creeping or swimming. Appearing already in Cambrian times, the Ostracoda rapidly became numerous in both individuals and species, and still abound in the seas and lakes of to-day. They live usually in shallow water, and occur in such crowds that their remains sometimes form considerable beds of rock. *Leperditia*, *Beyrichia*, *Thlipsura*, *Entomis*, and the four families which they represent, are all Palaeozoic. *Cytherella*, *Cypridina*, and the relatives of *Cypris*, range from Ordovician to modern times. *Cythere* and its family are Mesozoic to Recent. Most ostra-
cods are too small for satisfactory exhibition, so that only a small selected series is shown.

The **Cirripedia**, or barnacles and their allies, long the delight of folk-loreists, are of special interest to British naturalists as having formed the subject of two monographs, on the recent and fossil forms respectively, by Charles Darwin. They have also a peculiar interest as representatives of a free and actively moving group of animals that, so soon as their early wandering days are over, settle down to a fixed existence, becoming permanently attached by the fore-

![Fig. 46.—Fossil Crustacea.](image)

Fig. 46.—Fossil Crustacea. 1, a Brachyuran Decapod of the tribe Dromiacea, *Dromilites Lamarcki*, London Clay. 2, a Brachyuran Decapod of the tribe Oxystomata, *Palaecorystes Stokesi*, Gault and Upper Greensand. 3, a Macruran Decapod of the tribe Eryonidea, *Eryon arcticiformis*, Solenhofen Stone. 4, a Macruran Decapod of the tribe Loricata, *Mecochirus longimanus*, Solenhofen Stone. 5, an Ostracod, *Cypridea tuberculata*, Wealden. 6, a Cirripede, *Loricula pulchella*, Turonian Chalk; the specimen was figured by C. Darwin. All figures except 5 and 6 are considerably less than natural size.

part of the head to rocks, shells, drift-wood, ships, and the like. Some, as the common *Balani* or acorn-shells of our coasts, are closely and immediately attached to the rock or wood; others, as the barnacles, hang from a long stalk. In either case they develop a calcified shell composed of a number of definite pieces. Six pairs of feathery cirrus-like limbs, to which the Sub-Class owes its name, stretch out from the shell, and continuously sweep food-particles to the mouth within the shell. This return to a mode of feeding characteristic of creatures fixed to one spot is accompanied, as in all such creatures, by a tendency towards radiate symmetry. In spite of this remarkable modification, Cirri-
pedes differing but little from advanced modern types are already found in the Silurian; these are *Pollicipes* and *Scalpellum*. More common during that Epoch was *Turritela*, in which the body was covered with from four to six columns of overlapping scales. A still older form, *Lepidocoela*, had but two columns of scales. Not very unlike these is *Loricula* (Fig. 46, 6), found in Cretaceous rocks, which furnish many other genera. Most of the exhibited series, however, is from the Pliocene Crags of East Anglia.

The rest of the fossil Crustacea belong to the Sub-Class **MALACOSTRACA** (soft shells), an old name originally intended to distinguish these “shell-fish” from those with hard calcareous shells. There are nineteen body-segments, of which eight form the thorax, and six, rarely seven, form the abdomen. Most of the larger and better known modern Crustacea fall within this Sub-Class. Into the unsettled question of their classification we shall not here enter, but merely allude to those Orders or other groups that are represented by fossils. First comes the group **Phyllocarida**, in which are doubtfully but conveniently placed a number of Palaeozoic Crustacea, which may or may not be related to the recent *Nebalia*. These have over the head and thorax a large shield, which may be folded as in the Phyllopoda, and may bear a narrow beak-like plate loosely joined to it in front. The abdomen consists of ring-like segments (seven in modern forms), and the telson has side-spines. In *Hymenocaris*, from the Cambrian rocks of Wales, the shield is in one piece. In *Caryocaris* from the same rocks it is bivalve, as also in the Ordovician and Silurian *Ceratiocaris*, which was sometimes two feet long, and in the Devonian genera, *Echinocaris* from North America and *Aristozae* from Bohemia. In *Rhinocaris*, from the Devonian of New York, a third plate arose in the middle of the back between the two valves. *Discinocaris* and allied forms, ranging from Ordovician to Trias, had an almost circular divided shield, much like the brachiopod shell *Discina*, while *Aptychopsis* and others have been confused with the similarly shaped opercula of ammonites (compare Fig. 83).

There have long been known from Carboniferous and Permian rocks some genera differing greatly from their contemporaries and placed in a division **Syncarida**. These are now considered to resemble and to be related to a remarkable Crustacean called *Anaspides*, which lives in fresh-water pools near the top of Mt. Wellington, Tasmania.
This animal has the segments of the thorax all distinct, covered by no carapace, and bearing limbs with swimming branches and plate-like gills; its eyes are stalked as those of a lobster. Among British fossils the Syncarida are represented only by poor specimens of *Palaeocaris*, also found in the Coal Measures of Illinois. To show the characters more plainly, a specimen of *Uronectes* [Gampsonyx] from the Lower Permian of Saarbrücken, Germany, has been placed in the Table-case.

Several Crustaceans now referred to independent Orders of Malacostraca were till recently united under the name **Schizopoda** (cleft-feet). Since certain fossils, placed originally in the same supposed Order, are too obscure to be referred with certainty to any other position, the division Schizopoda is provisionally retained for them in the exhibited series. Most are from the Coal Measures, and among these *Pygocephalus*, *Palaeocaron*, and *Anthrapaemon* are the better known forms. Some of these outwardly resemble the Decapoda, but appear to have some thoracic segments still unfused with the carapace.

Next come fossils referred without doubt to the clearly defined Order **Isopoda**. Of this Order the most familiar representative is the wood-louse, but most isopods are marine and some inhabit fresh water. The flattened body has a small head-shield (not a carapace) to which are flexibly joined seven thoracic segments, bearing each a pair of walking legs; plates attached to the bases of these limbs form a brood-pouch for eggs and young; the abdomen, which bears gill-plates, is reduced in size, its segments partially fused, with a relatively large tail-shield. Here are shown fragments of the large *Prearcturus*, from the Devonian near Hereford, whose isopod nature may be questioned. *Cyclosphaeroma*, however, from the Jurassic rocks of Northampton and Solenhofen, is an undoubted
isopod, as also is the little *Archaeoniscus Brodiei* found in quantities in the Purbeck Beds of Wilts and Dorset. Several forms are found in Cretaceous and Tertiary rocks, as *Palaega Carteri* (Fig. 48) in the Cenomanian Chalk of Dover, and *Eosphaeroma Smithi* in the Eocene of the Isle of Wight. One tribe of Isopoda, the Epicaridea, live as parasites on other Crustacea, notably on prawns, causing distortion of their carapaces. The distorted carapace seen in some specimens of *Palaecorystes*, a crab from the Cambridge Greensand, suggests that they harboured these parasites.

The Order **Amphipoda** contains the sand-hoppers and fresh-water shrimps; small animals with a body flattened from side to side, and with gills attached to the thoracic feet. A few have been found in Tertiary rocks, but are not represented in the Museum.

The Order **Stomatopoda** comprises but a single family, the Squillidae, of which all living representatives are marine. Resembling lobsters in general form, they differ in having the carapace so short as to leave the hinder segments of the thorax uncovered, in having none of the thoracic limbs modified as jaws, but the first five pairs bearing pincer-claws which are especially large on the second pair, and in having gills borne only on the limbs of the abdomen. *Squilla*, well-known in modern seas, is found in the London Clay, in the Cretaceous of Lebanon, and in the Kimmeridgian of Solenhofen. *Necroscilla*, based on an abdominal fragment from the Middle Coal Measures of Derbyshire, is placed in this Order provisionally.

The large Order **Decapoda** (lobsters, prawns, crabs) owes its name (ten feet) to the fact that the hinder five pairs of thoracic limbs are strongly developed as either walking or swimming legs or as pincers (k-o in Fig. 45); gills attached to these limbs are covered by the carapace. Three pairs of
limbs in front of these are called maxillipeds, because they assist the mandibles and the two pairs of maxillae in the work of jaws; and the two pairs in front of all these act as feelers. The fossil Decapoda belong to two Sub-Orders, the Macrura (long-tails) and the Brachyura (short-tails). These represent two grades of structure, the former being the older; and the most interesting among fossil Decapoda are those that cast light on the evolution of the short-tails from the long-tails, or, as one may put it broadly, the change of lobsters into crabs.

The British series of fossil decapods is arranged under the time divisions: Jurassic, Cretaceous, and Tertiary; and within each of these divisions the Macrura precede the Brachyura. True decapods are first found with certainty in Lower Triassic rocks, but these are only represented in the foreign series.

We shall now take those tribes of Decapoda that are found fossil, in an order corresponding approximately to that of their appearance in the rocks. This order agrees with an arrangement according to grades of structure, the most simple and primitive coming first.

In the foreign Trias and in the Lower Lias of England is found a long-tailed genus Aeger, in which the first three pairs of thoracic legs bear pincer-claws as in the lobster, but here the third pair is much the largest. For this reason and because of its general form, Aeger is held to be an ancestor of the tribe Stenopidea. It is also found in the Kimmeridgian lithographic stone of Solenhofen.

A large prawn, common in the Mediterranean and called Penaeus, differs from true prawns in having the first three pairs of thoracic legs all much of a size and all, as a rule, with pincer-claws. The tribe Penaeidea of which this is typical is also supposed to be represented in the Trias. The early fossils, however, are rather doubtful, and it is in the Solenhofen stone that we first certainly meet with Penaeidea in the genera Atrimpos, Acanthochirus, Bylgia, Drobona, Dusa, and others. Penaeus itself is found in the Senonian rocks of Westphalia. A few examples of the tribe occur in Tertiary strata.

The true prawns and shrimps, which with their allies form the tribe Caridea, have pincer-claws on the first two pairs of legs, and have the side-plates of the second abdominal segment broadened so as to overlap those of the segments in front and behind. Owing to their comparatively tender
outer covering they are not common as fossils. Udorella of the Cretaceous of Lebanon is probably an ancestral form derived from the Penaeidae. The common British prawn and the river-prawns of the tropics belong to a family Palaemonidae. Some fossils of this family found in the Oligocene Osborne Beds of the Isle of Wight are exhibited.

Among the more conspicuous crustaceans in the Liassic rocks of England is the genus Eryon, of which the first representatives occur in the foreign Trias. These are the earliest examples of the still extant tribe Eryonidea, in which the first four pairs, and sometimes all five pairs, of legs bear pincer-claws. The existing genera are mainly confined to the deep sea, and, like many other deep-sea animals, are blind. It is, however, plain that the fossil Eryon lived in much shallower waters. The genus is found also in the Solenhofen Stone, which was apparently deposited in a lagoon (Fig. 46, 3). Its latest species is of Neo-comian age.

Next we notice the many fossils of Glyphaea (Fig. 45) representing the extinct family Glyphaeidae. These are of much interest as the ancestors of all the recent Loricata, a tribe represented in modern seas by the Palinuridae and Scyllaridae. The Palinuridae are familiar through the rock-lobster, the langouste of the French (Palinurus vulgaris); the body is more or less cylindrical, and the antennæ are long, cylindrical, and jointed. In the Scyllaridae the body is flattened, and the antennæ are expanded into broad plates, which are said to be used as shovels in burrowing. All these lobster-like forms may be distinguished from true lobsters by the absence of pincer-claws, though in the female the last pair of legs has them imperfectly developed. Now the Glyphaeidae have antennæ still of primitive form, with a stalk of many segments none of which are joined to the upper lip-plate as they are in recent Loricata, and with a feeler-portion only moderately developed; the legs have no regular pincer-claws, but there is a tendency towards their formation in the first pair, which is larger than the rest; they all have a small pointed rostrum. Along with Glyphaea we find Soaphenus and Preatya in the Lower Lias of England. These were preceded by Pemphix, of which there are specimens from the foreign Trias, and were followed by Pseudoglyphaea in Liassic and Oolitic rocks, by the long-limbed Mecochirus of Oxfordian and Kimmeridgian age.
Gallery VIII. (Fig. 46, 4), and by *Meyeria* in the Lower Cretaceous series. The earliest representative of Palinuridae appears to be *Palinurina* from the Lower Lias of Lyme Regis. Following this in Upper Cretaceous beds and in the English Eocene is *Podocrates* [*Thenops*], scarcely to be distinguished from *Linuparus* now living in Japanese waters. *Cancerinus*, also a Solenhofen genus, possibly led to the Scyllaridae, which are represented in the English Gault and London Clay by *Scyllaridia*, and in the Chalk by *Scyllarus*.

The true lobsters and crayfish are examples of the tribe Astacidea. These forms have pincer-claws on the first three pairs of legs, and the first pair is very large. Already in the Lias we meet with *Eryma*, which is also found with *Pseudoastacus* in the Solenhofen stone; these two, especially the latter, are very like the freshwater crayfish. In the Chalk, *Enoplocyti* is fairly common and strikingly lobster-like. But *Hoploparia*, found in Cretaceous and Tertiary rocks, is even more closely allied to the modern lobster (*Homarus*) and its near relation *Nephrops*, of which the Norway lobster is a familiar example.

*Callianassa* is a characteristic genus of the tribe Thalassinidea, burrowing forms, with a soft, loosely built body. In recent species of the genus the end segments of the first leg with its pincer-claw are greatly enlarged and flattened for shovelling; but this is only in one leg of the pair. From the Kimmeridge Clay of England comes *Callianassa isochela* in which this flattened claw is not so disproportionately enlarged and is found in both legs of the first pair; the preservation of the abdominal segments in this fossil suggests that they were not so thin-skinned as in later forms. In Cretaceous and Tertiary rocks the characteristic claws are found, but not the abdominal segments as a rule.

We come now to the Brachyura. The typical crabs of the present day differ from the decapods thus far described in the following characters among others: the abdomen is short and so bent up under the body that it is quite or almost invisible from above; its last segment bears no tail-appendages; there are at most nine pairs of gills; the maxillipeds of the third pair are broad and flattened, so as to cover the other mouth-parts; the front feelers are set in cavities formed by partitions that connect the front margin of the carapace with the hard parts of the under surface; the whole body is rarely longer than broad. There is, however, a primitive tribe of Brachyura, the Dromiacea, in which these characters
are not developed so constantly or to such an extent. Thus, the outline of the body is longer, more of the abdomen can be seen from above, its last segment often bears small tail-appendages, and in the female the first abdominal segment also has traces of a pair of limbs, the gills are more numerous, the cavities for the first feelers are not so clearly defined. In these characters the Dromiacea approach the Macrura, and it has been shown that they more particularly resemble the true lobsters. A peculiarity, not derived from the lobsters, is that in the Dromiacea the last pair or two pairs of thoracic legs are turned on to the back, where, by their hooks or claws, they hold a bit of sponge or some such object, under which the animal is completely hidden. It is to this tribe that most of the early fossil Brachyura probably belong, although this cannot be determined with certainty, because the tender-skinned abdomen is very rarely preserved. A starting-point for the tribe is furnished by Prosopon, British specimens of which from the Great Oolite are here exhibited. First found in Bajocian beds, it persisted to Neocomian times. The carapace, which alone is known, is elongate and closely resembles that of Homalodromia, now living in the West Indies, and apparently the most primitive of recent Dromiidae. Another precursor of this family was

![An ancestor of the Crabs, Palaeinachus longipes of the tribe Dromiacea, Great Oolite, Wiltshire. (After H. Woodward.)](image)

*Palaeinachus*, found in the Forest Marble, fortunately with its limbs; unfortunately, however, the unique specimen is not in the Museum, and its present whereabouts is unknown. The restoration (Fig. 49) shows the primitive nature of the carapace, abdomen, and limbs. The English Gault has yielded *Homolopsis*, which appears to lead from *Prosopon* towards the family Homolidae. *Dromiopsis*, found in the
Upper Chalk of Denmark, is an obvious representative of the Dromiidae, and Dromilites (Fig. 46, 1), of which there is shown a good series from the London Clay, is scarcely different from the modern Dromia.

Nearer to the true crabs is the tribe Oxystomata (the Sand-crabs), which owes its technical name to the fact that the mouth-frame is narrowed in front and projects forward between the eyes. In most crabs the mouth-frame is square, and the channels that carry the outward stream of water from the gills open at its two front corners; here, however, they are carried forward to the front of the head, and closed below by plates connected with the first maxillipeds. The reason for these changes is, that the Sand-crabs bury themselves in sand, leaving only the eyes exposed. This tribe appears later than the Dromiacea, its earliest representative being Mithracites vectensis, from the Lower Greensand of Atherfield. In the English Gault and Upper Greensand are found Palaeocorystes (Fig. 44, 2), Eucorystes, Necrocarcinus, Orithopsis, and Trachynotus. The precise relations of these to modern families are doubtful. The family Raninidae, found first in the Cenomanian Chalk, and not rare in Tertiary rocks, is represented in the foreign series by Eocene specimens from Kressenberg and Scinde. Other families, represented by genera still living, also appear in Tertiary times; Calappa, for instance, is represented in the British series by fragments of Eocene, Oligocene, and Pliocene age.

The tribe Oxyrhyncha, characterised by a triangular carapace, with the apex produced in front as a rostrum, is not common in the fossil state. The common genus Maia—the spider-crab—is, however, represented by specimens from the Coralline Crag of Suffolk, where it must have lived as nowadays, covering itself with masses of bryozoans and seaweed.

The crabs of the tribe Catometopa have a squarish carapace with front strongly bent downwards. Their earliest representative is the Upper Cretaceous Lithophylax, a doubtful and rare form not shown here. In Eocene rocks they are less rare, and here is to be seen the original of Gonioeypoda Edwardsi from Hampshire.

The majority of fossil crabs belong to the tribe Cyclo-
metopa, in which the carapace is, as a rule, broader than long, with the front curved and not produced into a rostrum. With the exception of the river-crabs, all modern forms of this tribe live in the sea. Among the oldest genera are
Etyius from the Gault and Cambridge Greensand, and Xanthosia from the Greensand of Warminster and Cambridge. Xanthopsis is common in the London Clay, and specimens are also shown from the Tertiary rocks of Bavaria and China. In the British Tertiary series are to be noted remains of a spiny Rhachiosoma (Fig. 50), of the swimming crab Portunus, and of the edible crab Cancer. A very large species of the latter comes from Patagonia, and there are some large examples of a Scylla from the Philippines. The river-crabs are represented by Thelphusa from the freshwater Miocene beds of Oeningen.

Class DIPLOPODA (Millipedes).

The millipedes of to-day are inhabitants of the land, with a distinct head and a worm-like body of many similar segments, each enclosed in a horny ring. Many of the segments bear two pairs of legs apiece, and thus represent two primitive segments fused together. In some millipedes the under part of the ring still consists of two plates, one to each pair of legs. In many Palaeozoic millipedes not only does each segment bear two pairs of legs, but, both on the back and on the under side, the ring is composed of two plates. On the side of each segment, near the attachment of the legs, is an opening. These openings, called stigmata, lead to the tracheae or air-tubes by which the animal breathes. The head bears two clusters of eyes and a pair of short feelers or antennae, and is furnished underneath with two pairs of
jawi-limbs, the mandibles and the maxillae, the latter joining to form a lower lip.

Millipedes of modern character are found in Tertiary rocks, especially in amber, e.g. *Julus* and *Polyxenus*. The Mesozoic rocks furnish a single doubtful form, *Julopsis*, of Cretaceous age, and a few from the Trias. The Palaeozoic forms above referred to occur in the Devonian and Carboniferous rocks. In the British series are exhibited *Eupholeria* (Fig. 51) and *Xylobius* from the Coal Measures.

**Class CHILOPODA** (Centipedes).

Superficially like the millipedes, the centipedes differ in never bearing more than one pair of legs on a body-segment, and in having four pairs of jaw-limbs. The stigmata lie in the membrane between the horny rings of the segments. These also being dwellers on land, are not common as fossils, and up to the present are not represented in the Museum.

**Class INSECTA.**

The insects are the most highly modified of Arthropoda. Like the centipedes they breathe by tracheae and have a long segmented body; but in addition to a distinct head, the hinder part of the body is sharply divided into a thorax of three segments in front and an abdomen of nine or ten segments behind, and it is only the thorax that retains legs, these being always in three pairs. The head bears a pair of compound eyes, a pair of antennae, and three pairs of jaw-limbs. Most insects have on the thorax two pairs of wings.

Owing to certain difficulties attending the study of fossil
insects, notably the slight appeal which their crushed fragments make to entomologists, it has not as yet proved possible to arrange an exhibited series in such a manner as either to indicate the riches of the Museum or to give a clear view of the palaeontological history of the Class. Here then it can only be stated that, although the earlier insects of Palaeozoic age have a primitive character, still they can be connected with some of the Orders into which modern insects are divided. Other of those Orders first appear at a rather later date. The Orders are: 1. **Aptera**, wingless insects, including the spring-tails, first found in the Carboniferous 2. **Orthoptera**, including cockroaches, possibly from Silurian, certainly from Carboniferous onwards; earwigs, beginning in Lias; grasshoppers and the like, from Lias onwards. 3. **Neuroptera**, including may-flies, dragon-flies, caddis-flies, and white ants; ancestral forms are found as far back as the Devonian if not before; more modern types come in with the Mesozoic Era. 4. **Hemiptera**, including bugs, plant-lice (Aphidae), scale insects, and harvest flies; an hemipterous wing is found in the Upper Ordovician of Sweden, and more nearly complete fossils from the Carboniferous onwards, while modern families come in Mesozoic rocks. 5. **Coleoptera** or beetles are not certainly known before the Triassic Epoch. 6. **Diptera** or flies are first found in the Lias, but are neither numerous nor readily identified before Tertiary times. 7. **Lepidoptera**, or butterflies and moths, are as yet known only from Tertiary strata. 8. **Hymenoptera**, including bees, wasps, ants, and gall-flies, are represented by ants in the Lias, but are mostly found in later Tertiary beds.

In the **British** series, the Orthoptera of the **Coal Measures** include forms allied to cockroaches, among which the specimens of *Eteoblattina* (Fig. 52) and *Leptoblattina* are noteworthy. Nodules of the same age contain wings of the Neuroptera *Lithosialis* (Fig. 53 a), *Brodiea*, showing bands of colour (Fig. 53 b), and *Lithomantis* with its expanded prothorax (Fig. 54), formerly considered an ally of the recent praying insect *Mantis*. The insects found in the **Liassic** rocks are for the most part small and insignificant, but there is a moderate-sized dragon-fly, *Libellula*, from both Lower and Upper Lias, and from the Lower Lias of Barrow-on-Soar near Leicester comes a Neuropteron allied to the white ants and called *Palaeotermes*. Elytra of beetles, sometimes with a metallic lustre, are common in the **Stonesfield Slate**,
Fig. 52.—Cockroaches of the Coal Measures.  

*a*, *Eloblattina mazona* from Illinois, enlarged 2 diameters. (After Scudder.)  

*b*, restoration of *Progonoblattina helvetica* from Switzerland. Natural size. (After Heer.)

Fig. 53.—Wings of Neuropterous insects from the English Coal Measures.  

*a*, *Lithosialis Bronquarti*;  

*b*, *Brodiea priscotincta*, with colour-bands. Natural size. (Table-case 20.)
and from this rock comes the almost complete specimen of *Blapsium Egertoni*. *Libellula* may also be noticed from the *Purbeck Beds* and from the Bagshot Beds of *Upper Eocene* age near Bournemouth. These last-mentioned beds, as well as the *Oligocene* Bembridge Beds of Gurnet Bay in the Isle of Wight, have furnished a number of insects belonging to many modern families.

Among the fossil insects from foreign localities we notice first some Orthoptera allied to the cockroaches, and some large Neuroptera, from the *Coal Measures* of Commentry, Allier, France; a locality famous for the beautiful examples that it has yielded. The next series of importance is that from the Lithographic Stone of *Solenhofen*; here

![Fig. 54.—An Insect from the Coal Measures of Scotland, *Lithomantis carbonarius*, probably a Neuropteron. \(\frac{3}{4}\) natural size. (After H. Woodward.)](image-url)

we may imagine that numerous insects lived on the islands around a lagoon, into whose placid waters they were constantly being blown. The Orthoptera are represented by the cricket *Pseudogryllacris* [*Gryllacris*], by the locust *Pycnophlebia*, and by *Chresmoda* [*Pygolampis*] *obscura*, a precursor of the Mantidae and Phasmatidae or stick-insects. Among the Neuroptera is the dragon-fly *Cymatophlebia*. Water-bugs allied to *Nepa* and *Belostoma* represent the Hemiptera. Beetles are numerous. The chief example of the Hymenoptera is *Pseudosirex*, one of the tailed wasps that bore into trees. Cretaceous insects are unrepresented, but there are a number from *Tertiary* rocks, all of modern type. The Oligocene deposits of Aix in Provence and of Florissant in Colorado, the Miocene Beds of Oeningen on
Lake Constance, and the Brown Coal of Rott near Bonn, also Miocene, all furnish their quota. The Indusial Limestone of Lower Miocene Age from Offenbach is composed of the cast-off cases of caddis-worms, Phryganea. Many insects also come from the Miocene deposits of Radoboj in Croatia, and among these a cricket (Gryllacris Ugeri) is preserved in a most life-like attitude. Other Miocene insects are seen in amber cast up on the shores of the Baltic, just as they are preserved in the hardened gums of later age from Zanzibar and elsewhere.

Further information may be obtained from "The Fossil Insects of North America," by S. H. Scudder (New York, 1890), and from the handbook by A. Handlirsch, "Die Fossilen Insecten" (Leipzig, 1906-07).

**BRACHIOPODA.**

Following on the Arthropoda, are exhibited the commonest of all fossils, the Brachiopods or Lamp-shells. The important specimens contained in the Davidson, Sowerby, and Gilbertson Collections have already been noticed. (See pp. 9, 10.)

The Brachiopoda are animals that live in the sea, and have a soft body enclosed in an external shell with two valves (Fig. 55). They thus look something like bivalve Mollusca; but both the shell and the soft parts have really a very different structure from those of the Mollusca. So much of the anatomy of the Brachiopoda as is important to the student of fossils, is illustrated by the large coloured diagrams in the wall-cases.

The two valves of the shell lie on the back and front of the animal, not on its sides as in bivalve molluscs. Each

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**Fig. 55.—Shell of a common Silurian Brachiopod, Atrypa reticularis." a, ventral or peduncular valve; b, dorsal or brachial valve. Shows bilateral symmetry, and slightly greater size of ventral valve.**
valve is symmetrical in itself, its right and left halves resembling one another. On the other hand, one valve is nearly always larger than the other (Figs. 55 a and 60 a). By the larger valve the adult animal is usually attached to rocks or other objects. In a few forms, such as Crania (Diagram 1), the valve is directly cemented to the rock by its own substance. In others, such as Lingula (Diagram 2 and Fig. 57), the valve is attached by a long muscular stalk, the “peduncle” or “pedicle,” which is capable of waving movement and of contraction. There are also intermediate stages, with shorter peduncles, such as Hemithyris (Diagram 8) and Orbiculoidea (Fig. 58 b). This larger, attached valve is therefore often called the “peduncle valve”: by English writers it is called the “ventral valve,” although in the natural position of stalked forms it is always the uppermost. Similarly the smaller valve is called the “dorsal valve”; but a better name is “brachial valve,” which reminds one that this valve often bears a calcareous skeleton supporting the so-called “arms” (Diagram 9 and Figs. 59, 60).

In microscopic structure also the shell differs from that of the Mollusca (Diagram 3). It is mainly composed of small prisms of calcite (carbonate of lime), which usually lie at right angles to the surface of the shell. In the harder-shelled forms these make up the greater part of the shell, merely being coated on the surface with a layer of slightly different texture and with a thin horny membrane. In the softer-shelled forms, such as Lingula, horny substance occurs not merely on the surface, but in layers between the prisms, which in this case are of phosphate of lime. In many genera, such as Terebratella, Crania, Cistella, the shell is perforated by a number of small canals; these contain processes of the mantle, the arrangement of which is shown in Diagram 3 a. In fossils this structure gives to the exterior of the shell a pitted or “punctate” appearance under a magnifying glass, and thus enables one to distinguish such shells from those which are “impunctate,” as the shells of Atrypidae and most Rhynchoselonellidae.

The shell-valves are secreted by the two mantle-folds which line them. These are extensions of the body-walls, and they contain prolongations of the body-cavity, in which is a blood-like fluid and in which the generative products are formed. These vessels often produce impressions on the inside of the shell, and so can be traced in the fossils, e.g. Camarophoria (Fig. 56) and a fine specimen of Orthis Wall-case 11c.
Gallery VIII. (Schizophoria) striatula. The outer edges of the mantle-folds are set with bristles (setæ). All these structures are shown in Diagrams 5 and 11.

Muscles pass across the body of the animal from one shell-valve to the other (Diagram 6); they serve to open and to close the valves, and to move them sideways. The attachment of these muscles to the shell forms scars, which in fossil brachiopods are the only evidence we have as to the arrangement of the muscles. The drawing from life of the inner surface of a Lingula shell (Diagram 7 and Fig. 57) should be compared with Diagram 8 showing the muscles.

The viscera lie near the hinder or peduncular end of the shell, with the mouth directed towards its forepart or opening. The mouth is surrounded by a somewhat horseshoe-shaped disc; this bears tentacles, furnished with minute, rapidly-moving processes (cilia), which drive towards the mouth currents of water containing food-particles; it is
called the "lophophore" (tuft-bearer), and resembles the structure of the same name in the Bryozoa (Diagram 8). The lophophore is generally produced into two so-called "arms," which fill the forepart and sides of the shell-cavity and are often spirally coiled. Since they were formerly supposed to represent the "foot" of the Mollusca, their presence suggested the name Brachiopoda (arm-feet). The mouth leads to a slightly coiled intestine (Diagram 10), which in the simpler genera is provided with an anus, whence the Class comprising them received the name Tretenterata (pierced guts). Fossils indicate that some of the earlier genera of the other Class were also provided with an anus; but in its later genera this structure became degenerate, and no longer exists in the living representatives; for this Class, therefore, the name Clisterenterata (closed guts) was proposed. These names are now generally supplanted by Inarticulata and Articulata (see p. 112).

The Brachiopoda are found in seas all over the world, and usually at depths of less than 100 fathoms, but they have been dredged at a depth of 2,900 fathoms. Most kinds attach themselves permanently to a hard bottom by the peduncle, open their shell so far as the hinge permits, and collect minute food-particles in the currents of water that flow down the lophophore; some protrude and even unroll the arms. Lingula, as shown in Diagram 2, lives in a tube in the sand, forming a case of agglutinated sand round the lower end of its peduncle; it stretches its shell to the opening of the tube, and the projecting setæ guide the currents of water down to the lophophore; but when disturbed, the peduncle contracts and the shell is withdrawn into the tube, which closes in above. It is not, however, to be inferred that all extinct species of Lingula and of similar genera lived in this way.

Though brachiopods usually occur in great numbers wherever found, they are not so numerous now as they were in past ages. In the Carboniferous Epoch especially, the number of species and individuals was very great, and the Producti then living reached a larger size than any brachiopod before or since. Terebratula grandis, of the Coralline Crag, is the largest brachiopod found in later rocks. Many examples of masses of brachiopod shells are exhibited, and among them may be mentioned a slab covered with Lingulella Davisi, from the Lower Lingula Flags of Upper Cambrian Age, near Tremadoc; Ungulite Grit with Obolus and
Gallery VIII. 
Table-case 18.

Schmidtia from the Upper Cambrian of Russia; a mass of Lower Llandovery sandstone with Pentamerus oblongus; groups of Conchidium Aylesfordi from the Aymestry Limestone, of its ally Sieberella galeata, and of Atrypa reticularis from the Wenlock Limestone of Dudley; blocks of Jurassic Age with Aulacothyris and various species of Rhynchonella; and slabs of Middle Lias, Great Oolite, and Cornbrash, with their characteristic brachiopods.

The classification adopted in this Department is that used in the English edition of Von Zittel's 'Palaeontology.' This is based on the supposed evolution of the group, and may be followed in the Wall-cases, where the fossils from foreign localities are arranged in zoological order, beginning with the top shelf and following its whole length from left to right, and continuing on the middle and bottom shelves in the same manner. By attending to this, the genera mentioned in the following account will readily be found.

Fig. 58.—Types of Brachiopod Orders. a, Order Atrema, the Cambrian Iphidea labradorica, one of the simplest forms of brachiopod shell. b, Order Neotremata, the Ordovician Orbiculoidea lamellosa, the peduncular valve, showing the delthyrium, which has become surrounded by the valve and partly filled in from underneath by a shelly deposit. c, Order Protremata, the Ordovician Clitambonites Verneuili, seen from the side of the brachial valve, above which is the lofty hinge-area of the peduncle valve. The delthyrium is covered by a single pseudodeltidium, through which the peduncle passed by the foramen.

In those Brachiopoda that appear to be the simplest and oldest, the shells are not as a rule joined by any hinge (Diagram 7). These have therefore been called Inarticulata or ECARDINES (e, without; cardo, hinge): they include Lingula, Discina, Obolus, Crania, Trimerella, and their allies. In more advanced forms, such as Orthis, Leptaena, Atrypa, Terebratula, a hinge is developed at the hinder end of the shell (Diagram 4), and these have been called Articulata or TESTICARDINES (testa, shell; cardo, hinge). As classificatory divisions, however, these are not altogether satisfactory, for the Articulata are necessarily derived from the Inarticulata, and intermediate forms are not rare.
The relations of the peduncle to the valves manifest a gradual evolution. The simplest type is seen in Iphidea (from the Cambrian, Fig. 58, a) and Lingula (Cambrian to present day, Fig. 57), where the peduncle simply passes out between the valves and is not enclosed by either of them; such genera are therefore without a hole (trema) through which the peduncle may pass, and constitute the Order Atremata. In Trematis, Orbiculoidea (Fig. 58, b), Siphonotreta, and their relatives (mostly Ordovician and Silurian), the peduncle is restricted to the ventral valve; it lies in a groove or fissure (delthyrium), which remains open in primitive forms, but closes round the peduncle (forming a trema) in later forms: such genera constitute the Order Neotremata. Next, the fissure in which the peduncle lies assumes a triangular shape; the peduncle is towards the apex of the triangle, and itself secretes a single shelly plate (pseudodeltidium), which gradually fills up the triangular fissure till only a small foramen is left at its apex, as in Olitambonites (Fig. 58, c) and Rafinesquina (Diagram 4); later in life the pseudodeltidium may be reabsorbed, as in the Orthidae: such genera constitute the Order Protremata. Some forms have taken another line of evolution: in them the pseudodeltidium is either absent or soon reabsorbed, so that the delthyrium is open in early life, but at a later period it becomes partly or entirely closed by two "deltidial plates," which are secreted by the edges of the mantle along the sides of the delthyrium, and which may subsequently meet either above or below the peduncle, and may even fuse into one plate, the "deltidium"; these plates are well seen in Atrypa and Stringocephalus, and occur in Rhynchochelus, Spirifer (Fig. 59), Terebratula, and allied forms: such genera constitute the Order Telotremata. The arms are often supported by a calcareous skeleton, the shape of which is of great importance in classifying fossil brachiopods. Thus, the Telotremata branched into loop-bearing forms (Diagram 9 and Fig. 60, Magellania) and spire-bearing forms (Diagram 9, Atrypa, and Fig. 59, Spirifer).

The Orders Atremata and Neotremata are equivalent to successive stages of the Inarticulata, and are most abundant in the earlier Palæozoic rocks. From them the Protremata and Telotremata arise as divergent groups, which together are the equivalent of the Articulata. The Protremata were dominant in later Palæozoic time; the Telotremata in Mesozoic.
Gallery VIII.
Wall-cases 10 & 11.

In addition to the particular specimens already pointed out, attention may be directed to the following. Specimens brought from the Arctic Regions by various British expedi-

Fig. 59.—An example of the spire-bearing Telotremata, *Spirifer striatus*. Carboniferous Limestone. The shell is seen from the side of the brachial valve, and portions of that valve are broken away, exposing the spires that support the arms of the lophophore. Between the umbones of the peduncular and brachial valves is seen the delthyrium, partly filled in by the deltidial plates that have met and fused above the foramen into a single deltidium. (From the “Cambridge Natural History.”)

Fig. 60.—An example of the loop-bearing Telotremata, *Magellania flaves-
cens*. Recent seas, Australia. A.—Interior of peduncle valve. f, foramen for peduncle, below which are seen the two deltidial plates; t, teeth of hinge; a, b, c, muscle-scars; b', scar of peduncular attachment. B.—Interior of brachial valve. c, e', cardinal process for attachment of muscles; b, hinge-plate, supporting cardinal process and prolonged below into p, the median septum; s, sockets for the teeth of the peduncle valve; l, loop, supporting lophophore; a, muscle-scars.

tions; a good series, chiefly of *Spirifer*, from the Permo-
Carboniferous rocks of Australia; Silurian and Carboniferous specimens from near Niti in the Northern Himalaya; and a Devonian series from the Hindu Khoosh. Many of these
have been specially described in technical publications. The British series includes the originals or type-specimens of several species from the Lower Greensand of Upware described by Mr. J. F. Walker. Otherwise this portion of the collection appears poor in type-specimens, owing to the fact that so many are exhibited among the historical collections. Some attempt is therefore being made to show here specimens interesting from their habit or structure. Such are the examples of *Spirifer*, in which the arm-spires were worked out with a needle and acid by the Rev. Norman Glass, and specimens of *Terebratula* and *Ornithella* from the Cornbrash, showing the arm-loop. The interior of the shell is also well shown in a *Liothyrina* from the Norwich Chalk on the other side of the Case. Another Jurassic brachiopod *Acanthothyris spinosa* owes its name to the numerous spines borne by its shell; these were very long and are generally broken off, but there are exhibited both British and foreign examples in which they are most wonderfully preserved and displayed. *Productus* is another genus richly provided with spines, and these owing to their size are often found scattered in the rock as thin tubes; a slab of Carboniferous Limestone covered with them is shown. A pretty impression in flint of the rare *Trigonosemus elegans* is among the Cretaceous fossils. Various Jurassic *Terebratulae* and terebratuloid shells reflect changes of growth, due to old age or illness, in variations of their ornament. The same outer form has frequently been assumed by brachiopods of different internal structure at widely separated geological periods, a circumstance very perplexing to the field-geologist; but here is exhibited a set of four different species, belonging to at least three genera, all coming from the Inferior Oolite Marl, and so much alike that a casual observer could hardly tell them apart. A similar independent recurrence of form is displayed by *Pygope* and other "diphyoid" genera, in which the fore-part of the shell has grown out into two lobes that ultimately meet and enclose a vacant space. These various modes of growth and of evolution are by no means confined to Brachiopoda.

Further information on the Brachiopoda may be sought in the memoirs by T. Davidson already referred to, and in "An Introduction to the Study of the Brachiopoda," by J. Hall and J. M. Clarke (Albany, 1904–5).
BRYOZOA or POLYZOA.

The fossil remains of the Bryozoa (moss-animals), or Polyzoa, as they are often called, a group of which the modern Sea-mats are familiar examples, are exhibited in three cases in the middle of the gallery, adjoining the Brachiopoda, with which they are by some supposed to be allied. These specimens are all British, and form an adequate representation of these widely spread fossils. In Wall-case 12 are placed a few foreign specimens of particular interest, as well as certain colonies too large to find room in the British series proper.

![Diagram of structure of a typical Bryozoa zooid](image)

Fig. 61.—Diagram of structure of a typical Bryozoa zooid, as seen in a vertical cut down the middle. *an*, anus; *ap*, aperture; *b.c*, body-cavity; *c.p*, communication pore; *d*, diaphragm; *ect*, ectoderm; *end*, endoderm; *f*, funiculi; *n*, nerve-ganglion; *o*, orifice; *œs*, œosphagus; *op*, operculum; *r.m.*, retractor muscle; *st*, stomach; *T*, tentacles; *t.s.*, tentacle sheath.

Though often mistaken by wanderers on the sea-shore for sea-weeds, the Bryozoa are really animals. They live in either fresh or salt water, mostly in the latter, and with one exception (*Loxosoma*), always form colonies, which are generally fixed. A colony consists of a large number of individuals (or zooids), each of which is completely separated from the rest and enclosed in a double-walled sac. The digestive tube is U-shaped, the mouth and anus being placed close together. A band of tentacles occurs around the mouth in most forms, but in one group (Entoprocta) this surrounds both the mouth and the anus (Fig. 61). These colonies may spread in delicate gauze-like sheets over weeds, shells,
and stones, rise in hard shrub-like tufts, forming hemispherical masses, or stretch out flexible horny branches. In the fossils the soft body of the animal is of course destroyed, and there remain only the hardened walls of the little chamber in which each zooid lived. Fortunately the shapes of these chambers afford characters by which the species can be classified. They must, however, be studied under the microscope. In the arrangement of the collection, therefore, specimens are exhibited to show the general form and habit of the colony, and drawings are placed beside them to show the minute structure of the chambers. The specimens are in most cases near enough to the front of the case to admit of the use of a magnifying glass, and thus the main features of the chambers can be recognised.

Setting aside the aberrant *Loxosoma* and its allies, modern Bryozoa are divided into two Sub-Classes: (1) **PHYLACTOLAEMATA**, in which the mouth of the zooid has a lip, and the crown of tentacles or lophophore is horse-shoe-shaped; (2) **GYMNOLAEMATA**, in which there is no lip to the mouth, and the tentacles form a complete circle. Since the chambers of the Phylactolaemata are either soft or horny, they are not preserved as fossils, so that we are concerned only with the Gymnolaemata. Omitting the doubtful Ctenostomata, of which no fossils are certainly known, these last are divided, according to the structure of the chamber-opening, into four Orders: (1) **Trepostomata** (turned mouths); (2) **Cryptostomata** (hidden mouths); (3) **Cyclostomata** (round mouths); (4) **Cheilostomata** (lip mouths).

The Trepostomata and Cryptostomata are all extinct, but some of each Order seem to have given rise to the Cyclostomata and Cheilostomata respectively. The Trepostomata, which were dominant in early Palaeozoic times, generally form massive colonies, composed of the chambers drawn out into long tubes and set side by side; the tubes turn upwards towards the openings at their ends; as the colony grew each animal moved up in its tube, the lower part of which was cut off by a platform (diaphragm) like the tabula in Tabulate Corals. Examples are *Monticulipora, Stenopora*, and *Callopora* (Fig. 62). In the Cryptostomata, which became dominant in later Palaeozoic times, the opening of the chamber is hidden at the bottom of a tubular shaft (vestibule); the chambers grow up into continuous tubes, as in the Trepostomes. Examples are *Ptilodictya, Rhabdomeson,*
and *Fenestella* (Fig. 63). The *Cyclostomata* were dominant during the Mesozoic Epoch; their chambers have simple round openings at their ends, with no covering. Examples:

![Diagram of Trepostomatous Bryozoan](image1)

**Fig. 62.**—A Trepostomatous Bryozoan, *Callopora submodosa*, Ordovician (Cincinnati group), N. America. *a*, fragment of a colony, natural size. *b*, part of the surface, magnified 12 diameters. *c*, part of a vertical section, showing tubes of differing size crossed by diaphragms, magnified 18 diameters. (After E. O. Ulrich.)

![Diagram of Cryptostomatous Bryozoan](image2)

**Fig. 63.**—A Cryptostomatous Bryozoan, *Fenestella vera*, Devonian (Hamilton group), N. America. *a*, fragment of a colony, obverse, showing the fenestrae, to which the genus owes its name, as white spaces, and the chamber-openings as small dark holes. *b*, reverse of same, showing fenestrae only. *c*, section of same, passing from near obverse surface on left to near reverse surface on right. *a* and *b* magnified 9 diameters; *c*, 18 diameters. (After E. O. Ulrich.)

*Stomatopora, Berenicea* (Fig. 64). In the *Cheilostomata*, which reached their present dominant position in Cainozoic time, the opening is removed from the end of the chamber, is constricted, and provided with a movable lid (operculum).
Examples: Flustra (Fig. 61), Membranipora (Fig. 65a), Onychocella (Fig. 65b).

In many Bryozoa certain individuals are modified for special duties. Thus in the Cheilostomes the growth of the operculum has produced snapping beaks, called avicularia, and long movable bristles, called vibracula; the positions occupied by these appendages can be detected in the fossils (Fig. 65). Sometimes individuals are set apart and modified for reproduction, sometimes special pouches for the reception of the developing eggs are attached to the chambers (Fig. 65a).

These modifications suggest explanations for the smaller chambers and tubes interspersed among the normal ones in the fossils of extinct Orders.

We may now briefly review the exhibited specimens. The British series begins with some from the Ordovician rocks of Wales. These are so poorly preserved that they can only be determined by the help of better specimens from American rocks of the same age. Most of them are Trepostomes or Cryptostomes.

The Silurian Bryozoa are in a better state, and the specimens from the Wenlock Limestone include some
interesting and elegant Cryptostomes, such as *Ptilodictya sublanceolata*, *Penniretepora Lonsdalei*, and several species of *Fenestella* and *Polypora*.

The Devonian fauna is small, and its representatives not well preserved; a specimen of *Fenestella prisca*, figured by Phillips, and a *Polypora populata* are worth notice.

The Carboniferous Bryozoa, on the other hand, are numerous and show their structure well. The commonest genus is *Fenestella*, and the size to which its colonies may attain is exemplified by two specimens mounted on blocks.

The fan-shaped species of this genus and of *Ptilopora*, the feather-shaped *Penniretepora*, and the cylindrical *Rhombopeora*, are the most interesting forms in the British Case. *Hemitrupa* should not be overlooked, as it consists of a colony formed of two layers, of which the upper was once regarded as a coral growing as a parasite on a *Fenestella*. All the genera just mentioned are Cryptostomes. The adjacent case contains a selection from the Carboniferous Bryozoa collected by Mr. G. W. Shrubsole, mostly from Halkin Mountain in North Wales.

The Permian species are few; but three larger masses

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**Fig. 65.**—Cheilostomatous Bryozoa, from the Bathonian (Calcaire à poly-pliers) of Normandy. *a*, *Membranipora jurassica*, portion of a colony; in front of the large chamber-openings are in some cases small depressions left by avicularia, in others round openings leading into egg-pouches. Enlarged 13 diameters. *b*, *Onychocella flabelliformis*, portion of a colony; on the right is a chamber much narrower than the others, indicating that the individual was turned into an avicularium. Enlarged 24 diameters. (From J. W. Gregory.)
of *Fenestella* and *Synocladia*, another Cryptostome, are mounted on blocks.

The next fauna represented in England is that of the Jurassic system. This shows a great advance on that of the Palaeozoic; old types such as *Fenestella*, *Penniviretepora*, *Acanthocladia* disappear, and species belonging to existing genera form the larger part of the fauna. Among these the Cyclostomes, *Stomatopora*, *Berenicea*, and *Diastopora* are important, and are illustrated by an extensive series of specimens, but *Theonoa* and *Apsendesia* are typically Jurassic. These are associated with forms such as *Ceriopora*, which are survivals of the Trepostomata.

The ensuing Cretaceous fauna in many ways resembles the Jurassic. Trepostomata survive, and the Cyclostomata are still in the ascendant. Examination, however, of the specimens exhibited shows that the Cheilostomata are now fairly abundant, as we may see by the numerous species of *Membranipora* and *Onychocella* (Fig. 65) and the presence of more specialised genera such as *Cribrilina*.

Passing to the Eocene we find that the fauna in England became smaller, though that of the Mediterranean Basin at the same period was very large. The forms, moreover, were scarce and dwarfed, as they lived in a sea exposed to the north and cut off from the warm waters of the Mediterranean by a land barrier across Central France and Germany. Hence in our series genera such as *Idmonea* are represented only by small delicate colonies (see also *Idmonea coronopus*, from the Paris Basin), which are in striking contrast to the massive growths exhibited from Italy. In addition to these general differences, the fauna includes a high percentage of peculiar species, among which *Orbitulipora petiolus*, consisting of a disc supported on a short stem (see specimen B. 4349), *Aedonellopsis Wetherelli*, and *Notamia Wetherelli*, are the most remarkable. The species of the still existing genera *Schizoporella*, *Smittia*, and *Entalophora*, on the other hand, are quite modern in form. Except for the Cyclostomes *Idmonea* and *Entalophora*, all the genera mentioned in this paragraph are Cheilostomes.

The Pliocene in England is much richer than the Eocene, and comparison need only be made between the small fragile specimens from the latter and such massive colonies as those from the Crags, to realise that the Bryozoa were then living under more favourable conditions; the Arctic Ocean was probably cut off by a land barrier to the
north, while there was free communication with the seas to the south. The Museum possesses the specimens used by Busk in writing his monograph on the "Crag Polyzoa," and many of them are exhibited. The most interesting forms found in the Crag are some massive Cyclostomes, including the three species known as *Alveolaria semiovata*, *Faseicularia aurantium*, and *F. tubipora*. Among the Cheilostomes, the most remarkable forms are two species of *Cellaria* [*Salicornaria*] and one of *Melicerita*. The numerous species of *Schizoporella*, *Mucronella*, and *Membranipora* are closely allied to or identical with living forms.

A small collection of Pleistocene species from the Clyde and from Selsea Bill is shown; but all these species still live on the English coast.

The Bryozoa from foreign localities are not yet completely arranged, and only a few representative species are exhibited. The lowest slope of the wall-case is devoted to the Palaeozoic faunas, chief of which are those from the Ordovician and Silurian rocks of North America. The Carboniferous fauna of the same continent furnishes some remarkable forms, notably *Archimedes* *Wortheni*, which is like a *Fenestella* twisted into a screw, and *Evactinopora quinqueradiata*, another Cryptostome with a star-shaped colony.

A collection from the Bathonian deposits of Northern France on the middle slope contains several interesting forms, notably *Membranipora jurassica* and *Onychocella flabelliformis*, which are the oldest true Cheilostomes known (Fig. 65).

Among the Tertiary Bryozoa on the top slope, the large specimens from the Miocene deposits of the Mediterranean are most worthy of notice.

The Trustees have published a Catalogue of the Jurassic Bryozoa (1896), and the first volume of a Catalogue of the Cretaceous Bryozoa (1899), both by Dr. J. W. Gregory.

MOLLUSCA.

These animals derive their name from their soft bodies (*mollis*, soft), which never have any internal skeleton, and rarely any hard appendages capable of preservation as fossils. The glandular skin, however, usually secretes, on a portion of the outer surface called the mantle, a hard shell, sometimes horny in appearance, but usually thickened by a deposit of
carbonate of lime. The fossil remains of the Mollusca consist therefore chiefly of shells, and these in substance and appearance may not differ from shells that one picks up in the fields or on the sea-shore. Such are many of the Post-Pliocene non-marine shells, and *Neritina concava* from the Oligocene of Headon. Usually, however, the horny layers of the shell have entirely disappeared, and are replaced by a secondary deposit of lime that has soaked in from the surrounding rock and hardened or petrified the shell. Sometimes the original lime itself has been replaced by another mineral, such as flint or iron pyrites, and sometimes the whole shell has disappeared, leaving only a cavity in the rock, with an imprint (or external cast) of the shell on its outer walls, and sometimes containing a mass of rock that filled the interior of the shell, and is called an internal cast. This last method of preservation is clearly shown in a large block from the Roach bed in the quarries of Portland, Dorset. Shells preserved in flint are common among fossils from the Chalk, and pyritised shells are generally found in clays of all ages. A few Mollusca possess other hard parts capable of fossilisation; thus several of the cuttle-fish and their allies have horny beaks, which form the fossils called *Rhyncholithes* (beak-stones, Fig. 78), some have horny hooks on their arms, and some, buried in clay, have even left traces of skin, muscles, and eyes (Fig. 86). We can also infer the former presence of certain boring molluscs, such as *Pholas*, by the burrows they have made in the rocks of the old sea-floor.

Mollusca are in some of their forms, such as oysters and snails, so familiar that we need here only recall the fact that they are among the more highly organised invertebrates, having in the simpler types a distinct head, a mouth and complete digestive system, and a thickening on the under side of the body forming a muscular organ called the foot. All molluscs have a nervous system, blood-vessels, a heart, gills of varying origin, and excretory and reproductive organs. Further information on these matters may be sought in the Zoological Department.

The shells of molluscs are built on various plans; the oyster, for instance, has a shell of two valves; the whelk or the snail-shell is all in one piece, and is besides coiled and open only at one end; the nautilus shell is also coiled, but when cut through it is seen to be divided up by partitions into a number of chambers; the tooth-shell, *Dentalium*, is a tube shaped like a tusk, but open at both ends; *Chiton*,

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**Gallery VIII.**

**Upright Case A1.**

**Table-case 3.**

**Centre of Gallery VIII.**

**Gallery VII.**

**Table-case 1.**

**Gallery VIII.**

**Table-case 9.**
which may be found creeping on the rocks at low tide, has a jointed shell, and looks like a wood-louse. Examination of the animals that live in these five types of shell shows that they are built on as many plans of structure, and to one or other of these plans all molluscs, except the Palaeozoic Conularida (p. 141), can easily be referred. Therefore the Mollusca are divided into the following five Classes:—

I. **AMPHINEURA**, of which *Chiton* is an example, owe their name to the two nerve-cords that run down each side of the body, which is elongated and symmetrical. The mantle always secretes little plates or spicules of shell-substance. They are divided into two Orders: (a) *Aplacophora*, which have no shell other than the spicules, and therefore are not found fossil; (b) *Polyplacophora*, which have a shell of eight larger pieces, surrounded by a flexible girdle formed of the mantle-edge, in which are usually smaller plates or spicules (Fig. 66). Appearing first in Ordovician rocks, they have persisted till the present day, with increasing elaboration of the shell, but with no changes of sufficient importance to mention here.

II. **GASTROPODA**, of which the snail *Helix* is an example, derive their name, meaning Belly-foot, from the position of the large foot beneath the stomach and viscera, which are contained in a hump on the animal's back; the surface and folded edges of this hump constitute the mantle that forms the shell. Thus the shell is a cone, sometimes short, as in the limpet (*Patella*), but generally long, and coiled either in one plane as in the ram's-horn snail (*Planorbis*), or spirally as in *Helix* (Fig. 68, 7). In the common snail it may easily be seen that the edges of the mantle form a cavity (the pallial chamber) on the right side of the animal; and into this open the anus and genital duct, which have been brought towards the mouth end of the body by the curving upwards of the viscera into the hump. In many gastropods this twisting of the end of the gut forwards and to the right side has affected other organs and notably the nerve-cords. This affords a basis for dividing the Gastropoda into two Sub-Classes:—

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![Figure 66](image-url)
(a) **STREPTONEURA** (twisted nerves), in which the loop formed by the visceral nerves is twisted. Since the gill-cavity is in front of the heart, they are also known as the Order **Prosobranchia**. Example: the whelk, *Buccinum*. To this Order are now referred certain small gastropods (e.g. *Carinaria*), with very thin shells and a muscular expansion forming a vertical fin by which they swim near the surface of the sea. They are sometimes distinguished as **Heteropoda**; some are found in Tertiary rocks.

(b) **EUTHYNEURA** (straight nerves), in which the nerve-loop has untwisted. These are again divided into the Orders: **Opisthobranchia** (hind-gills), marine forms, e.g. *Bulla*; and **Pulmonata** (with a lung, i.e. air-breathers), terrestrial or fresh-water forms, e.g. *Helix*. With the Opisthobranchs are now placed some small thin-shelled gastropods in which the foot is changed into two wing-like fins, one on each side (Fig. 67). They were formerly considered as a separate Class, **Pteropoda** (wing-foot), to which the Conularida were also referred.

Gastropods already existed in early Cambrian times, but all through the Paleozoic Era they remained of simple type, being mostly marine Streptonera. The number of genera and species increased greatly during Mesozoic time, but it was not till the later Jurassic and earlier Cretaceous rocks that non-marine forms were preserved in any abundance. In the older Tertiaries most of the genera are the same as those now living, though the species are different. During the Miocene Epoch a few modern species made their appearance, and of the Pliocene species about 85 per cent. have persisted to the present day. On this gradually increasing proportion of recent species of Mollusca Lyell based his division of Tertiary time into Eocene, Miocene, and Pliocene.

III. **SCAPHOPODA** (digger-foot), of which *Dentalium* is an example, have a foot adapted for burrowing in mud or sand. The shell is tubular, since the folds of the mantle have grown together on the under side of the animal; its broader front end, from which the foot can stretch out, is sunk in the sand, while the narrower hind end projects above it and discharges the waste products. Scaphopods
have endured with little change from the Silurian Epoch to the present day.

IV. **LAMELLIBRANCHIA**, of which the oyster and cockle are examples, usually have complicated gills (branchiae) formed of many lamellae or plates; the foot is rarely used for crawling, but is generally wedge-shaped, whence they are also called **PELECYPODA** (hatchet-foot). These features are naturally associated with a sedentary habit of life and with the suppression of the head-region. The shell is not deposited by a visceral hump, but by two flaps of the mantle, placed on the right and left sides; hence it consists of two valves, which are joined along one edge by a ligament and generally a hinge, and can be closed by powerful muscles (adductors). The lamellibranchs are confined to the water, and most are marine. Some, like the oyster, are fixed; most burrow in mud or sand, and a few bore into wood or rock. This Class, as a whole, presents a somewhat uniform structure, and it can hardly be said that any of the numerous attempts to divide it into Orders has met with general acceptance. Therefore we shall only indicate some of the chief variations that can be seen in the shell. These are: (1) the adductor muscle-scars, whether two equal, two unequal, or only one; (2) the outline of the mantle-attachment (Fig. 68, 2), whether simple or indented by a sinus due to certain muscles that work tubular extensions of the mantle called siphons (but the absence of a sinus does not imply the absence of siphons); (3) perfect or deficient symmetry of the shell-valves; (4) shell-structure, whether porcellaneous or nacreous; (5) the arrangement of the ligament; (6) the hinge (Fig. 69, g), whether plain or toothed, and the varying numbers and development of the hinge-teeth.

A few ill-preserved shells, apparently of simple Lamellibranchs have been found in the Cambrian rocks of Wales, Thuringia, and North America. In Ordovician rocks they are still rare, but in Silurian times a score of families existed, mostly with thin shells of simple type. The Devonian saw the beginning of brackish and fresh-water lamellibranchs; these increased in Carboniferous times, when also appeared **Allorisma**, the first form known to have a retractile siphon. With the Trias many of the older genera disappeared and new families came in, followed by others in the Jurassic period, when also **Trigonia** arose and soon flourished in numbers (Plate V.). Among Cretaceous lamelli-
branches the Rudistae are remarkable and characteristic. The early Tertiary fossils are evidence of a gradual change to present conditions, both in the character of the molluscan fauna and its eventual distribution in the same provinces as are at present recognised.

V. **CEPHALOPODA** (Head-foot), of which the Nautilus and cuttle-fish are examples. These are exhibited in Gallery VII. We shall therefore leave them for the present, and confine our attention to Gallery VIII., which contains fossil shells of only the first four Classes, viz.:—

**AMPHINEURA, GASTROPODA, SCAPHOPODA, LAMELLIBRANCHIA.**

It will be most convenient to start with the British specimens. Beginning at the main entrance to the Gallery, and following the numerical order of the Table-cases, these are arranged under the larger stratigraphical divisions, with the newest first.

**Post-Pliocene.** The marine forms are separated from the non-marine, and are arranged under geographical districts in the following order: S. England, Selsea, &c.; S.E. England, including Norfolk; N.E. England, chiefly Yorkshire; W. England, Gloppa, near Oswestry; Wales, Moel Tryfan; Ireland, the Wexford gravels and Belfast; Scotland, Clyde series; E. Scotland, chiefly Dyer's Burn and Golspie; W. Scotland, including the Lewis. Some of these date from the Glacial period, others are later, as indicated on the label in each instance. They come from raised beaches, glacial drift, and other deposits, and are often found far above present sea-level, as much as 1350 feet at Moel Tryfan, 1120 feet at Gloppa, and 1200 feet at Calbeck Castle, near Dublin. Some writers have supposed that they were carried to these heights by moving ice. Most of the shells belong to species still existing, and they are as a rule characteristic of northern seas, but southern forms are sometimes associated with them. In the following list of the more important species, the letters N, S, and E respectively denote the Northern, Southern, and Extinct forms:—Gas tropoda: *Bittium reticulatum* (S), *Boreotrophon [Trophon] clathratum* (N), *Littorina littorea* (N), *Neptunea antiqua* (N), *Tricolor [Phasianella] pulla* (S), *Tritonofusus Leckenbyi* (E), *Turritella communis* (N). Lamellibranchia: *Acila [Nucula] Cobbolidae* (E), *Area lactea* (S), *Callista [Venus] chione* (S), *Cerastoderma [Cardium]*
edule (N), Chlamys [Pecten] opercularis (S), Cyprina islandica (N), Macoma [Tellina] balthica (N), Mya truncata (N), Serobicularia plana (S), Tridonta [Astarte] arctica (N), Ven-tricola [Venus] cosina (S).

The **non-marine** series comprises shells of land and fresh-water mollusca, found chiefly in river drifts and alluvium. These also are arranged under localities, of which the following are the more important: London and Essex (Fulham, Blackfriars, Admiralty Buildings, Lea Valley,

![Shells of Non-marine Mollusca from the River-Drift of the London District.](image)

Copford, Walthamstow, Witham, Ilford, and Grays); Kent (Crayford and Crossness); Cambridgeshire (Barnwell and Barrington). Most of the species still live in England, but a few are altogether extinct. Thus of those depicted in Fig. 68 only *Pisidium amnicum* remains; other extant species are *Helix nemoralis*, *Jaminia [Pupa] muscorum*, *Pomatias [Cyclostoma] elegans*, *Bithynia tentaculata*, *Ancylus fluviatilis*, and *Neritina fluviatilis*, among gastropods, with *Unio tumidus* and *Anodonta cygnaea* among lamellibranchs.
Pliocene. The shells are primarily arranged under the formations from which they come: the Norwich Crag and the Red Crag, both of Astian age; the Coralline Crag, of Plaisiancian age; the Lenham beds of Kent and the St. Erth beds of Cornwall, both probably Plaisiancian. No rocks of Sicilian age occur in Britain. The Crags are well known for their abundant molluscan fauna, of which further examples, mostly collected by the late Robert Bell, are exhibited in Cases A 2 (Lamellibranchia) and A 3 (Amphineura, Gastropoda, Scaphopoda). The fauna is essentially marine, but a few non-marine shells invite speculation as to the cause of their occurrence. From the Coralline Crag of Gomer, Suffolk, come some fine examples of the gastropod Voluta Lambertii, one of them 9 1/2 inches (24 centimetres) long. The Red Crag furnishes Neptunea antiqua wound in a reverse direction to the usual one, Cyprina rustic, and Pholas cylindrica. In the Norwich Crag again appears Acila Cobbolidae, and Bittium reticulatum is again found in the St. Erth beds.

Oligocene. Shells of this Epoch come from the Hamstead or Hempstead Beds of Rupelian age, the Bembridge Limestone (Tongrian), and the Osborne and Headon Beds (Priabonian). These formations were deposited in an estuary where now are the Isle of Wight and the south of Hampshire. One may notice particularly Amphidromus [Bulimus] ellipticus wound in a reversed or left-handed coil, with many species of the gastropod genera Viviparus, Melania, Limnaea, and Planorbis evomphalus, and the lamellibranchs, Ostrea, Corbicula, and Volsella [Modiola]. The eggs of some large gastropod, Bulimus or an ally, will be seen preserved in Bembridge Limestone. A slab of Headon Limestone almost entirely composed of the shells of a fresh-water snail, Limnaea longiscata, is fixed on the wall.

Eocene. Here are the Bartonian shells of Barton, the Lutetian of Bracklesham, and the Landenian shells from the London Clay, Oldhaven Beds, Woolwich and Reading Beds, and the Thanet Sands. Except for the Oldhaven and Woolwich Beds, which contain both estuarine and marine fossils, all these formations are purely marine. Many of the specimens exhibited have been figured by G. A. Mantell, James Sowerby, F. E. Edwards, and others. A large specimen of Cardita planicosta is marked so as to explain the terms applied to the various parts of a lamellibranch shell. Here also is the curious burrowing lamellibranch

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Stirpulina [Clavagella] coronata, in which a shelly tube is formed round the long siphons while the shell-valves remain quite small and attached to the side of the tube. In addition to the gastropods here figured (Fig. 69), one may note the beautiful Voluta luctatrix, the large Hippochochrenes [Rostellaria] ampla, and Xenophora agglutinans.

Fig. 69.—Shells of Eocene marine Gastropods and Lamellibranchs. 

- a, Cyprina Morrisi, Thanet Sands of Herne Bay, Kent; the specimen figured by J. de C. Sowerby; the hole is due to a boring animal.
- b, Pholadomya virgulosa, London Clay of Bognor, Sussex; a small specimen. 
- c, inside, and d, outside of left valve of Crassatellites sulcatus, Barton Beds of Hampshire. 
- e, Cypraea Bowerbanki, Lutetian of Bracklesham, Sussex.
- f, Clavalithes longaeus, Barton Beds of Hampshire.
- g, the hinge, and h, the outside of right valve of Cardita planicosta, a small specimen, Lutetian of Bracklesham.

One of the carrier-shells that become covered with bits of stone, shell, and other foreign substances. Among London
Clay fossils are the long tubes formed by the boring lamellibranchs *Teredo* and *Teredina*, allies of the modern “ship-worm.” *Pyrula Smithi* and *Xenophora extensa* are noticeable gastropods. *Aporrhais Sowerbyi* is a gastropod from Oldhaven Gap. Characteristic of the Woolwich beds are *Corbicula [Cyrena] dulwichiensis*, *Corbulina regulbiensis*, *Pitharella Rickmani*, a freshwater shell, and *Ostrea bellocvacina*, of which a fine mass is exhibited. From the Thanet Sands come *Cyprina Morrisi*, *Pholadomya euneata*, and *Meretrix orbicularis*. In the middle of the Gallery stands a slab of “Bognor Rock” from the London Clay of Bognor, largely composed of shells or the casts of shells of *Cardita Brongniarti*, *Glycimeris [Acanacca] brevirostris*, *Volusella [Modiola] elegans*, *Voluta dannata*, *Pyrula Smithi*, *Natia hantoniensis*, and the annelid *Vermetus bognoriensis*. Other masses of the same age from Fareham, Hants, are placed on the wall; one is full of shells of the lamellibranch *Pinna affinis*, the other (Plate V.) contains in addition *Glycimeris brevirostris* and the smaller *G. decussata*.

**Cretaceous.** This series is rich in specimens described by G. A. Mantell, Samuel Woodward, James Sowerby, J. de C. Sowerby in Dixon’s “Geology of Sussex,” H. Woods, and others. Except those from the estuarine or lacustrine deposits of the Wealden series, all are marine. First come the shells from the Upper and Middle Chalk and Chalk Rock, in other words from those zones that are of *Senonian* and *Turonian* age. Among these the lamellibranchs of the genus *Inoceramus* are conspicuous and their disintegrated shells enter largely into the composition of the Chalk: *I. Cuvieri* may attain a width of 18 inches, some large pieces of its hinge are shown; *I. labiatus* gives its name to a Turonian zone; the original specimen of the widely distributed *I. Crippsi* is of historical importance. *Spondylus spinosus* with the long spines well preserved (Fig. 70), and the five-ribbed fan-shell *Neithica quinquecostata* (Fig. 71a), are characteristic of the Upper Chalk; the gastropod *Pleurrotomaria perspectiva* is Turonian. Here too are tubes of *Teredo amphibia*, and several specimens of the curious *Radiolites Mortonii*, one of the *Tudiste* (see p. 143).

Following on these are shells of *Cenomanian* age, coming from the Lower Chalk, Upper Greensand, and Chalk Marl. Here may be noted the cocks-comb oyster *Alectryonia carinata*, the scollops *Chlamys Beaveri* and *C. asper*, of which the latter gives a name to the “Pecten asper zone,” the
four-ribbed *Neithea quadricostata*, and *Syncyclonema* [*Pecten*] *orbicularis*. Among the Grey Chalk fossils from Folkestone, *Rostellaria Pricei* and *Aporrhais Mantelli* are remarkable.

Fig. 70.—A Lamellibranch shell, *Spondylus spinosus*, common in the Senonian Chalk of England. Natural size.

A beautiful series of silicified shells from Blackdown and Haldon in Devonshire, a set from the Red Chalk of Hunstanton in Norfolk, and another from the Cambridge Greensand, are all of Albian age. Of the same age are the
shells from the Gault clay of Folkestone, of Black Ven, Charmouth, and of Okeford Fitzpaine also in Dorset, all which are on the other side of the Case. Among the Gastropods the Aporrhaidae with their winged lips and the Scalaridae with their transverse ribs furnish the most striking forms. Among Lamellibranchs one may notice species of Actinoceramus (Fig. 71 b), Trigonia, Cucullaea, Liopistha [Thetis], Thracia, Pinna, Perna, and Protoeardia.

During the Aptian age was deposited the Lower Greensand of Hythe, Sandgate, and Faringdon. Here are casts of the borings made by Lithodoms and Pholas. Pleurotomaria gigantea, Alcetryonia macroptera, and Toveasia Lonsdalei are noteworthy forms. The last is a representative of the Diceratidae (see p. 143).

The Barremian or Urgonian age is represented in the South of England by the marine Atherfield Clay of the Isle of Wight and the brackish-water Punfield Beds of Dorset, both of which belong to its lower division. Exogyra sinuata was then a common oyster; among the shells exhibited one has been cut through to show its great thickness in old age. The long-hinged Perna Mulleti occurs in quantities in a special layer at Atherfield. Here are also shown the elongate Gervillia anceps, with species of Trigonia, Sphaera, Protoeardia, Cyprina, and Astarte, and, among the Gastropods, Ceratosiphon [Aporrhais] Fittoni, and Vicarya Pizudetana.

Next come fossils from various rocks in the east of England: the Speeton Clay of Yorkshire, the Tealby Series, and Claxby Ironstone, of Lincolnshire, all which correspond partly to the Barremian and partly to the preceding Neocomian Age. Among the more interesting lamellibranchs are Exogyra subsinuata (= E. Couloni), Trigonia ingens, and the large Camptoneotes [Peeten] cinctus from the Claxby Ironstone and Cucullaea donningtonensis. The Lincolnshire formations run over into the next Case, and are followed by fossils from the Wealden of Sussex and the Isle of Wight. The name Wealden is applied to a series of freshwater and estuarine formations deposited mainly during the Neocomian Age, and perhaps beginning at the end of the Jurassic Epoch. The river-shells include examples of Unio, of which the largest are the U. valdensis from Sussex and the Isle of Wight. In the middle of the Gallery is a polished slab of Petworth or Sussex Marble, composed of the shells of a freshwater snail, Viviparus...
Gallery VIII.

[Paludina] fluviorn, in a greenish calcareous cement. The marble occurs in layers, from a few inches to a foot in thickness, and is used for chimney-pieces, slabs, and columns. It may be seen in Canterbury, Chichester, and Salisbury Cathedrals, York Minster, Westminster Abbey, and the Temple Church. A similar stone is found at Bethersden in Kent. Other bands of limestone, often red, are full of Corbica [Cyrena] media. Pleurocera strombiformis (= Potamides carbonarius) is also noticeable.

Table-case 9.

Further west in the same great estuary were deposited the Purbeck Beds, the uppermost of which are by some geologists regarded as Cretaceous, while the lower are Jurassic. They extend from the Isle of Purbeck in Dorset to Brill in Buckinghamshire, and are found also at Brightling and Pounceford in Sussex. A small set of shells from them is shown, and in some species already observed in the Wealden, contains Corbica [Cyrena] parva. Unio Martini, Physa Bristovi, and the marine form Mytilus Lyelli. One of the characteristic beds is the Purbeck Marble, very like the Petworth Marble, and another is the Cinder Bed composed of Ostrea distorta, an oyster that probably owes its peculiar shape to brackish-water conditions.

Table-cases 9-14.

Jurassic. Leaving the estuarine formations of intermediate and uncertain age, we return to the marine series, exhibited under a number of stratigraphical groups, which will be taken in order, beginning with the highest. Many of these British Jurassic specimens in the Museum are of historical importance, having been described and figured by J. Sowerby, J. Phillips, Morris & Lycett, S. Stutchbury, R. Damon, W. H. Hulston, J. F. Blake, and others.

The Portland Oolite, from which is derived the name of the Portlandian Age, is worked for building stone in Dorset and Wiltshire. Two slabs of the hard Portland stone are shown; one of them contains shells of the lamellibranchs, Perna Bouchardi and Chlamys [Pecten] lamellosus; the other is almost entirely composed of shells of Cerithium concavum. In the middle of the gallery is a large block of the "Roach-bed," which is full of hollows, whence the shells have been dissolved by percolating water, leaving behind impressions and internal casts of the following species: Trigonia gibboa, Chlamys lamellosus, Ostrea expansa, Lucina portlandica, and Protocardiida dissimilis among lamellibranchs; Natica elegans and Cerithium portlandicum among gastropods. All these shells may be better studied in the Table-case, as well as
Fossil Shells.
Trigonia in Corallian rock from Weymouth.
Part of a Shell-bank in London Clay at Fareham, Hants.

[To face p. 135.]
MOLLUSCA (EXCEPT CEPHALOPODA).

Pterocera oceani, Pleurotomaria rugata, Sowerbya Dekei, Astarte rugosa, and many species of Trigonia. Note also the large borings of Lithodomus.

The shells of the Kimmeridge Clay are mainly from Weymouth, Dorset; Wootton Bassett, Wilts; the neighbourhood of Oxford; and Hartwell, Bucks. The following are noteworthy: the large Pleurotomaria reticulata, the D-shaped Ostrea deltoidea, the common and characteristic Exogyra virgula, Gryphaea dilatata with its thickened hinge, well shown in sections, a Gryphaea with a supposed pearl, Protocardia [Cardium] striatula, Astarte hartwellensis, various Trigonias, Goniomya literata, and Thracia depressa. A slab of Kimmeridge Clay filled with Ostrea laeviuscola is on the wall, between and below it is the fine block of Coralline Oolite with over 110 shells of Trigonia clavellata figured in Damon's "Geology of Weymouth" (see our Plate VI.).

The Coral Rag and Calcareous Grit are the chief British rocks formed during Sequanian or Corallian time; they stretch across England, with occasional breaks, from Weymouth to Buckinghamshire, reappearing in Yorkshire, and are rich in fossils. Among the gastropods are Bourguetia [Piasianella] striata, Pseudomelania heddingtonensis, Nerinea Goodhalli, and a spiny winkle Littorina muricata. Lamellibranchs are represented by Alectryonia gregaria, Chlamys vimineus, Ctenostreon pectiniformis, Mytilus pectinatus, Trigonia triqueta, and many others. The numerous borings of Gastrochaena and Lithodomus in the coral masses bear witness to a shallow sea. It is interesting to contrast the richness and variety of the molluscan fauna that lived in the Corallian sea with the comparatively few species found in the clays above and below.

The clay below is the Oxford Clay, well developed in Oxfordshire, and forming a more continuous band across England from Dorset to Yorkshire than do the Corallian limestones. Within this tract the fossils are entirely marine, and those exhibited come chiefly from Weymouth, Christian Malford and Chippenham in Wiltshire, and Scarborough. The delicate character and shelly constitution of the Wiltshire specimens contrast with the coarser and stony appearance of the others. Characteristic forms are Alaria trifida with its long processes, the delicately spined Spinigera spinosa, Nucula ornata, Volsella [Modiola] cuneata, and Pleuromya recurva. Under Ostrea and Gryphaea are to be seen shells that have grown upon Trigonias and an ammonite and have assumed...
their ornament. Adjoining these are shells from the Kellaways Rock, a brownish sandstone at the base of the Oxford Clay, giving a name to the Callovian Age. Among these are *Gryphaea bilobata* with a curious fold on its side, *Isocardia minima*, and *Goniomya V-scripta*.

The shells of *Bathonian* age are arranged under the formations from which they come. First are those from the widely distributed rubbly limestone called *Cornbrash*, which has yielded the scaphopod *Dentalium annulatum*, the gastropods *Bulla undulata*, *Pseudomelania vittata*, and the lamellibranchs *Chlamys* [*Pecten*] *vagans*, *Lima duplicata*, *Pseudomonotis* [*Avicula*] *echinata*, and several Trigonias. Many of the specimens are figured in Blake's "Monograph of Cornbrash Fossils." There are very few shells from the *Forest Marble*, but among them is a fine example of *Trigonia detrata* figured by Lycett. In the equally small series from the *Bradford Clay*, one may notice a set of *Oxytoma* [*Avicula*] *costata*. The *Great Oolite* shells are mostly from the Oolitic freestone. Here is the type-specimen of *Pterocra* *Wrighti*, a fine winged shell; then several species of *Purpuroides*, among which *P. morrisea* with its heavy spines is conspicuous; several limpets, *Patella*, testify to the rocky nature of the sea-floor; *Nerita rugosa* seems to show bands of colour, and such are still more evident in a large form of *Natica*. Among lamellibranchs one may note *Lima cardiiformis*, *Pteroperna costatula*, *Pinna ampla*, *Parallelodon* [*Macrodon*] *hirsonensis*, and the massive *Pachyrisina* ("thick support") *grande*. The Stonesfield Slate yields *Trigonia impressa* and *Pinna cuneata*. The small series from the *Fuller's Earth* includes *Volsella* [*Modiola*] *imbriicate* and *Ceromya plicata*.

Next come shells from the variable series of marine limestones known as the *Inferior Oolite*, deposited mainly in *Bajocian* time. The chief localities in the south-west of England here represented are Dundry near Bristol, Halfway House near Yeovil, Bradford Abbas, and Leckhampton Hill near Cheltenham. These have yielded a fine series of *Pleurotomaria*, many species of *Amberleya*, *Pseudomelania*, *Purpurina*, *Delphinula*, *Cerithium*, *Cirrus*, *Nerinaea*, *Alaria*, and other gastropods. Certain species of *Amberleya*, *Cerithium*, *Onustus*, and *Neridomus* acquired an interest a few years ago from their resemblance to some shells now living in Lake Tanganyika; but it is not now imagined that the animals themselves had the same structure. Similar
repetition of outward form at different epochs has already been noticed among the Brachiopoda (p. 115). Among lamellibranchs from the same rocks are Alectryonia, Lima, Ctenostreon, Inoceramus and the large Trichites with its thick shell, a favourite haunt of Lithodomus, as various specimens show. Another specimen shows a Lithodomus burrow in the floor of the Oolitic sea, here formed of black Carboniferous Limestone. Then follow many species of Trigonia, Astarte, Pholadomya, Ceromya and other genera. Among all these specimens may be noticed others from the Ironstone of Duston, Northamptonshire, and from the Collyweston Slate. In Yorkshire the Bajocian series includes beds of estuarine origin, furnishing such forms as Unio and Anodon.

The lamellibranchs from the **Lias** are arranged under the three divisions of that marine formation: the Upper, of Toarcian age, the Middle, of Pliensbachian, and the Lower, of Sinemurian. Here one should notice Leda ovum, which gives its name to a horizon in the Upper Lias, Volscella [Modiola] scalprum from the Middle and Lower Lias, the very familiar Gryphaea incurva, and the equally massive Hippopodium ponderosum, both from the Lower Lias (Fig. 72).

![Fig. 72.—Lamellibranch shells from the Lower Lias. a, Hippopodium ponderosum; b, Gryphaea incurva. Natural size.](image)

**Oxytoma** [Avicula] cygnipes, from the Cleveland ironstone beds of Yorkshire is a fine shell. The gastropods are all placed together, since they are few in number, and nearly all comprised within three genera: Encyclocyst, Cryptaenid, and Pleurotomaria (Fig. 73). This last contains some large shells, those of *P. anglica* being most numerous.

**Trias.** The Mollusca come chiefly from the **Rhaetic**
Gallery VIII. beds of Beer, in Devonshire, and Westbury, near Bristol. 

Avicula contorta is the best known, as giving a name to a widely distributed horizon of Rhaetic Age. Like the Rhaetic shells generally, it is relatively small, perhaps in consequence of brackish water. Monotis decussata and Chlamys vuloniensis are also important. There is a small but interesting series from the Keuper Marls of Warwickshire. The Conchylial Age has no shell-bearing rock in this country.

Table-case 14. Permian. This Epoch is represented by marine shells from the Magnesian Limestone of Durham and the red marls near Manchester. Note Monotis speluncaria and Byssarca striata from the former, and the tiny Rissoa and Turbo from the latter. Bakewellia antiqua comes from both localities, and from Tyrone as well.

![Recent and fossil shells of Pleurotomaria.](image)

Fig. 73.—Recent and fossil shells of Pleurotomaria. a, P. Quoyana, now living in the West Indies; b, P. platyspira from the Middle Lias of France. The slit s receives the projecting anus, and, as the shell grows forward, is filled up by shell-substance. Both figures are less than natural size.

Table-case 15. Carboniferous. The shells of this Epoch come mainly from the Coal Measures and the Mountain Limestone, the rocks of Middle Carboniferous or Moscovian age having yielded few mollusca in this country. The Coal Measures, though largely of fresh or brackish water origin, contain many marine bands; the Lower Carboniferous rocks are all marine. The fossils have not here been separated according to age or rock or habitat. It will, however, be noticed that the Coal Measure fauna, and particularly the freshwater elements in it, occur among the lamellibranchs, whereas the gastropods are almost all from the Mountain Limestone. Among the Coal Measure fossils are many described by Sowerby in Prestwich's classical memoir on Coalbrookdale, and many described by Dr. Wheelton Hind in the Monographs of the Palaeontographical Society. The fresh-water forms include Anthracomya and Carbonicola [Anthracosia]
MOLLUSCA (EXCEPT CEPHALOPODA). 139

(Fig. 74a), which had a very wide distribution during this Age (Ouralian). Marine forms are Aviculopecten papyrus, Pinna costata, and Nuculana attenuata from Scotland. Among lamellibranchs from the Mountain Limestone are the peculiarly shaped Conocardiium and various species of Cardiommorpha. Posidonomya Becheri is widely distributed in the Lower and Middle Carboniferous, often occurring in abundance and giving its name to certain beds in this country and on the Continent (Fig. 74b). The gastropods of Lower Carboni-

Fig. 74.—Carboniferous Lamellibranchs. a, Carbonicola [Anthracosia] robusta, Upper Coal Measures; \( \frac{1}{2} \) natural size. b, Posidonomya Becheri, Middle Carboniferous, North Devon; natural size. (Table-case 15.)

ferous age contain many genera not hitherto noticed; among these Macracheilus, the loose-coiled Phanerotinus, Euomphalus, and the wide-mouthed Bellerophon are conspicuous; a large shell of Naticopsis retains the operculum, lofty-spired forms are represented only by Murchisonia and Loxonema; Morulonia is a fore-runner of Pleurotomaria; Platyceras [Capulus] often occurs on the ventral surface of crinoids, living on the food-particles that they reject.

Devonian. During this epoch the sea existed where now are found fossils similar to those found in the Devonian limestones and slates of the Continent. The rocks that were at the same time being deposited over the rest of the British area are for the most part known as Old Red Sandstone, and seem to have been formed in lakes or estuaries. Evidence of this is furnished by the large shells related to the modern fresh-water mussel, and called Archanodon Jukesi. Specimens of this, the oldest non-marine lamellibranch as yet known, come from rocks of Upper Devonian age in Northumberland, Monmouth, and Kiltoran in S. Ireland. The marine forms from the south-west of England are mostly Middle Devonian gastropods from the collections of the late W. Vicary and
J. E. Lee, and have been described by the Rev. G. F. Whidborne in the Monographs of the Palaeontographical Society. Here are most of the genera already observed in the Carboniferous series. Among lamellibranchs Cardiola retrostriata is important to the stratigrapher.

Silurian. The Ludlovian Age is represented by fossils from Ledbury, Ludlow, and Kendal, the Wenlockian by fossils from Dudley and Benthall Edge, the Valentian by a few Bellerophons from the Llandovery beds. In addition to the Palaeozoic genera already mentioned, one may see here Pterinacea (Fig. 75 b), Orthonota, and Grammysia among the common lamellibranchs. The gastropods include numerous forms allied to Euomphalus, one of them, Polytopina, preserving the operculum, also the slightly curved Ecculimphalus, and Tremattonotus (Fig. 75 f), the "pierced back" ally of Bellerophon. The Amphineura are represented by Helminthochiton.

![Fig. 75.—Lower Palaeozoic Lamellibranchs and Gastropods. a, Cardiola interrupta; b, Pterinacea Danbyi; c, Bellerophon cambriensis, a reconstructed side view; d, Platyceras [Acroculia] haliotis; e, Maclurea Logani, the type-specimen. a, b, d, and f, are Silurian from the Wenlock beds of Dudley; e is Ordovician, from the lower Llandeilo of Ayrshire; c, Cambrian, from the Upper Festiniog beds of Dolgelly. a and c are nat. size; b, d, e, f, \( \frac{3}{2} \) natural size. (Table-case 16.)](image)

Ordovician. The rocks of this Epoch, formerly classed as Lower Silurian, have furnished molluscan fossils from the Caradoc, Llandeilo, and Arenig divisions. Those exhibited come mostly from Wales and Shropshire, but there are a
few from Tyrone, and a fine series of the loosely-coiled Malheurea from Scotland (Fig. 75 e).

Cambrian Mollusca are few, and in Wales, where the rocks are exposed, the shells are poorly preserved. Glyptarca primaeva and Bellerophon cambriensis (Fig. 75 e) may be noted.

In a special Table-case at this end of the Gallery are placed together the British and foreign specimens of CONULARIDA. These, which are almost exclusively Palaeozoic, are sometimes regarded as forming a Sub-Order of Pteropoda; but now that the origin of the Pteropoda from Opisthobranch Gastropoda appears certain, this view is scarcely tenable. Some of the simplest forms are found in the Lower Cambrian of North America, and of these Salterella, Helena, and Hyolithes are exhibited. They are thick-walled tubes, straight or bent, smooth or striated, ending in a point, and are generally placed in a family with Torellella from Sweden, which has a shell of phosphate of lime. Salterella has also been found in West Australia, and Hyolithes occurs in Ordovician rocks. The Devonian Coleoprion is probably an allied form. The shell of Hyolithes [Theca] and a few similar forms is composed of carbonate of lime, is conical, straight, or curved, triangular, elliptical, or flattened in cross section. Its surface is smooth or striated, and it is closed by an operculum. The narrow end of the shell is often divided by cross partitions. Species of this genus lived throughout the Palaeozoic Era, but were most abundant in its earlier half. The shell of Tentaculites is an elongate cone, ornamented with rings. It often begins in a small bulb, and the earlier portion may be partitioned as in Hyolithes. Its thick wall is composed of two layers. The genus abounds in Silurian and Devonian rocks, often being the chief constituent of certain beds. Conularia is the most widely distributed in space and time, and contains the greatest number of species and the largest individuals of all Conularida. The shell is elongate and four-sided. Each face is divided lengthwise by a groove (Fig. 76 e) and ornamented with parallel ridges, which slope upwards toward
this groove. The margin of the shell opening is folded in-
wards (Fig. 76 b). A few specimens of Cambrian age are
shown. The largest are Ordovician, those from the Grés de
May in Calvados being noteworthy. Several species occur
in Silurian rocks, and there will be noticed some specimens
well preserved in nodules from the British Coal Measures.
Some from the Permo-Carboniferous rocks of New South
Wales are exhibited. A single species is also found in the
Trias and in the Lias.

We return now up the side of the Gallery, taking the Wall-
cases with shells from foreign localities in reverse order.

Wall-cases 8 & 9. **Palaeozoic.** On the bottom slope of Case 9 are fossils
of Cambrian, Ordovician, and Silurian age, mainly con-
sisting of the Barrande Collection from Bohemia. Among
the few North American specimens is *Megalomus*, which
forms thick banks in the Guelph Limestone. Devonian
fossils, mostly from Germany, follow on the same slope.
The irregularly-shaped *Platyceras [Aeroculina] Protei* from
Mayenne is worth notice. The middle slope contains the
Carboniferous shells, among which those of the De Koninck
Collection from Belgium are the most numerous. Here the
*Chiton* claims attention. These are succeeded by a series
from the Permo-Carboniferous rocks of Tasmania and New
South Wales, many described in Strzlecki’s book on the
latter colony (1845). Then come Permian fossils from the
Zechstein of Saxony, with a few from Africa, among which
is a specimen of a freshwater Lamellibranch collected by
Henry Drummond on the N.W. shore of Lake Nyassa.

Wall-cases 9-5. **Mesozoic.** The **Trias** begins with a series of marine
lamellibranchs from the Malay Peninsula, described in the
Proceedings of the Malacological Society (London, 1900); the
next is the valuable collection from St. Cassian in the
Tyrol, formed and described by A. v. Klipstein. The **Lias**
of Germany and Northern France is fairly well represented;
among the gastropods is also a set of *Lithothrochus Humboldti*
from Peru. Wall-case 7 contains the shells of **Oolitic**
(Aalenian to Portlandian) age. At the end are placed
three interesting series recently described: from Singapore
(Geological Magazine, 1906); from Borneo (Proceedings of
the Malacological Society, 1903); and from Madagascar
the bottom slope is a fine series from the Bajocian and
Bathonian rocks of Normandy agreeing closely with that
from British localities of corresponding age; the large
**Pleurotomaria Proterus** and the great limpet **Patella Tessoni** cannot fail to attract attention. The shells on the middle slope range from Callovian to Corallian, and among the lamellibranchs **Diceras** with its two curly horns should particularly be noticed. Shells from the Solenhofen lithographic stone and other rocks of Kimmeridgian age, with many from the Portlandian, occupy the top slope; most are from French localities.

The Cretaceous series from Neocomian to Cenomanian comprises specimens from all quarters of the globe. The most remarkable are the Diceratidae and their descendants. **Diceras** and **Toucasia**, already noticed, are fixed to the sea-floor by one valve, just as is their more ordinary relation **Chama**. In them, however, this fixed valve becomes very long, and is often twisted, while the other valve is smaller and may be reduced to a simple lid, as in **Requienia**. In **Monopleura** and **Caprotina** it will be seen that the fixed valve is less twisted, and that it grows upwards like a cup-coral. Like a coral, too, its lower part is often cut off by partitions. The likeness to a coral is still greater in **Hippurites**, of which a simple form is on the bottom slope, and more complicated forms on the top slope. Here too is **Radiolites**, with a massive shell-wall broken up into cubical cavities. Meanwhile the other valve has become a stout lid with great projections inside, presumably for the attachment of muscles. The **Rudistae**, as these later types are called, reach their acme in Turonian and Senonian times with marvellous forms whose true nature must for ever have remained a mystery had we not been able to trace their gradual evolution. In Southern Europe and in the East and West Indies (e.g. **Barrettia**) they formed reef-like masses, now known as **Hippurite Limestone**. The remainder of the Upper Cretaceous Mollusca are in Case 5A, and adjoining are specimens of **Inoceramus expansus** from S. Africa.

**Tertiary.** The very complete series from the Paris Basin has an added interest from the fact that most of the specimens were purchased from the eminent fossil conchologist, G. P. Deshayes. The Eocene shells are closely allied to those found in England, but are better preserved, and the species are more numerous. The map showing the formations and localities is taken from "The Eocene and Oligocene Beds of the Paris Basin," by G. F. Harris and H. W. Burrows (London, 1891), a work that gives also a full list of the Mollusca. Below it are fine specimens of the...
Gallery VIII. Middle Eocene *Campanile* (*Cerithium*) *giganteum*, while in Case 4 is a longitudinal section of the same fossil, showing the shelly pleats upon the columnella. Next come the Tertiary shells from Bordeaux, from Muddy Creek, Victoria, and from South Australia. Pliocene shells from Italy are followed by Post-Pliocene shells from raised beaches in Florida, Australia, and elsewhere. A "Catalogue of the Australasian Tertiary Mollusca," by G. F. Harris, was issued by the Trustees in 1897. On the bottom slopes of Wall-cases 1–3 are temporarily placed small series recently acquired: Eocene shells from Northern Nigeria (Annals and Mag. Nat. Hist., 1905); Eocene shells from Somaliland (Quart. Journ. Geol. Soc., London, 1905); Miocene shells from the Azores and from Malta; Post-Tertiary and Tertiary shells from the region of the Dardanelles (Quart. Journ. Geol. Soc., London, 1904); Tertiary shells from Patagonia; Post-Tertiary shells from Angola, W. Africa; and Tertiary shells from Alaska. All the preceding series are marine, but there is also a series of Post-Pliocene shells of estuarine character from the Pampean formation near Buenos Ayres. Adjoining these is the collection of marine Miocene shells from Maryland, described by Thos. Say in 1824 (Journ. Acad. Nat. Sci., Philadelphia).

Finally, next the entrance to the Gallery, is a small representative set of Pteropod shells (see p. 125), mostly from the Upper Tertiary rocks of Italy.

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**Class CEPHALOPODA.**

A whole Gallery is occupied by fossils belonging to the remaining Molluscan Class, the Cephalopoda (Head-feet). The meaning of this name is obvious to any one considering such well-known living examples of these marine molluscs as the octopus, the cuttle-fish, the squid, and the nautilus (Fig. 78), in all of which the mouth is encircled by arms or tentacles, the altered representatives of part of the molluscan foot. Just behind these are the two eyes. The hind-part of the body consists of a rounded or sometimes elongate sack containing the viscera and called the visceral hump. Part
of the skin forms a mantle-fold on the under surface and encloses the gills. In front of the gills another part of the foot is folded together to form a funnel through which the water that has passed into the gill-chamber can be forcibly squirted out. By this means the animal can be driven backwards through the water (Fig. 87). The mouth is armed with strong horny jaws, shaped like a parrot's beak (Fig. 78 a–c). The power of locomotion the concentration of the chief organs in the forepart of the body, and the strong jaws, show that these animals can prey upon others, even on those with hard coverings, and lead one to regard them as more highly developed than the other Mollusca. All these features can be recognised in the exhibited specimens and models of living cephalopods. Further study of these specimens will show certain differences between the forms that have been mentioned. Thus the cuttle-fish, squid, and octopus have rather long distinct arms, and these are furnished with suckers strengthened by horny rims; in the nautilus

![Diagram of Cephalopoda]

Fig. 78.—Cephalopoda. a, b, c, Rhyncholiths, fossil beaks of a nautilus from the Miocene of Malta; nat. size; d, the Pearly Nautilus in its shell, part of which has been cut away so as to show the chambers and the siphuncle running through them. On the animal are seen the visceral hump (a), the muscle and ring of attachment (g), the eye (s), the funnel (k), the tentacles (p), and the hood (n). e and f, the squid, Loligo; the shell (e) is seen from the back, the animal (f) from the under side; d, e, f, are much less than natural size.
(Fig. 78 d) the arms are represented by short lobes, and these are furnished with tentacles that can be withdrawn into sheaths. The eyes in the long-armed forms are conspicuous and are as highly developed as those of a vertebrate animal; in the nautilus one can detect only a small opening leading into a chamber that acts like the simple optical apparatus known as a pin-hole camera. If the mantle-folds be separated and the gills exposed, it will be seen that there are four gill-plumes in the nautilus, but only two in the other forms mentioned, and indeed in all other cephalopods now living; for this reason, and in view of the other differences, Sir R. Owen divided the Cephalopoda into two divisions: Tetra-branchia (four gills) and Dibranchia (two gills). We know that some fossil cephalopods, clearly related to the modern Dibranchia, had long arms, suckers, and large eyes; but we cannot be certain that all of them had only two gill-plumes. Other fossil shells resemble that of the nautilus, and in some of these cases the animal most probably had four gill-plumes; but in other cases there are no grounds for any such assertion. Therefore Owen's classification is unsuitable for fossil cephalopods.

In order to arrange systematically the large number of extinct cephalopods one must consider chiefly those parts of the animal that can be preserved in the rocks. The beaks, the horny rings of the suckers, and similar structures have already been mentioned (p. 123), as also the fact that some fossils found in very fine clays have preserved even the muscles of the mantle (Fig. 86 a). As a rule, however, one finds only the shell with which most of the Cephalopoda are provided. This shell has in the course of geological time undergone many changes, and has been modified in several directions. It is moreover intimately related to the structure and development of the whole animal, and therefore furnishes an excellent basis of classification.

Specimens, models, and drawings have been arranged to show the history and relations of the cephalopod shell, and to these attention should first be directed. As in all Mollusca the shell is primitively secreted by the mantle or skin of the visceral hump, and, at its edges, by the backwardly turned folds of the mantle. Originally then the shell follows the shape of the visceral hump, and we may suppose that, in cephalopods older than any which are known to us, it was a somewhat conical cap, not unlike the shell of some uncoiled gastropods. Whether these were
active predatory animals, or whether they were sedentary and possibly attached by the shell, is uncertain. But certain it is that at an early period the hump was drawn out into a long visceral cone, and that the shell acquired a similar shape. Then followed a mode of growth very common in sedentary animals that form a tubular shell, and already observed in corals, worms, bryozoans, and gastropods. The mollusc continued to build up the shell around its opening, and thus formed a long tube. As the animal moved along this tube, the visceral cone was pulled away from the shell-

![Figure 79](image)

**Fig. 79.**—Primitive Cephalopod shells: *Endoceras*. *a*, The end of a shell, broken at its apex; *b*, the same cut in half showing the chambers (sc), the swollen end of the wide neck-tube (cc), and calcareous substance deposited by the end of the siphuncle (cd). *c*. Fragment of another shell cut in half, showing the chambers separated by septa (s), the large neck-tube (sc), a sheath (sh), and the endosiphon (en). Natural size. (*a*, *b*, after Holm; *c*, after Foord.)

wall; but its skin went on secreting shell-substance, and formed a partition shutting off the space between the visceral cone and the outer wall. In *Endoceras*, here shown (Fig. 79), the visceral cone remained attached to the end of the outer shell, and shrank at a little distance from the apex, so that the partition or septum does not go right across the shell, but shuts off a chamber at the side. Shrinkage then took place a little higher up, and another chamber was formed. By the continuation of this process there arose a series of chambers entirely shut off by septa, and the down-turned portions or necks of these septa formed a long tube in which
lay the much constricted visceral cone. Examination of polished sections across *Endoceras* and allied forms shows within this septal neck-tube yet another series of structures, called sheaths, and somewhat like a pile of funnels stuck one inside the other (Figs. 79 c, 80 a). These indicate that, as the animal advanced in its shell, its viscera naturally went with it, and towards the void thus left the walls of the visceral cone were still further sucked in. Thus there tended to

arise a narrow and empty tube—the siphuncle. The walls, however, were stiffened with lime and did not completely yield to the suction, so that, when the animal again advanced, the inner layers of the skin were torn away from the outer ones. These inner layers thickened and stiffened in their turn, and the process was repeated. Thus arose the thin tube, sometimes called the endosiphon, and a series of sheaths attached to it.

In *Orthoceras* and similar forms there is an advance on
this mode of growth (Fig. 80 b). The shrinkage appears to have been greater from the first, and thus the septa stretch across the conical shell, dividing it into a series of chambers, and leaving only a narrow neck-tube in which there are no sheaths. The part of the shell-cavity in front of the last-formed septum is called the body-chamber, and in it was the main mass of the animal. From the visceral hump, however, proceeded the fleshy siphuncle, passing through each septum to the apex of the shell. The line along which the septum is attached to the shell-wall is called a suture, and in these simple forms passes regularly all round the shell.

There are numbers of straight shells of this simple type, but those that are the most completely known may be divided into two groups by the presence or absence of a small, more or less globular, initial chamber. This, which is generally separated from the next chamber by a slight constriction, is called the protoconch (first shell) and believed to be the shell of the embryonic cephalopod. It is well seen in some specimens of Bactrites (Fig. 81 b). Often this protoconch seems to have been lost in the adult, and in its place is seen only a scar or cicatrix denoting its former presence (Fig. 81 c, c, m).

At an early period in the history of the cephalopod race the shell began to curve, and this curvature increased until the shell was coiled on itself. Such a coiled shell was far more manageable than the long shell of an Orthoceras, and was less liable to damage. And so it is found that the long straight shells gradually die out and give place to coiled shells. Now, just as there were two types of straight shells, so were there two of coiled shells: one with a protoconch, as may be seen in models of early goniatites (Fig. 81 n); the other without a protoconch, as shown by the model of Nautilus (Fig. 81 a–e). Further examination of the shells of these two types reveals other differences. The early coiled shells with a protoconch are long, narrow, smooth, with septa usually far apart, and with a long deep body-chamber (Fig. 81 n). Those without a protoconch are short, broad, often with a longitudinal ornament, with septa relatively close together, and with a shallow body-chamber (Fig. 81 a). In later forms of these two types other differences appear, such as will be realised by comparing an ammonite (which is one of the former series) with a Nautilus (Fig. 82). Generally speaking the siphuncle of an ammonite is close to the outside of the coiled shell; the edges of the septa are folded, so that the sutures are compli-
Fig. 81.—The first-formed portions of various chambered Cephalopod shells. a, b, c, *Nautilus pompilius*: a, side view; b, front view, the apex broken to show siphuncle passing through septa; c, apex, showing scar left by protoconch. d, a straight shell of Triassic age, with a plug (p) at its apex. e, a curved Carboniferous nautiloid, *Meloceras*, with scar (ci). f, g, h, front, side, and upper views of protoconch of an ammonite, *Cosmoceras*. i, Protoconch and first chamber of a belemnite. k, Protoconch and five chambers of *Bactrites*, the shell partly removed, showing siphuncle. l, m, the first two chambers of an allied form from which the protoconch has disappeared, leaving a cicatrix, shown in m. n, Protoconch and first five chambers of *Mimoceras compressum*, a goniatite. o, Protoconch and part of two chambers of *Spirula*, the shell-wall partly broken away. In g, i, k, o, the siphuncle is denoted by si. c is natural size, the rest greatly enlarged. (From Foord & Crick. After Barrande, Hyatt, & Branco.)

cated; the edges of the septa round the siphuncle form a collar projecting towards the opening of the shell; the outside of the shell is ornamented by ribs, folds, or tubercles, radiating from the centre of the coil. In a nautilus, on the other hand, the siphuncle is near the centre of the septum; the edges of the septum are but slightly curved or bent, so
that the sutures are simple; there are backwardly projecting septal necks; the outside of the shell is smooth or slightly ribbed. For these, and other reasons, the cephalopods of which we have been speaking are divided into two Orders: the Ammonoidea and the Nautiloidea.

Returning now to the straight shells, we shall note that none of those without a protoconch survive the Palaeozoic Era, but that all give place to coiled Nautiloidea. Those with a protoconch, however, do not all give place to Ammonoidea, but some of them begin another line of evolution. In these the chambered shell becomes shorter, and it is believed that folds of the mantle were turned back right over the shell to its very apex, thus affording a

![Diagram of a Nautilus and Ammonite]

**Fig. 82.**—A nautilus and an ammonite. *a* is a plaster cast of the inside of the shell of *N. pompilius*, and may be compared directly with the similar internal cast of (*b*) the ammonite. *s*, suture; *m*, mark left by muscle of attachment; *l*, lines left by this mark in its previous positions. Less than natural size.

... protection to the shell and to its protoconch. These mantle-folds continued their activity in secreting shell-substance, and so there were deposited outside the apex of the shell a number of layers, forming a solid guard (Fig. 85 *a*). Such forms first appear with certainty in the Trias, e.g. *Aulacoceras*. In these the guard is short, and the chambered shell-cone, when found without it, might be taken for an Orthoceras; but it must be noted that the siphuncle is at the margin as in Ammonoidea. It is probable that, like their living descendants, these animals lived an active life and frequently shot backwards by discharge of water from the funnel; thus the value of the guard is obvious. The further history of these forms shows the rise of various...
modifications assisting in one way or another this actively moving life. The main result in all, however, is the enclosure of the shell-cone by the mantle-folds and its reduction in size, so that from being an external protection to the animal it becomes an internal support for the mantle and for fin-like appendages. This does not mean that it becomes an internal skeleton in the true sense of the word; for its relations to the visceral hump always remain the same: it is always outside the true body-wall. Cephalopods in which the shell is thus ensheathed by the mantle constitute a third Order, corresponding, so far as living genera are concerned, to the Dibranchia of Owen. For this Order the name COLEOIDEA (sheath-forms) has been proposed; the name BELEMNOIDEA has also been used for it, but is more frequently restricted to one of its subdivisions.

In their further history the Nautiloidea undergo no changes of importance. It will be well, however, to study the preparations showing some points in their structure. Among these are some internal casts and some portions of shell, showing the scar made by the attachment of the muscles that fix the body to the wall of the shell (Fig. 82 a).

The Ammonoidea also do not diverge greatly from the general outlines sketched above. They do, however, break up into a number of lines of descent which it has proved exceedingly difficult to unravel. In this study great importance is attached to the foldings of the suture, and some attention may profitably be given to the models showing how the chief types of suture are gradually developed from the simple type of the older forms. A fold of the suture directed towards the opening of the shell is called a saddle, and a fold in the opposite direction is called a lobe. The absence of folding in the earlier sutures characterises forms known as Asellati (unsaddled). Folding, when once started, begins at an early stage in the life-history of each shell, and is manifest even in the suture between the protoconch and the first chamber (Fig. 81 f-h). Forms in which this suture has one broad saddle, and in which external lobes and saddles appear gradually in the later sutures, are called Latisellati (broad-saddled). Those in which this suture has a narrow saddle in the middle line, bounded on each side by a lateral lobe, and this again by a lateral saddle, are called Angustisellati (narrow-saddled). The further developments of these foldings must be studied in the models and specimens, and in professed text-books of palæontology.
The same models show the change from septal necks to the collars characteristic of later Ammonoidea.

Among other specimens shown in this Case are those illustrating the lid or operculum of the ammonite to which, when it was supposed to be an independent shell, the name *Aptychus* was applied (Fig. 83). Some of these specimens preserve the aptychus in its natural position completely closing the opening of the shell. The aptychus consists of two equal halves (sometimes not divided), and there is no good reason for doubting that these were formed as calcifications of that structure which in a nautilus is called the hood (see the preparation of *Nautilus* between Wall-cases 13 and 14).

Returning to the specimens and models that illustrate the structure of the *Coleoidea* or Belemnoida, we study first a belemnite from the Lower Lias of Charmouth that was described by Huxley in 1864 (Memoirs of the Geological Survey). This shows that the guard is small as compared with the size of the whole animal, and that only a small part of the shell-cone is contained within the guard (compare Figs. 84 and 85 a). The wall of the body-chamber extends a considerable distance in front of the chambered portion of the shell (or phragmocone) and is known as the pro-ostracum. In front of this is the head of the animal, in which may be seen the beaks; and in front of this again are six arms bearing hooks. Other hookless arms may or may not have been present. Upon this and other specimens is based the exhibited diagram of a belemnite animal. This is seen from the under side, with a part of the skin removed from the middle region and the shell sliced down the middle. In addition to the short hooked arms seen in the fossil just mentioned, are shown two long arms, of which, however, traces have not been found. The funnel is seen emerging in front of the mantle-folds, and on each side of it is an eye. Behind this is the body-chamber, protected on the back by the pro-ostracum, not seen in this view. The viscera are not shown, since nothing definite is known about them, with the exception of a small bag, from which a tube leads forward and opens into the funnel. This bag is filled with a carbonaceous substance which, as known in living cephalopods, can be ejected in the form of a dark ink. It serves to
obscure the water and protect the animal in its flight from an enemy. A paint made from this ink is called sepia, after the Latin name of the cuttle-fish. The ink-bag is often found in these fossils, and its contents can still be used as a paint. Behind the body-chamber are seen the phragmocone and the guard, and stretching along the sides of the whole shell are expansions of the mantle, forming fins. Belemnites

![Diagram](attachment:image.png)

Fig. 84.—Restoration of the animal and shell of belemnites. a, back view; b, front or under view. (After d’Orbigny.)

having this general structure and a solid guard lived to the close of the Cretaceous Epoch, when they disappeared.

Whereas the Ammonites left no descendants, the belemnites appear to have become changed into other forms. One of these, Belosepia, is found in the Eocene London Clay (Fig. 85 d). Here the guard has become reduced in size, and the septa stretch in an upward curve from the apex of the shell (corresponding to the belemnite phragmocone) to the front of the pro-ostracum. They are numerous and close together. This form leads to the ordinary Sepia or cuttle-fish, of which two glass models are shown. The shell of this
animal is the familiar cuttle-bone (Fig. 85 e). Viewing it from the back one sees at its end a small point (the muco) corresponding to the guard, and in front stretches a broad shelly plate, like a pro-ostracum. This, however, when viewed from the other side, is seen to be covered by a mass of thin shelly plates, which correspond to the septa more clearly seen in Belosepia.

Another Eocene fossil that appears to be descended from the belemnites is Spirulostra (Fig. 85 b). In some of the later belemnites may already be observed a shortening of the guard and a curvature of the phragmocone, processes which tend to reduce the unnecessary length of the shell. In
Spirulirostra the apex of the chambered shell is coiled quite round, and a part of the sheath has come to lie above it; the rest of the sheath is greatly reduced and ends in a sharp point. By further coiling of the shell, by reduction of the pro-ostracum, and by the final disappearance of the guard, was produced the coiled shell of Spirula (Fig. 85 c). This is very like a shell of the earlier coiled Ammonoidea, but the siphuncle is on the inner side of the coil, the septal necks pass backwards, and the shell is very thin. It is almost entirely enclosed by the mantle.

Meanwhile a distinct line of evolution was progressing. Among the beautifully preserved fossils of the Oxford Clay of Christian Malford is the specimen of Belemnoteuthis antiqua described by G. A. Mantell (Fig. 86 a). Here are to be seen 10 short hook-bearing arms, the head with its well-developed eyes, the folds of the mantle, the fins, the ink-bag,
and the phragmocone. This last is relatively much wider than in the belemnite and is covered by no solid guard, but by a thin coating of horny substance deposited by the mantle, which is still seen surrounding it (Fig. 85 f). The reduction of guard and phragmocone has progressed even further in *Geoteuthis brevipinsis*. A specimen of this from the same rock preserves the 10 very short arms, mantle, and tail-fins, and shows a broad thin shell such as in the squids and calamaries now living is called a pen, with very little trace of any chambered portion (Fig. 85 g, h, i). A similar but narrower pen is found in *Plesiotethis prisca*, of which specimens from the Solenhofen Stone are exhibited, one of them with the soft parts also indicated (compare Fig. 86 b). These pens should be compared with that of the squid, *Loligo vulgaris*, which is shown beside a glass model of the complete animal (Fig. 78 e). Other models of living cephalopods with similar shells are exhibited, but cannot here be discussed. It is enough to realise that in one line of descent of these forms with ensheathed shell the chambered cone, long protected by a stout covering, retained its calcareous and septate nature, while in another line it became horny, and ultimately lost its septa.

There was however another line of evolution, the origin of which is difficult to trace, because one of its conspicuous characters was the complete loss of the shell. Another character was the absence of the long arms, reducing the number to eight. These forms therefore are called Octopoda, in opposition to the ten-armed squids and cuttle-fishes, which have been called Decapoda. Study of the early development of an octopus teaches us that its ancestors must have had a shell, and it seems probable that the loss was due to a reduction like that which took place in the squids, but greater in extent. A drawing of the oldest Octopod known is placed in this Case. Beside it are glass models of living species of Octopoda, among which *Argonauta* deserves mention. As shown by the models and by the exhibited specimens, the well-known Argonaut shell is of a different

![Fig. 87.—Female Argonaut swimming from left to right.](Image)
nature to that of the other cephalopod shells. It is confined to the female and is secreted mainly by her arms, with which she enfolds her body. Their inner surfaces deposit this paper-like shell, which serves as a protection for the brood. A few examples have been found fossil.

We pass now to the **General Collection**, which is divided into the three Orders, Nautiloidea, Ammonoidea, and Coleoidea or Belemnoidea, the smaller specimens as a rule being in the Table-cases and the larger ones in the Wall-cases. The collection is rich in types and figured specimens, of which a list was published by the Trustees in 1898.

**Order—** **Nautiloidea.** Among the uncoiled Palæozoic fossils placed in this Order, there are many which increased knowledge will probably cause us to ally with the Ammonoidea or with the Coleoidea. *Endoceras* and its allies, for instance, are generally admitted to be among these (Fig. 79). Many exhibited specimens come from Sweden and the Baltic provinces, where they are common in a reddish-green limestone, of Ordovician age, which, owing to thin layers of shale, splits readily into flagstones. Thus fine specimens may be seen in the pavements of Swedish towns. A fine example, showing the extraordinary length of the shell, is in a framed slab on the wall by the door. Adjoining *Endoceras* are the obscure but deeply interesting specimens of *Piloceras* from the Durness Limestone of Sutherland and Tremadocian beds of Canada (Fig. 80 a). Other genera of which the truly Nautiloid character may be questioned are *Actinoceras* and *Hurovia*. In *Actinoceras* (Fig. 88) the visceral cone seems to have been constricted by the septa into a series of beadlike swellings. The wall of these was stiffened, and as the siphuncle gradually shrank
away from it towards the centre, the space between the two became filled with calcite. These structures are very peculiar and often puzzling, especially when the inner tubes are found apart from the outer shells; in Huronia from the Silurian rocks of North America they look like fossil backbones, whence one species is named *Huronia vertebalis*. Some of the shells of Actinoceras must have been immense, probably exceeding 10 feet (3 metres) in length, for there is exhibited the body-chamber of *Actinoceras giganteum*, which has a diameter of about 11 inches; and a huge fragment, the whole of which appears to be septate, measures 2 feet 5 inches in length, the diameter of the larger end being about 8 inches, that of the smaller about 4½ inches. This genus, which possibly occurs in the Cambrian, ranges through the Ordovician, Silurian, and Devonian to the Carboniferous Epoch.

Many of these Palaeozoic cephalopods are peculiar in that, after the shell has swollen out somewhat rapidly, perhaps to accommodate the ripening eggs, it again contracts in diameter, and not only this but in some genera the edges of the shell close over the opening of the body-chamber, leaving only narrow apertures. *Poterioceras* is one of the simpler types. The shell is pear-shaped, the earlier part being narrow and divided into chambers perforated by a siphuncle which is somewhat inflated between the septa; the upper and larger portion contained the animal. The whole shell is usually slightly curved, and its aperture not closed in. Fine examples of this genus come from the Carboniferous Limestone of Ireland. It must have attained considerable dimensions, for an example of *Poterioceras cordiforme* from the red Carboniferous Sandstone of Closeburn, Dumfriesshire, is 9 inches long and 7½ inches in its greatest diameter. The shells of the Silurian Gomphoceratide are more egg-shaped and the opening is narrowed to a T-shape by the ingrowth of the margin (Fig. 89). The ends of the crosspiece and upright of the T are widened, and we may suppose that through the lower opening the funnel could eject its stream of water while some arms could emerge through the upper paired openings. Some allied
forms had a larger number of paired openings and could, one supposes, stretch out more arms (Fig. 90). *Ascoceras*, which occurs in the Ordovician of North America and the Silurian of Europe, especially Bohemia, had a curious life-history. This may be followed in the enlarged and diagrammatically coloured model which is exhibited. Beginning with a narrow tubular shell, divided by transverse septa, and having a simple siphuncle near the margin (Fig. 91 *F*), it suddenly swelled out like a *Poterioceras*. This gave more

**Fig. 91.**—*Ascoceratidae. A*, upper part of shell of *Ascoceras manubrium*, cut down the middle, showing the upward-curved septa on the left; *B, C, D*, large curved septa of *Ascoceras fistula*; *E*, upper part of shell of *Ascoceras decipiens*, with septa of ordinary type formed after the deposition of the upward-curved septa; *F*, the shell of the same species completed, showing the simple nautiloid portion (*n*); *G, H*, fragments of an allied form, *Choanoceras*. (From Foord, after Lindström.)

room for the visceral cone with its contained genital glands, and naturally changed the character of the septation (Fig. 91 *A, F*). The body now had so much room that it ceased to advance to any great extent. On the contrary, as the animal grew older its body contracted, and first the opening of the body-chamber narrowed somewhat. Then the visceral hump shrank a little at its end, but more at the side, much as it did in the early stages of *Endoceras*. Thus the septa now produced remained close together at the apex of the body-chamber, and the siphuncle between them was swollen
as in *Poterioceras* (Fig. 91 A). But on one side the septa bent upwards shutting off curved chambers at the side of the shell (Fig. 91 B, C, D). Sometimes the hinder part of the visceral cone shrank a little more and again formed some ordinary septa above these curiously curved ones (Fig. 91 E).

Simpler forms of straight or slightly curved shell, with transverse septa, the siphuncle near the centre, and a plain opening in the body-chamber, are grouped under the name *Orthoceras* (Fig. 80). Shells of this nature are found in all rocks from Cambrian to Trias, but especially in those of Silurian age. Some of them were several feet long. Among the many species exhibited one may note the common *O. ludense* from the Lower Ludlow beds; several from Bohemia, among which is *O. truncatun*, a form that seems to have made a practice of dropping the earlier chambers and sealing up the broken end of the shell with a plug of shelly substance; some polished sections from the Middle Devonian limestones of South Devon; and elongate shells, such as *O. gracile*, from the Lower Devonian of Nassau. Of Devonian age are probably the large specimens of *O. chinense*, known to the Chinese as “Pagoda stones,” from the belief that they are formed underground where the shadow of a pagoda has fallen upon the surface. Polished slabs of rock containing these and other species of Orthoceras are on the wall by the door.

Slightly curved forms of simple Nautiloid type were formerly grouped under the name *Cyrtoceras*, but these are now distributed among several genera of early Palaeozoic age, the name *Cyrtoceras* being restricted to a purely Devonian genus. At an early period appear more closely coiled shells. Several in which the coils or whorls of the shell were scarcely, if at all, in contact, were formerly grouped as *Gyroceras*; but these also are now placed in several distinct genera. In some the earlier coils are close, but the last formed part of the shell is less close, as in the Silurian *Ophioceras* (Fig. 92 a), or even straight, as though unwound, as in the Ordovician *Lituites*. In these two genera the shell aperture is contracted. In *Trochoceras* also the shell is not closely coiled, its special feature, however, is that the coils are not in one plane, but rise in a spire, something like a snail-shell. Beginning in the Cambrian, this genus lasts to Devonian times, but is most abundant in the Silurian rocks of Bohemia, England, and the United States. *Trocholites*, an Ordovician genus from North America, Europe, and India,
has three or four whorls closely coiled in one plane. The Ordovician and Silurian rocks of Bohemia furnish several specimens of a flat closely coiled shell, called *Barrandeoceras*, after the great palaeontologist of Bohemia; an example from Dudley is also shown.

The coiled cone of the preceding shells is generally circular or elliptical in section, and has a smooth or slightly ornamented surface. There are others in which the cone is flattened or grooved, and the surface bears more marked ornament. Thus *Trigonoceras*, found in the Carboniferous

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Fig. 92.—Palaeozoic Nautiloidea. *a*, *Ophidioceras simplex*, Silurian, Bohemia. *b*, *Hercoceras mirum*, Devonian, Bohemia. *c*, *Apheleceras mutabile*, and *d*, *Vestinautilus multicarinatus*, Carboniferous Limestone, Ireland. All slightly less than natural size. (From Foord.)
Limestone of Ireland, has its outer face so flattened that the cone is almost triangular in section. *Hercoceras*, from the Devonian rocks of Bohemia, has a row of projecting tubercles arranged along its length; the shell is not very closely coiled, and its aperture is partly closed (Fig. 92 b). The Carboniferous *Apheleceras* [*Discites*] is flattened and so loosely coiled that there is usually an open space in the centre; ornament is often well marked on the inner whorls, but dies away, and gives place to transverse lines of growth in the later formed part of the shell (Fig. 92 c). The cone is sometimes grooved along its length, but such grooves are more pronounced in its contemporary, *Vestinautilus* (Fig. 92 d). Blunt tubercles are sometimes found on the sides of *Temnocoelites*, which occurs in Devonian, Carboniferous, and Triassic rocks, and has a flattened outer face with angular sides. *Pleuronauutilus*, which lasts from Devonian to Triassic times, is remarkable for its tubercles and transverse ribs, more like the ornament of an ammonite.

The Carboniferous *Asymptoceras* has a smooth cone, widening rapidly, and with the margin of the aperture drawn out into a narrow projection on each side; its septa are simple as in *Nautilus*, but the siphuncle is near the margin. In *Subclymenia*, on the other hand, the sutures are so folded as to resemble those of some goniatites.

From the Triassic Limestones of the Austrian Alps are shown a number of shells, generally resembling that of the recent *Nautilus*, but separated by various minor differences. Even more like the modern *Nautilus* are the shells of Jurassic age. Among these is a specimen of *Nautilus clausus* from the Bajocian rocks, near Caen, showing impressions of the shell-muscles. A large *Nautilus ornatus* from the Inferior Oolite of Sherborne, Dorset, shows the rapidly widening cone, the lines of growth, and longitudinal grooves. The Gault furnishes several nautili in which the shell is well preserved, and in a young shell of *Nautilus semiundatus* from the Upper Greensand of Warminster, the iridescence of the mother-o'-pearl septa may still be seen. This last species often attains a great size, as shown by specimens in the Wall-case. From the Upper Cretaceous Limestone of Lebanon comes *Nautilus libanoticus*, in two specimens of which the jaw-beaks may be seen. The nautili of the Chalk are often large, but the internal casts are usually all that is found.

Although it used to be said that the genus *Nautilus* persisted without change from early Palaeozoic times to the
present day, closer study has disclosed many differences between the Palaeozoic shells and those of the modern nautilus. Even the Mesozoic shells are not really the same, and not until Tertiary times do shells occur that can with more justice be referred to _Nautilus_ as now restricted. In the London Clay many of these are well preserved and show the internal structures; among them _N._ _imperialis_ often attains great size. _Aturia_, of which specimens from Eocene, Oligocene, and Miocene rocks are shown, differs from _Nautilus_ in the folding of its sutures, well seen in _A_. _ziezae_.

Further information concerning the fossil Nautiloidea in the Museum is given in the first two volumes of the "Catalogue of the Fossil Cephalopoda" issued by the Trustees in 1888–91.

Order—_AMMONOIDEA_. One of the earliest straight-shelled forms that can without doubt be referred to this Order is the Devonian _Badrites_ (Fig. 81 _b_), in which the septa are still unfolded, but which has a protoconch (see model, Table-case 1) and its siphuncle marginal, i.e., near the outer shell-wall. We have already seen, in such a form as _Mimoceras compressum_ (Fig. 81 _a_), how the straight shell became coiled first in its old age, and how in more advanced forms the coiling began at an earlier and earlier stage of the life-history, until even the protoconch was affected by it. In many of the earlier Ammonoidea the protoconch can still be seen distinctly (Fig. 93 _a_), being uncovered by later whorls,

![Fig. 93.—Goniatites. _a_, Pronorites cylindrobus, and _b_, Glyphioceras sphae-ricum, Carboniferous Limestone, England. _c_, Agathiceras Suessi, with shell preserved, Permo-Carboniferous, Sicily. Natural size. (From Foord and Crick.)](image)

and its globular shape is apparent. In others the shell soon became more tightly coiled (Fig. 93 _b_), till the protoconch is hidden by subsequent whorls (Fig. 93 _c_). As the shell
became coiled, its septa were thrown into folds, but these were generally of a relatively simple character with an angular suture (Fig. 93 b), whence these forms are collectively known as Goniatites. They are usually smooth, or with only fine lines of growth, rarely with tubercles or ribs. The goniatites are mostly of Devonian and Carboniferous age, but also occur in the Permian, after which they give place to the Ammonites. Among the genera here exhibited may be noticed the above-mentioned Mimoceras, and Agoniatites fecundus also showing the uncoiled initial portion. Most of the Devonian goniatites are from Germany, but some species have also been found in this country. Thus there is a Tornoceras from the Middle Devonian of Devonshire, and Gephyroceras intumescens from the Upper Devonian of that county. Among the Upper Devonian specimens from Germany are several aptychi. Some Devonian genera form a special group, distinguished by having the siphuncle on the inner side of the whorl. The best known of these is Clymenia, and a thin section here exhibited shows this character plainly. Greater complication is noticeable in several of the Carboniferous goniatites. Here, for instance, are the closely coiled Glyphioceras (Fig. 93 b) and Brancoceras, Gastrioceras with tubercles, Pericyclus with transverse ribs, and Prolecanites and Pronovites with many-lobed sutures (Fig 93 a). Near the last is the tiny Dimorphoceras discerpanis. The goniatites are fully dealt with in Vol. III. of the Catalogue of Fossil Cephalopoda (1897).

The transition from Goniatites to Ammonites took place gradually along many lines, which are being worked out by a study of the Triassic species. This study is based mainly on the complications of the suture, which are far too intricate to receive further explanation in this place. Broadly speaking, it is possible to trace lines of descent with some exactness, owing to the fact that each individual shell in its early stages still possesses the structure that characterised its adult ancestors. This is indeed only one instance of a general principle of growth affecting most living beings; but the principle can be more easily applied in the study of these coiled shells, since the early stages are always preserved and can often be clearly seen. We have, for example, already noticed how the straight Bactrites stage is repeated in Mimoceras. It is often found that two species closely resembling one another in adult stages differ so greatly in their earlier stages as to lead to the conclusion that they
Gallery VII. have descended from totally different ancestors, and they may thus be placed in distinct genera, perhaps in distinct families. Unfortunately this method of study has not yet been pursued long enough for investigators to have settled the numerous problems presented by the very large numbers of these fossils; nor are the solutions that have hitherto been published always found to agree. Hence the classification and nomenclature of the ammonites has for some years been in a state of transition, to the great perplexity of geologists who wish to use these widely distributed fossils for the identification of various strata and to quote names long familiar in that connection but now requiring emendation. It is impossible to alter the arrangement and naming of a great exhibited series like that of the British Museum to accord with a rapidly advancing classification. Those important questions must be studied in original memoirs, and these pages can only mention a few of the more conspicuous specimens.

The smaller ammonites in the Table-cases are arranged for the most part under the geological Ages, the foreign specimens of each Age being placed in the same Case as the British ones. The larger specimens in the Wall-cases begin with Liassic forms in Cases 12 and 11, pass to Bajocian in 10, and Oxfordian in 9; then, crossing the Gallery, continue with the uppermost Jurassic in Case 6, Lower Cretaceous in 5, and Upper Cretaceous species in 4 and 3. We begin with the oldest.

From the Himalayan Trias are shown Ptychites, Carnites, and Gymnites, smooth shells, with sutures not far removed from those of goniatites. Among many specimens from the Upper Trias of Hallstadt in Austria, Monophyllites and Rha-ycoplyphites, which have primitive sutures with leaf-like saddles, start the line of Phylloceratidae. In Pinacoceras Metternichii, on the other hand, the sutures have already acquired an extraordinary complexity, as best shown in some large specimens and a section in Wall-case 12. Sutures of rather simpler type are clearly shown in specimens of Cladiscites multilobatus. From St. Cassian in the Tyrol come the roughly ridged shells of
Ammonite Marble.
Marston Stone, from the Lower Lias near Yeovil, full of Amblycoceras planicosta.
The same Rock cut and polished.

[To face p. 107.]
Trachyceras Aon (Fig. 95 a), marking a well-known horizon rich in fossils. Ceratites with its many rounded saddles, recalling Prolecaneis, is also characteristic of the Trias (Fig. 94); a section of C. nodosus shows how the septa are at first concave as in older forms and then become convex towards the shell-aperture, as in ammonites generally. In Arcestes intuslabiatus the periodical constrictions of the coiled cone provoke enquiry as to their cause.

Passing to the Lias, one notes many type-specimens of species founded by the Sowerbys and by T. Wright, as well as examples of the ammonites that give their names to the successive horizons or zones into which the Lias has been divided. Of these the oldest is Psiloceras planorbis, the earliest English ammonite; beginning with slightly ribbed whorls, it reverts in the adult to the smooth shell of the older types. A fine slab covered with iridescent shells of this ammonite is placed on the wall. Among the British specimens from the Lower Lias, an example of Acroeceras heterogenes [C 1870] shows ribbed inner whorls like those of the adjoining A. capricornus, and outer whorls with tubercles, as in Liparoceras striatum. Such changes from smooth to ribbed, from ribbed to tuberculate, characterise many ascending lines of ammonite evolution. Here are to be seen shells of Amblyoceras planicosta, of which a thick bed was formerly worked as an ornamental marble at Marston near Yeovil. Slabs of this showing weathered and polished surfaces are placed on the adjoining wall (Plate VII). The larger specimens from the Lower Lias include a very fine example of Coroniceras Bucklandi showing the coronet of blunt spines from which the genus takes its name. Near this is a large Asteroeceras stellare from Lyme Regis, cut in half and showing the chambers dislocated during fossilisation. Asteroeceras obtusum shows the keel and the simple suture contrasting with the rather complex one of Coroniceras. Here is a Deroceras armatum with its big spines. Above are some large specimens of the rare Vermiceras Conybeari. Between the cases is the largest known Lias ammonite, about 1 metre (3 feet 4 inches) in diameter, possibly an old individual of the last species. Among species from the Middle Lias a noteworthy one is Lytoceras fimbriatum, with sharp ridges at intervals indicating that from time to time the aperture of the shell flared outwards, for reasons at which we can only guess (Fig. 95 e); these flares cut across the ordinary fine ribs of the shell; in
**Gallery VII.**

*Lytoceras lineatum* this habit was again lost, especially in old age, and only the fine ribs are seen. The two species *Amaltheus margaritatus* and *Patellopyrocceras [Amaltheus] spinatum* should be noted for their peculiar ornament, and because they are characteristic of the Marlstone.

Among specimens from the Upper Lias the black ammonites from Whitby catch the eye. They include the

![Ammonites](image)

**Fig. 95.—Ammonites.** *a*, *Trachyceras Aon*, Upper Trias, side view. *b*, the same, front view, showing folded septum. *c*, *Lytoceras fimbriatum*, Middle Lias. *d, e*, *Coeloceras Blagdeni*, Inferior Oolite. *f*, *Phylloceras heterophyllum*, Upper Lias, shell partly removed to expose sutures. Less than natural size.

**Table-case 7.**

**Wall-case 11A.**

Primitive form *Cymbites carinatus*. Close by is the type-specimen of *Harpoceras falciferum*. A specimen of this species in the Wall-case has an aptychus in the body-chamber. Here, too, are large specimens of *Phylloceras heterophyllum* in section, as well as polished ones showing the foliated sutures (Fig. 95f).
The Oolitic series includes *Macrocephalites* from India, and various genera from the Bajocian rocks of Calvados and the corresponding Inferior Oolite of this country. A specimen referred to *Stepheoceras Brodiquei* has a furrow down one side of the last whorl, the result apparently of injury to the secreting surface. *Cadomoceras [Oekotraustes] cadomense* is eccentrically coiled, and has projecting ears at the sides of the shell-aperture. *Haplopleuroceras subspinatum* is remarkable as repeating the general form of *Paltopleuroceras spinatum* in the Middle Lias; this is one of those similarities of form in different genera that has led to so much confusion in classification. Here we may note a fine aptychus from the Stonesfield Slate, and a beautiful *Stepheoceras* from the Great Oolite of Minchinhampton. The ensuing Cornbrash furnishes good specimens of *Macrocephalites*. The larger ammonites from the Lower Oolites include fine examples of species giving their name to well-known horizons, such as the type-specimens of *Ludwigia Murchisonae* and *Stepheoceras Humphriesianum*, also *Parkinsonia garantiana*, *Strigoceras Trudiei*, *Coeloceras [Stephanoceras] Blagdeni* (Fig. 95 d), and *Parkinsonia Parkinsoni*. There are some very large specimens of *Parkinsonia dorsetensis* and others, in section, or polished to show the sutures. The type-specimen of *Fontannesia Boweri* from the Inferior Oolite of Bradford Abbas preserves its long ears, and here is also the type-specimen of "*Ammonites* Tessonianus", d'Orbigny, from the Tesson collection. Among Oxfordian species the well-known *Cardioceras cordatum* repeats the general form of *Amaltheus margaritatus*, although derived from quite a different stock. Very fine series of this genus and of *Cadoceas* are in the Wall-case. The Callovian series furnishes several characteristic genera, among them the highly ornamented *Cosmoceras*. Many Oxfordian species of this genus have very long processes or ears from the sides of the aperture. From the Oxford Clay of Wiltshire come large specimens of these as well as of *Proplanulites Koenigi* with a smooth senile stage.

The Coral Rag of England is noted for the internal casts of the chambers of ammonites, from which the shell has been dissolved away. These belong mostly to *Aspidoceras perarmatum*, of which some specimens from Brora, in Sutherland, are also shown. Numerous aptychi are found in the Kimmeridge Clay, especially at Ely, and the same forms occur in connection with their ammonites in the contemporaneous Solenhofen Stone. Near these are specimens
of *Holocostephanus* from the Portlandian near Moscow, showing the iridescent shell. The large *Holocostephanus gigas*, and the still larger *Perispheles giganteus*, from the Portland Stone of England are in the Wall-case. A small specimen of the latter species had the shell turned into silica, and the infilling limestone has been dissolved out, exposing the shape of the septa and the position of the siphuncle.

The Lower Cretaceous rocks have not furnished many ammonites, but the large "*Hoplites*" *Deshayesi* from the Lower Greensand carries on the general plan of the genera just mentioned. Other new genera appear in the series from the Albian of Eseragnolles (Var.), and from the coeval Gault of Folkestone. The iridescent appearance of these and other ammonites previously noticed is due to the solution of the outer layers of the shell, by which the inner nacreous layers are exposed. Here we meet with *Hoplites*, characterised by a broad groove on the outer margin of its shell, similar to that previously seen in the otherwise unlike *Schlotheimia* of the Lower Lias and *Parkinsonia* of the Middle Oolites. The specimens of this genus from Folkestone form a series illustrating the decline of ornament from tuberculate, through ribbed, to smooth. Similarly *Schloenbachia* shows a decline from highly tuberculate to ribbed. *S. rostrata* marks a wide-spread horizon in the Albian, and there is a splendid series of it from both Gault and Upper Greensand, preserving the long rostrum at the shell-aperture. *Phylloceras Guettardi* and several species of *Holcodiscus* show periodical constrictions of the shell-aperture. The specimens from the Cambridge Greensand are derived from the underlying Gault, and those from the Red Chalk seen in the next case are also of that age.

The Cenomanian forms in the Table-case include the characteristic *Acanthoceras rotomagens* from Rouen, *Brahmaites* from Pondicherry, and *Pachydiscus* from various European localities. The Upper Cretaceous rocks of South Dakota yield specimens of *Placenticeras* with elaborate sutures. *Tissotia*, on the other hand, from the Senonian of Algeria, shows that return to a Ceratite form of suture which is found in many late Cretaceous ammonites. Among the British specimens from the Chalk Marl and Chalk are many figured in D. Sharpe’s monograph published by the Palaeontographical Society. Those from the Chalk Marl include a fine series of *Schloenbachia* ranging from the tuberculate *Schloenbachia Coupei*, through the more or less ribbed
H. varians, to the smooth H. Goupilianus. This again illustrates the change of ornament characteristic of a declining series. Among the larger specimens the most noteworthy are Puzosia Austeni, Pachydiscus peramplus, and Pachydiscus leptophyllus. Some specimens clearly show the complex suture of this last, but the most interesting is the very large one from Rottingdean, 3 feet 8 inches in diameter; the rapid increase in the width of the coiled cone may be contrasted with the very slow increase in the large Lias specimen opposite. The largest known ammonite is P. seppenradensis, from the Lower Senonian of Westphalia, with a diameter of about 2 metres (6 ft. 8 in.). A plaster reproduction of it is fixed at the north end of the Gallery.

We have already noticed the changes in ornament that characterise ammonite races as they advance to and recede from the acme of their development, and we have seen how the suture likewise becomes more complex and then returns to a simpler form. There is yet another and more obvious change. Just as the ascending stocks, beginning with straight forms, gradually coiled the shell more and more closely, so, having reached their acme, they begin to uncoil and may ultimately return to a straight condition, if they do not previously become extinct. Further, instead of merely unwinding, they may lose the regularity of the coil and become wound in an asymmetrical spire or turret, like that of most gastropod shells. Already in Triassic times the Ceratites (using the term in a broad sense) show all these retrogressive changes, ending in the straight Rhabdoceras. Of the various ammonite families that passed into Jurassic and Cretaceous times, the Stepheceratidae gave off a degenerate branch so early as the Bajocian Age. The eccentric Cadomoceras has already been noticed, and here are exhibited the further uncoiled Spiroceras [Crioceras] bifurcatum and similar forms, which led to the straight Baculina of the Cretaceous.

It was, however, chiefly towards the close of the Cretaceous Epoch that all the persisting races entered on this degeneration. Names have been given to the various stages of uncoiling, such as Crioceras, in which the whorls are partly separate (Fig. 96 a); Macroscaphites with the last whorl bent slightly back and then returning on itself (Fig. 96 d); Scaphites with a somewhat closer coil to start with and a more rapid return (Fig. 96 f); Hamites, which starts with a small coil, then goes straight for some distance,
Fig. 96.—Uncoiled and Asymmetric Cretaceous Ammonoidea.  

a, Crioceras Emerici, Neocomian.  
b, Heteroceras Emericianum, Neocomian.  
c, Turrilites catenatus, Albian.  
d, Macroscaphites Ivanii, Neocomian.  
e, Hamites elegans, Albian.  
f, Scaphites Hugoianus, Albian.  
g, Baculites anceps, Danian.  

All less than natural size.
then returns for a longer distance, and ends straight after yet another bend (Fig. 96 e); and Baculites, in which, after a small initial coil, the shell continues straight to its aperture right up to death (Fig. 96 g). By study of the ornament and suture, it has been found that these forms represent stages of degeneration common to more than one race, and therefore do not constitute true genera. Thus many of the Neocomian shells in the Crioceran stage are ribbed in the irregular manner characteristic of Lytoceras; see, for example, C. villersianum in the Table-case. Those named Pictetia Asticrana also show the peculiar Lytoceran suture. In many of those from the Speeton Clay, on the other hand, e.g. C. quadratum, the ribbing is very irregular, and appears to have been derived from such a form as Perisphinctes or Holocostephanus. The same appears to be the case with the large Aptian Crioceras Bowerbanki and its derivative Macroscaphites grandis, the earlier whorls of which resemble those of the Portlandian Holocostephanus gigas, while the later portion presents a remarkable exaggeration of certain ribs. In these massive shells may be seen a curious retention by the uncoiled portion of a character originally due to close coiling. The cone in the earlier Ammonoidea is circular or elliptical in section, but as they become coiled the inner side of the outer whorls is impressed or excavate, so as to fit closely over the inner whorls (Fig. 95 b); the closer the coil, the greater is the depth of the impressed zone. This impressed zone is clearly seen on the later whorls of many of these uncoiled forms and is clear evidence of their descent from more closely coiled ancestors. Eventually it disappears, and the long loops of a Hamite, for example, show no trace of it. In many shells called Baculites, the folded sutures alone distinguish the greater part of the shell from a smooth Devonian Bactrites; but even the sutures, as may clearly be seen in the specimens from the north of France, are much simpler than in most of the coiled or partly coiled ammonoids. The shell-aperture is oblique, suggesting that the animal had given up swimming for crawling.

The history of the turreted or helicoid shells is of the same general character. There is, however, reason to suppose that the tendency to this asymmetry usually arose at quite an early period in the life-history instead of coming in with old age as did the uncoiling. Something of the kind was noticed in Palaeozoic nautiloids, and it may be supposed that the tendency was always present, especially in weaker
individuals, but was kept in check by the close coiling. Turreted forms appeared first in the Upper Trias (e.g. Cochloceras), but it was in the Cretaceous Epoch that they were first a large element in the fauna. Turrilites (Fig. 96 c) in the strict sense and various genera of similar form appear to be derived from such antecedent forms as Cosmoceras and Dowilleiceras, which they resemble in ornament. In some of these the whorls are closer than in others. Helicoeceras and Heteroceras (Fig 96 b) begin as asymmetrical spirals, but turn off in a different direction in old age. They and some of the Hamites are supposed to be connected with Acanthoceras. The direction of the turretted coil varies: in the Senonian Heteroceras polyplecum from Westphalia it is dextral, as in most gastropods; in the Turrilites of the Chalk Marl it is generally reversed or sinistral; in those of the Gault it is indifferently dextral or sinistral in the same species.

The various changes in these Cretaceous Ammonoidea may be described as retrogressive, for they are in some respects a going back along the line followed in the previous evolution of the Order. They were followed by complete extinction, for the Order did not persist into the Cainozoic Era. Were these changes in accord with changes in the environment, and was the extinction of the Order due to inability to keep pace with change of conditions? Or were the changes inherent in the constitution of the ammonites, a necessary result of their previous history, and do they signify a true degeneration and decline, out of all accord with the surroundings? One fact not yet mentioned may have a bearing on this problem. It is that in some shells the last bend grows in such a direction that in old age its aperture was brought up against a preceding part of the shell, so that the arms of the animal can scarcely have emerged; by continuing its own growth, it seems that the individual killed itself. Did the race do the same?

Order COLEOIDEA or BELEMNOIDEA.—In modern times this has taken the place formerly occupied by the Ammonoidea and before that by the Nautiloidea. Whatever may be the affinities of certain straight-shelled Palæozoic cephalopods, the earliest fossils that show undoubted traces of the enclosing mantle are Aulacoceras and Atractites of the Upper Trias. These have guards, but the phragmocone is relatively large, and in the former retains traces of longitudinal ornament.
In the Lower Lias of Dorset are still found very large phragmocones, but other specimens show the Belemnite type fully developed, with pro-ostracum, ink-bag, and hooked arms. Of *Belemnites elongatus* there is the fine specimen described by Huxley in the Monographs of the Geological Survey and Sowerby’s original specimen from Crick tunnel near Daventry. The Middle Lias of Charmouth has yielded the slender *Xiphoteuthis elongata*, also described by Huxley. From the Upper Lias of Alderton, Gloucestershire, comes a well-preserved pro-ostracum. A monster phragmcone of the Bajocian species *Belemnites giganteus* comes from Germany. The Oxfordian of Trowbridge and Christian Malford in Wiltshire furnishes a large series of *B. Owenii* var. *Puzosianus*.

Among Cretaceous belemmites, *Duvalia dilatata* is remarkable for its guard, swollen in one direction and flattened in the other. *Actinocamax* is the usual form from Cenomanian to Senonian, being joined by the similar *Belemnitella* in the latter Age.

The belemmites did not die out at the close of the Cretaceous Epoch, but they changed in character. *Styrcwo-teuthis orientalis* from the Eocene of Syria is still of the older type, but in most the guard was reduced in length, thickness, and calcification. *Vasseuria* from French Eocene rocks has such a slender yet relatively short guard. In *Beloptera* and *Belopterina* the guard is short and somewhat swollen at its end, which makes a slight angle with the phragmcone; in the former it expands at the sides into two wings. The latter genus is not far removed from the Miocene *Spirulirostra*, already described (Fig. 85 b). In a later genus *Spirulirostrina* (not exhibited) the guard is more reduced, and in the modern *Spirula* it has disappeared (Fig. 85 c).

Another line of evolution leads, as previously explained, from *Beloptera* to *Belosepya* (Fig. 85 d) of which many specimens from the London Clay and Bracklesham Beds, are shown. *Sepia* itself is exhibited from later Tertiary rocks (Fig. 85 e).

Of those sheathed forms in which the calcification of the shell underwent a gradual reduction, the earliest known is *Phragmotethis* from the Upper Trias. The next in age is *Geoteuthis*, of which large specimens from the Lower Lias of Dorset are exhibited. These and the smaller specimens from the Upper Lias of Württemberg and Normandy show an expanded pro-ostracum, divided lengthwise into three areas,
and with no trace of a phragmocone. Many of these preserve the ink-bag, which in *G. brevipinnis* from the Oxford Clay of Christian Malford is sometimes of great size, while the ten arms are very short—an obvious correlation. From the Upper Liassic are also shown the similar shell of *Teuthopsis*, and that of *Beloteuthis* strengthened by a median keel. Shells of generally similar character are found in *Coccolithus* [Trachyteuthis] from the Solenhofen Stone. An admirable specimen preserves portions of the mantle and side-fins, and has eight well-developed arms bearing suckers; the two long arms found in recent Decapoda may have been present but retracted. A very large shell of this genus is at the bottom of the Case. In *Plesioteuthis prisca* from the same stratum the shell is reduced to a long narrow pen, with the side expansions at its hinder end and quite small. The same genus occurs in the Senonian rocks of the Lebanon, whence come *Plesioteuthis Fraasi* and the allied *Dorateuthis syriaca* (Fig. 86 b), both with eight short and possibly two long arms.

A shell in which the phragmocone is still preserved, as in *Phragmoloteuthis*, but in which the guard is reduced to a thin shiny coat, is that of *Belonephoteuthis antiqua* (Fig. 86 a). There is shown a fine series of this from the Oxford Clay, chiefly of Christian Malford (see p. 156). The ten short arms are well seen, and in one specimen seem to have caught a small fish. *Acanthoteuthis* from the Solenhofen Stone is said to have had a shell with more reduced phragmocone and larger pro-ostracum. Specimens are shown preserving the arms, eight or ten in number, with well-marked hooks; the mantle; and a membrane round the mouth like that of living Onychoteuthidae. *Conoteuthis* (Fig. 85 g), of which fossils are shown from Neocomian, Aptian, and Albian rocks, had a small curved phragmocone, suggesting the end of the *Ommastrephes* shell (Fig. 85 h).

The specimen of *Palacocotopus Newboldi*, from the Senonian of Lebanon, is the oldest fossil Octopod. There is no evidence to show from which of the races just described it may have been derived. With the Octopoda, which are the most highly specialised of Mollusca, and furnish some of the monsters of modern seas, we reach the end of this sketch of extinct invertebrate animals.
INDEX.

Lepidocerinus, (fig.) 64, 65
Leperditia, 93
Lepidaster, 68
Lepidocentrus, 72
Lepidoceras, 95
Lepidodesmus, 67
Lepidoptera, 105
Leptanaea, 112
Leptobiattina, 105
Leucandra, 52
Libellula, 105, 107
Lima, 136, 137
Limnaca, 19
Limpet, 124, 136, 143
Limulus, 88, (fig.) 89
Lingula, 109, (fig.) 110, 111, 112, 113
Lingula Flags, 111
Lingulella, 111
Liopista, 133
Liothyria, 115
Liparoceras, 167
Lister, J. J., 24, 25
Lithistida, 36
Lithodomus, 12, 133, 135, 137
Lithonanta, 105, (fig.) 107
Lithophylyx, 12
Lithosialis, 105, (fig.) 106
Lithostroton, 54
Lithothrochus, 145
Littorina, 127, 135
Lituites, 101
Lituola, 19, (fig.) 21, 23
Lituolidae, 19
Lobsters, 100
Locust, 107
Lodanella, 39
Lognogaster, 47
Loligo, (fig.) 145, 157
Lonsdalea, 54
Loricata, 99
Loricula, (fig.) 94, 95
Loxonea, 139
Lucina, 134
Ludwigia, 169
Lumbriecaria, 79, (fig.) 60
Lyctus, J., 124, 129
Lyelli, C., 125
Lyopora, 51
Lyssacia, 53
Lytocera, 167, (fig.) 168
MACULARIA, (fig.) 140, 141
Macoma, 128
Macrocephalites, 169
Macrocheilus, 139
Macrodon, 136
Macroscaphites, 171, (fig.) 172, 173
Macura, 98
Madreporaria, 49
Maeandrospiculidae, 35
Magellania, 116, (fig.) 114
Magnesian Limestone, 13
Malia, 102
Malacostraca, 95
Mantell, G. A., 129, 131, 156
Marginalia, 20, (fig.) 21
Markings, 14
Marsupites, 63
Mastigocerinus, 61
Mastoidea, (fig.) 37
Mecochirrus, (fig.) 94, 99
Megalomnus, 142
Megaronia, 37
Melania, 129
Melisnerita, 122
Mellitionidae, 35
Meloceras, (fig.) 150
Melonites, 72
Melongena, 119, (fig.) 129, 121, 122
Meretrix, 131
Microstomata, 81, 86
Mera, 100
Micracia, 57
Micraster, 75
Miliola, (fig.) 16, 23
Miliolidae, 19
Millipedes, 103
Millepora, 47
Millipora, 47
Miliocerinus, 63
Mimoceras, (fig.) 150, 164, 165
Minchin, E. A., 42
Mithracites, 102
Mitrobylites, 65
Modiolia, 129, 131, 135, 136, 137
Mollusca, 132
Monactinellida, 38
Monilopora, 54
Monophyllites, 166
Monopleure, 143
Monetis, 139
Monticulipora, 147
Montlivaltia, (fig.) 55
Mopsida, 57
Mopsisa, J., 134
Moseleya, 55
Mountain building, 12
Mourlonia, 139
Mucrophora, 19
Murchisonia, 139
Mya, 128
Mytilus, 134, 135
NAMEs, 2, 3
Nassellaria, 28
Natricula, 151, 154, 136
Naticopsis, 139
Nautiloidea, 151, 152, 158
Nautilus, (fig.) 145, 149, (fig.) 150, (fig.) 151, 158, 164
Neocarcinus, 102
Necrocallis, 97
Netheia, 131, (fig.) 132
Neolatomus, 89
Neotremata, 113
Neptunea, 127, 129
Nereitopsis, 79
Neritonus, 136
Nerinea, 135, 136
Nerita, 136
Neritina, 123, 128
Neuroptera, 105
Newton, R. B., 11
Nodosaria, 29, (fig.) 21, 23
Nodosinella, 19
Nozizonia, 20
Notamina, 151
Nucula, 135
Nuculana, 139
Nummulites, 29, 23, (fig.) 24
Nummulitidae, 99
OBOLUS, 111, 112
Octactinellida, 38
Octopoda, 157, 176
Octopus, 145
Ocydina, 57
Oekotraustes, 169
Ogygia, 85, 86
Olenellus, 85, 86
Olenus, (fig.) 85, 86
Ommastrephes, (fig.) 155, 176
Omphrya, 52, (fig.) 53
Oncobrochus, 57
Onustus, 136
Onychaster, 70
Onychocella, 119, (fig.) 129, 121, 122
Opereculina, 20
Ophioceratina, 161, (fig.) 162
Ophioglypha, 71
Ophiolepis, 71
Ophurella, 71
Ophiuroida, 70
Ophiolithbranchia, 125, 141
Ophiopathea, 55
Orbiculoides, 106, (fig.) 112
Orbigny, A. D. de, 18
Orbitoides, 20, 23
Orbitolites, 23
Orbitolipora, 121
Ortoceras, 102
Oreithyas, 115
Orthoceras, 151, (fig.) 148, 161
Orthohegra, 140
Orthoptera, 105
Ortonia, (fig.) 78
Ostacoda, 93
Ostrea, 129, 131, 134, 135
Owen, K., 146
Oxytoma, 102
Oxytoma, 136, 137
PACHASTRELLA, (fig.) 36, 41
Pachydiscus, 170, 171
Pachyphora, 53
Pachypoteron, 49
Pachyrhisma, 136
Palaeacis, 54
Palaeaster, 68
Palaeasteriscus, 69
Palaeochinus, 72, (fig.) 73
Palaeogyna, (fig.) 97
Palaeomachia, (fig.) 101
Palaeonidae, 99
Palaeocaris, 96
Palaeocystes, (fig.) 94, 97
102
Palaeocragon, 96
Palaeocystus, 176
Palaeocyculus, 58
Palaeodiscus, 72
Palaeonoman, 41
Palaeophonus, 89, (fig.) 90
Palaeoeternus, 105
Palasterina, (fig.) 69
Palinuridae, 90
Palinurna, 100
Palinurus, 99
Paltopleuroceras, 168, 169
Paludestrina, (fig.) 128
Paradoxides, 85
Parallelodon, 136
Paramoudras, 13
Paramusila, (fig.) 56
Parker, W. K., 18
Parkeria, 48
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